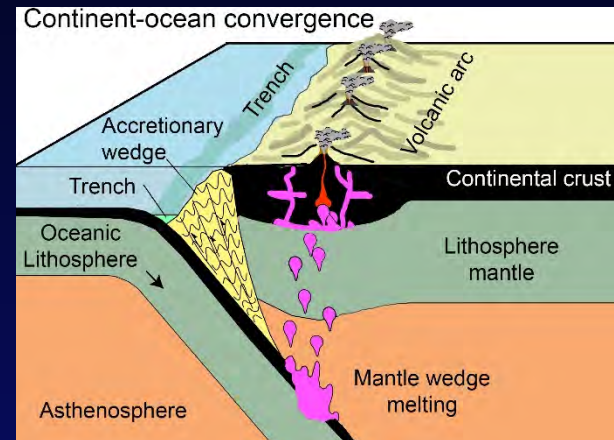


Geophysical constraints on geodynamic processes at convergent margins: A global perspective



Gondwana
Research,
2015

Irina Artemieva

Hans Thybo

Alexey Shulgin

Univ. Copenhagen, Denmark

Univ. Copenhagen, Denmark

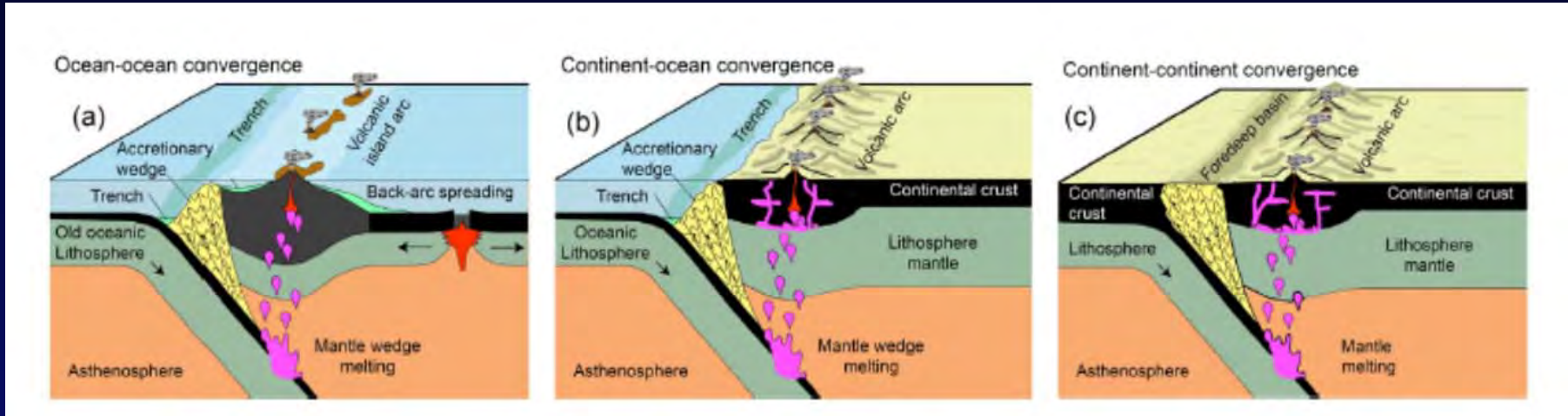
CEED, Univ. Oslo, Norway

Three types of convergent margins

O-O

C-O

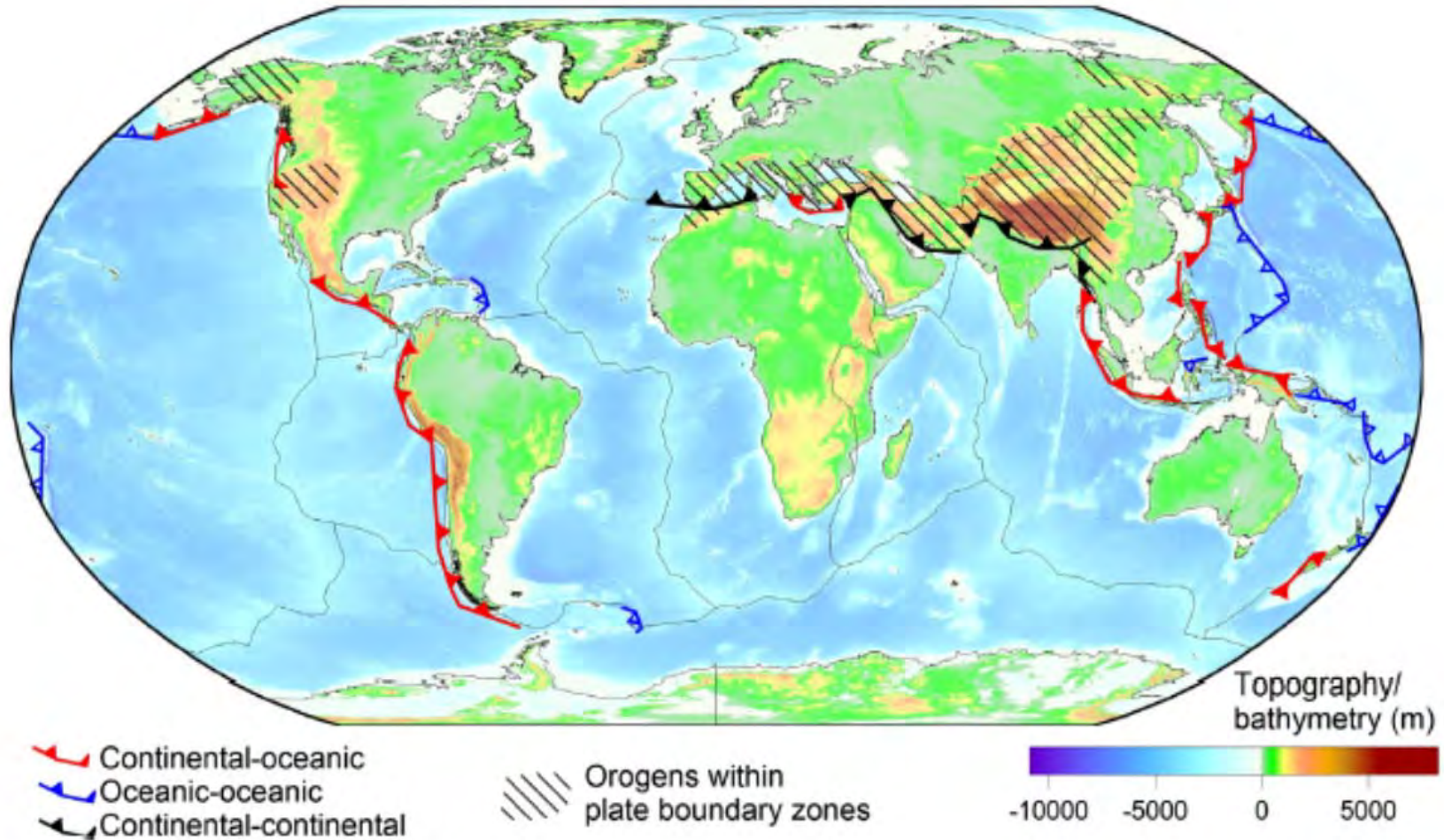
C-C



Incoming Oceanic

Overriding Continental

Three types of convergent margins



What happens at convergent margins and what are the controls?

"INPUT":

- **Structure of incoming plate**
(crust: Moho, seds, topography,
mantle: water, LAB, temperature
→ age for ocean plates...)
- **Structure of overriding plate**
 - (LAB, temperature,...)
- **Convergence rate**

What happens at convergent margins and what are the controls?

"OUTPUT":

- **Subduction angle**
may vary with depth
- **Roll-back**
-> back-arc basins
- **Seismicity**
Magnitude
Depth distribution
- **Volcanism**

What happens at convergent margins and what are the controls?

OBSERVABLES:

- Gravity (FA & B)
- Heat flow
- Seismic V_p , V_s , Q
- Seismicity & volcanism

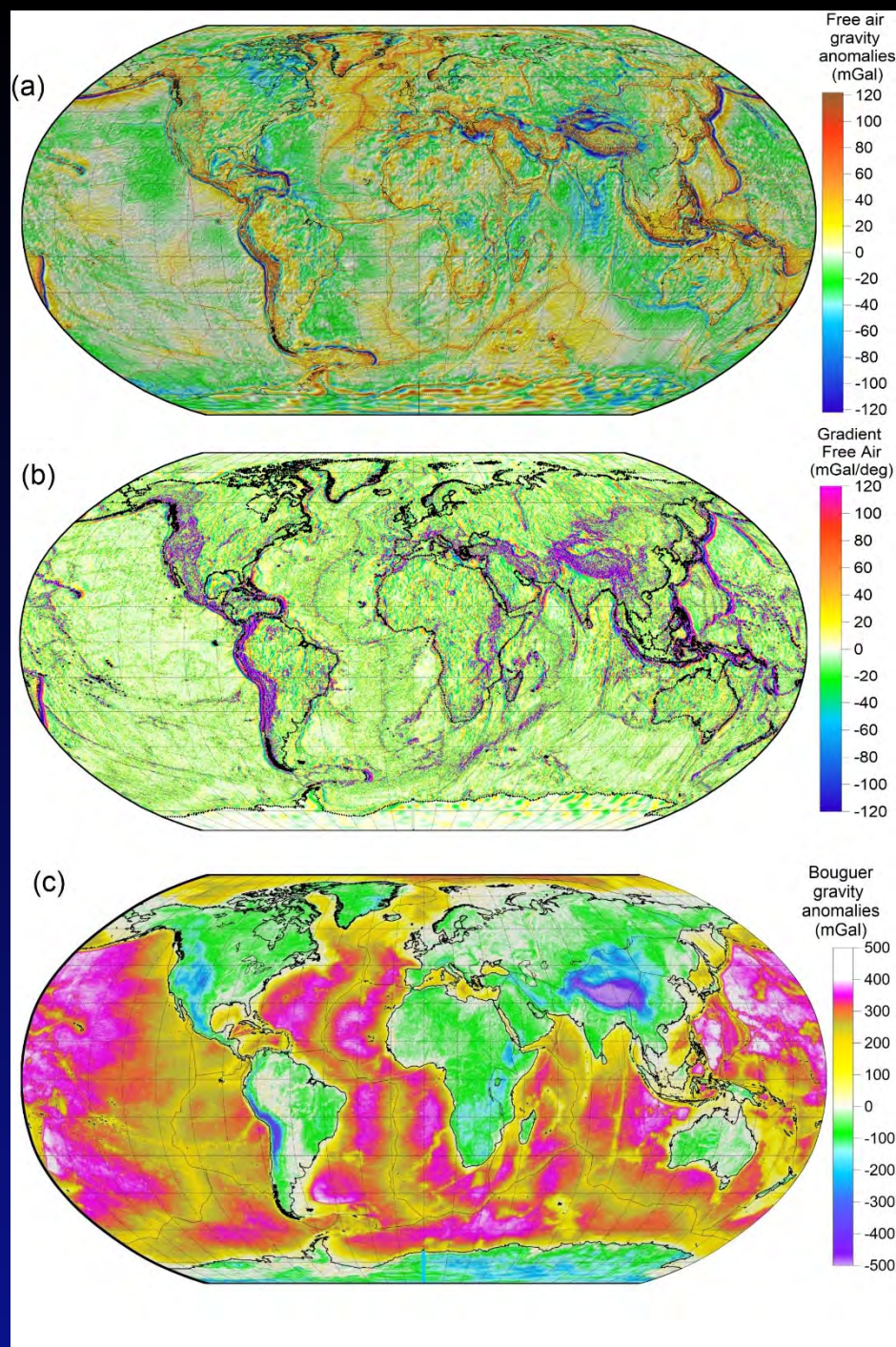
Gravity

Free Air

- Globally - proximity to isostasy
- No local isostasy at convergent margins

Extreme contrast in grad FA

Bouguer



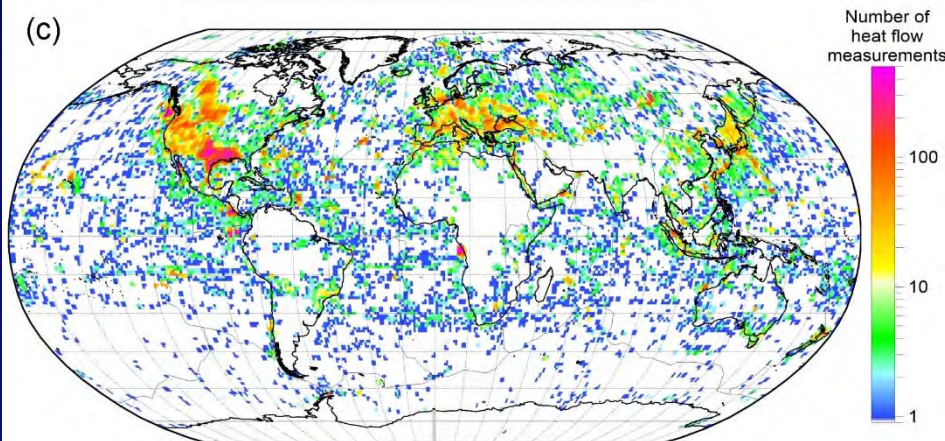
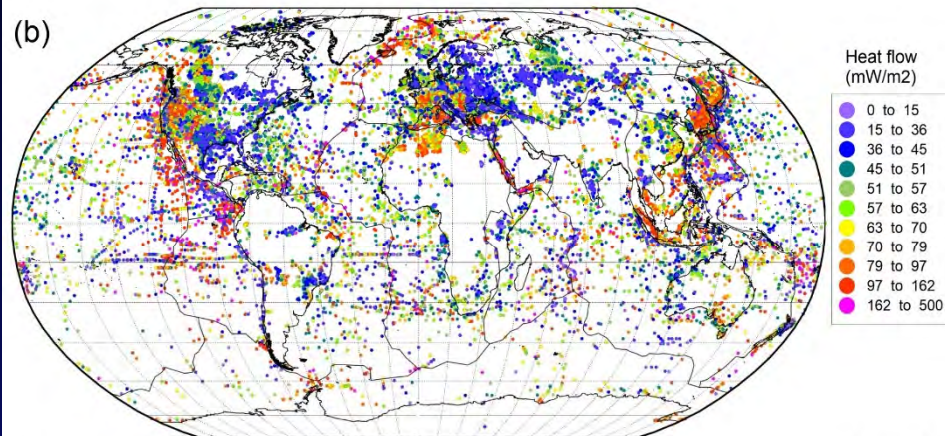
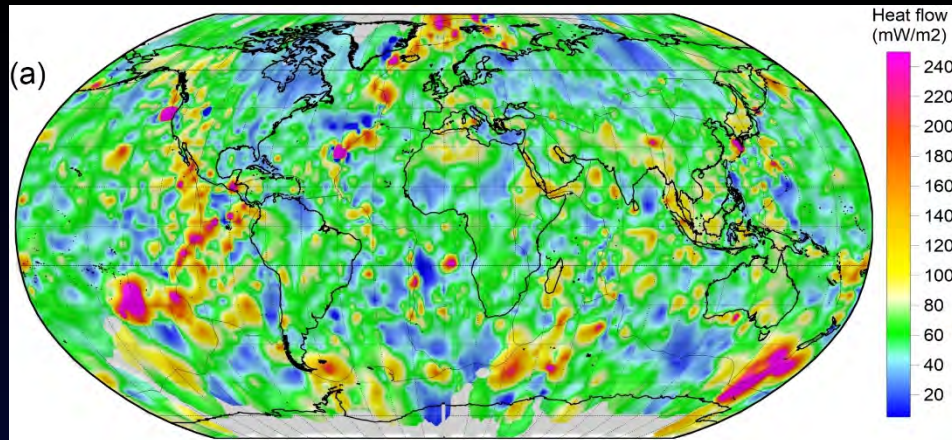
Heat flow

Interpolation

Point data

No systematic patterns are recognized in heat flow data due to strong heterogeneity of measured values, which are strongly affected by hydrothermal circulation, magmatic activity, crustal faulting, horizontal heat transfer, and also due to low number of heat flow measurements across many margins.

Data coverage



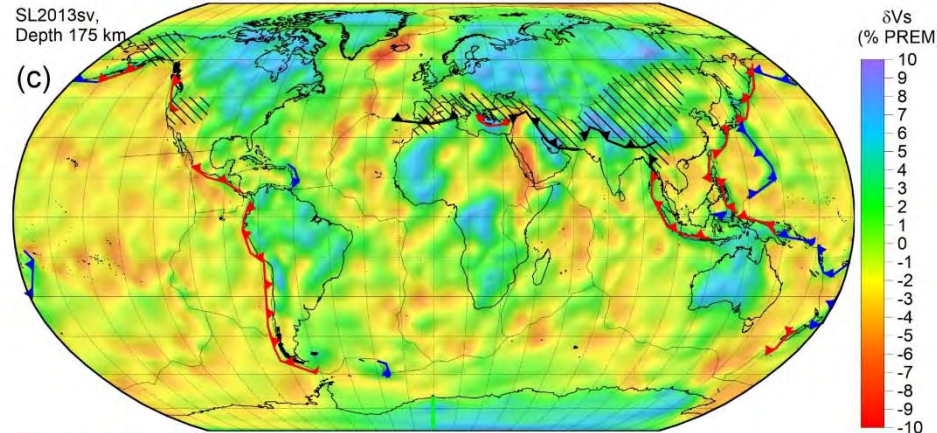
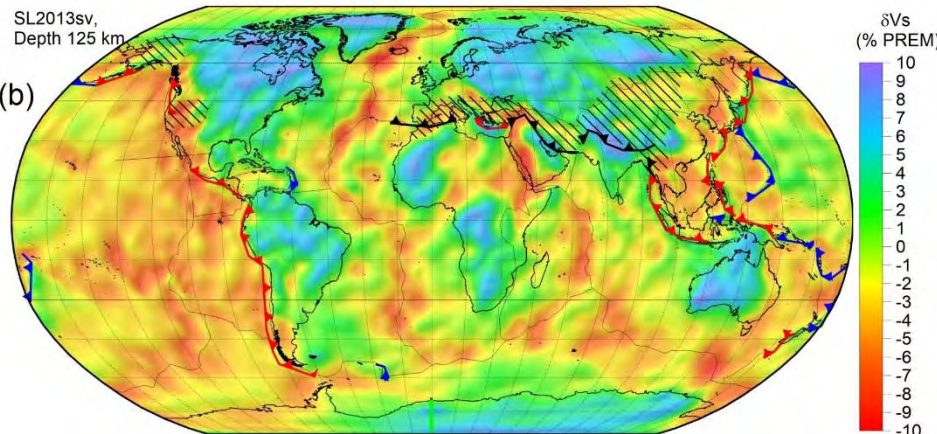
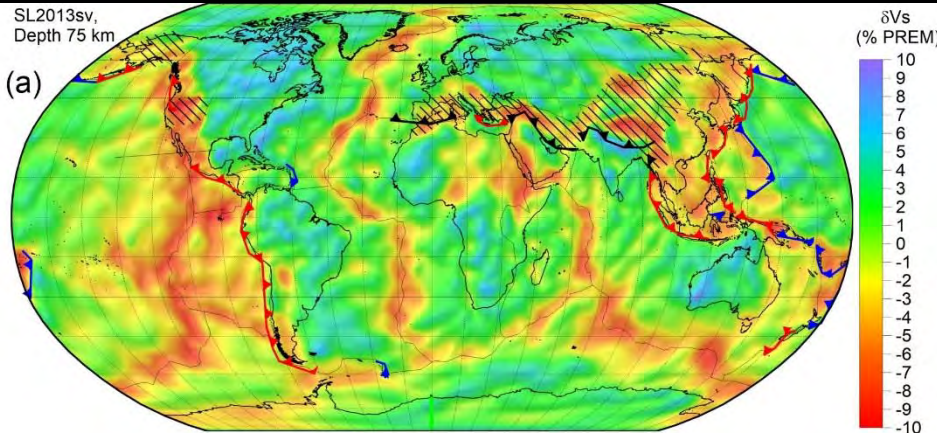
Vs in the mantle

75 km

Low upper mantle Vs seismic velocities beneath the convergent margins are restricted to the upper 150 km and may be related to mantle wedge melting which is confined to shallow mantle levels

125 km

175 km

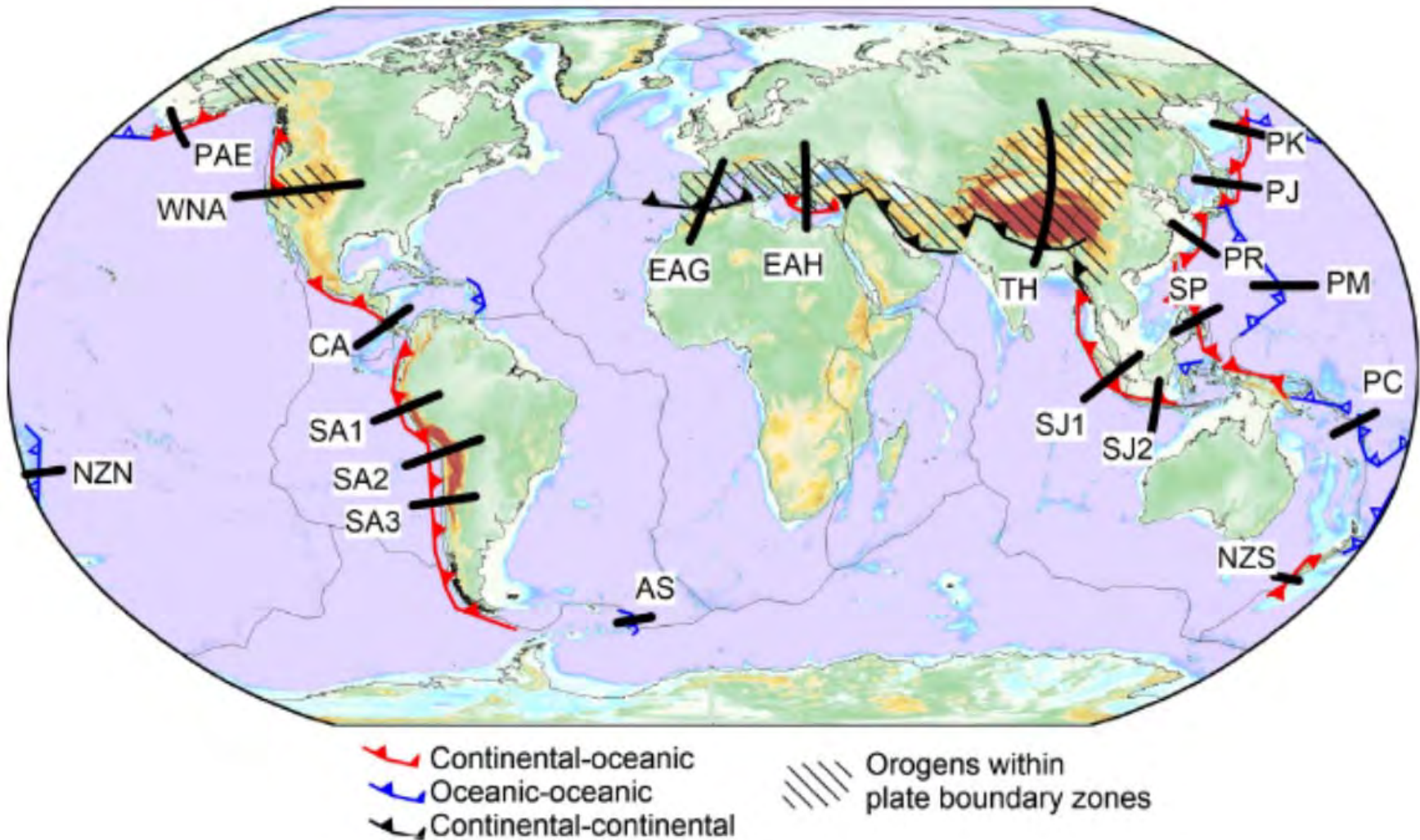


— Continental-oceanic
— Oceanic-oceanic
— Continental-continental

/// Orogens within
plate boundary zones

Model of Schaeffer and Lebedev, 2013

Locations of profiles

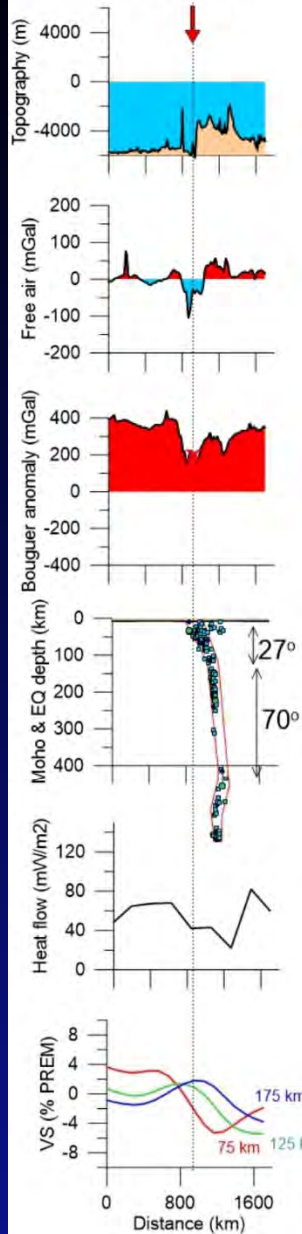


Ocean-ocean convergent margins

Marianna

OO, 27 mm/y

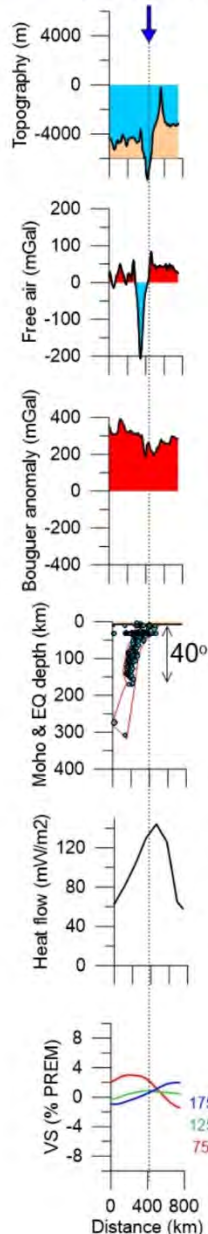
E <PM> W



Sandwich

OO, 57 mm/y

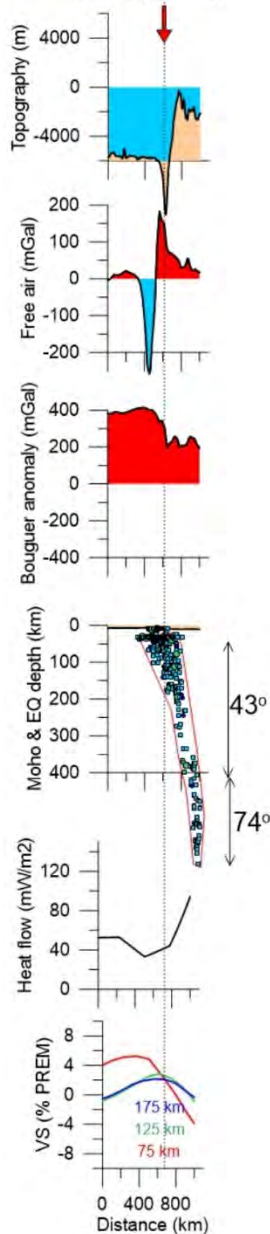
E <AS> W



New Zeal. N

OO, 72 mm/y

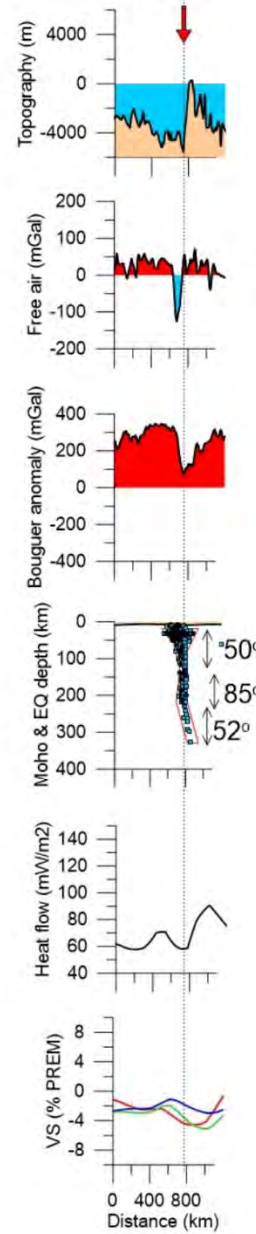
E <NZN> W



New Caled.

OO, 92 mm/y

E <PC> W



Topography

Free Air

Bouguer

Seismicity, $M > 4.0$

Heat flow

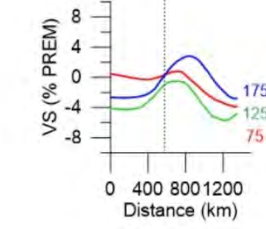
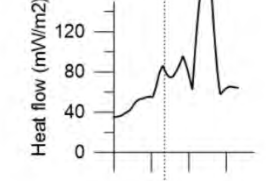
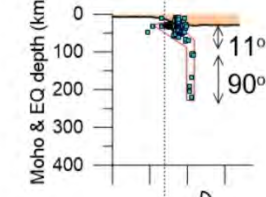
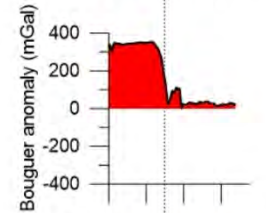
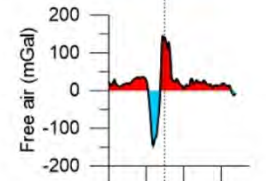
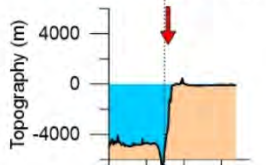
V_s @ 75, 125, 175 km

Continent-ocean convergent margins

Alaska

CO, 66 mm/y

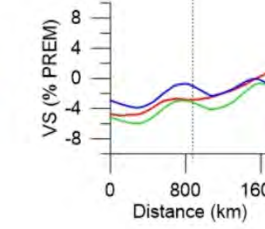
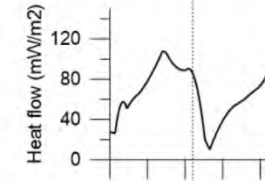
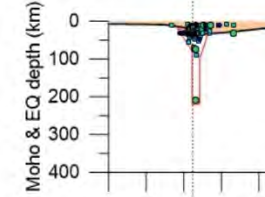
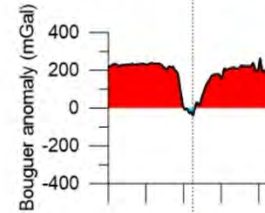
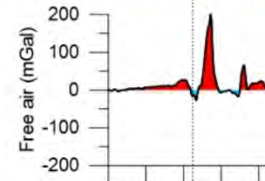
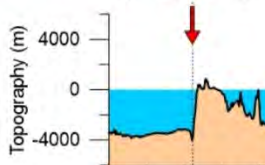
SE <PAE> NW



Central Am.

CO, 78 mm/y

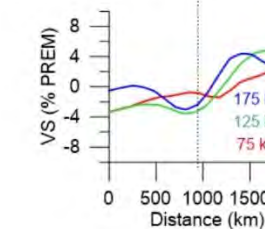
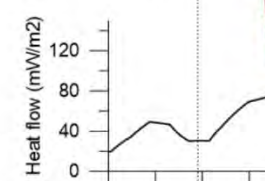
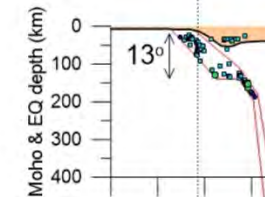
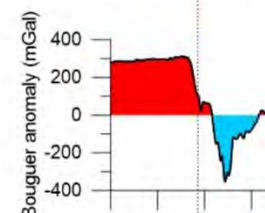
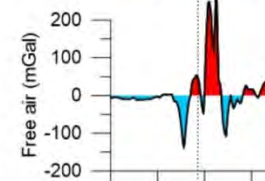
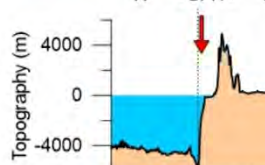
W <CA> E



Andes N

CO, 68 mm/y

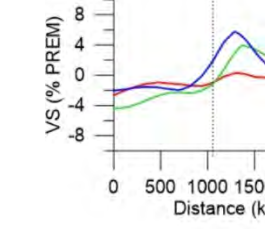
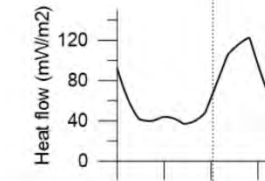
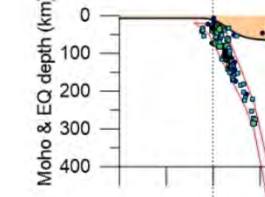
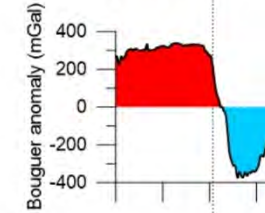
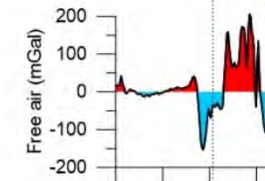
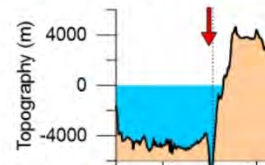
W <SA1> E



Andes C

CO, 72 mm/y

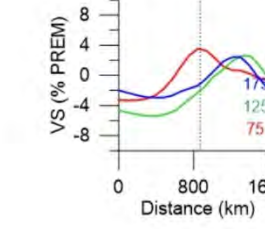
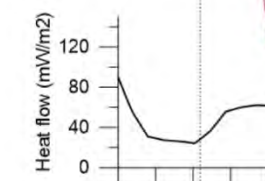
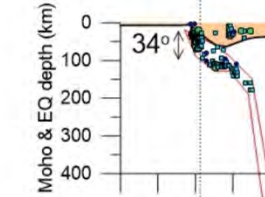
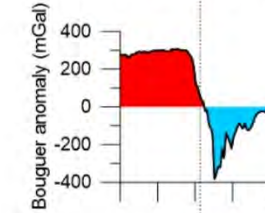
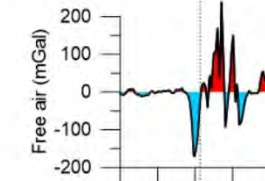
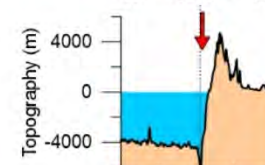
W <SA2> E



Andes S

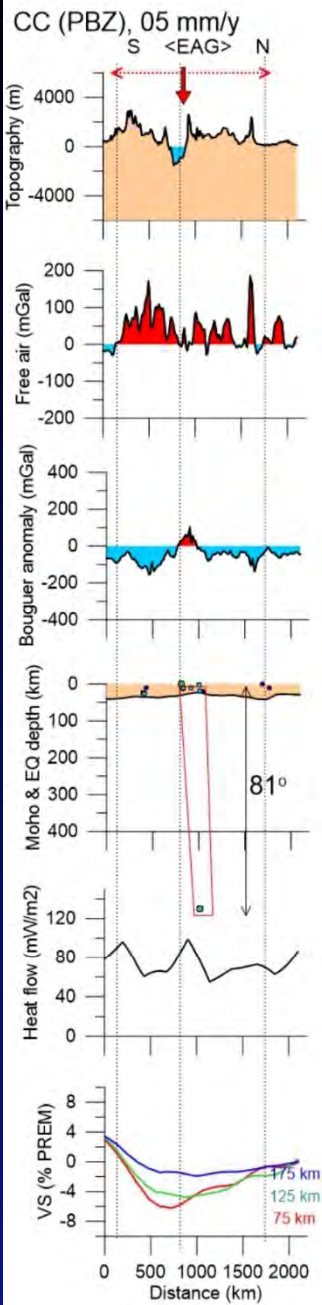
CO, 74 mm/y

W <SA3> E

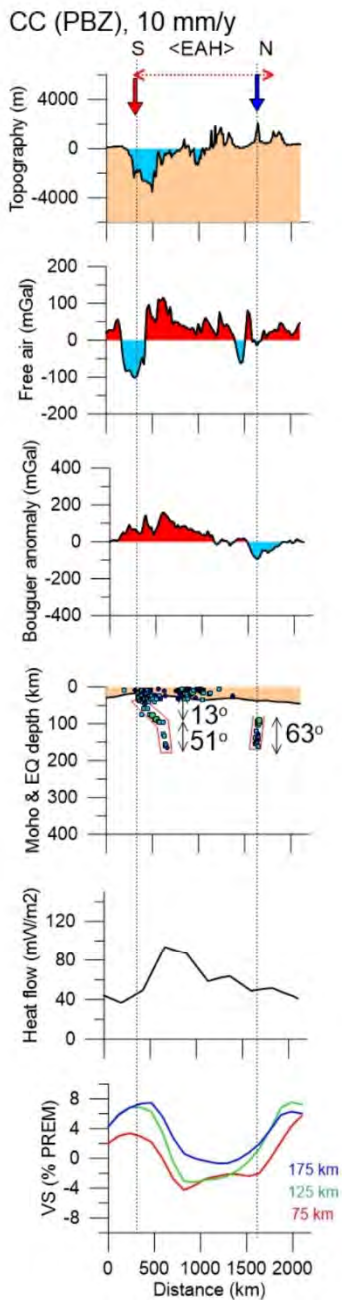


Continent-continent convergent margins and plate boundary zones

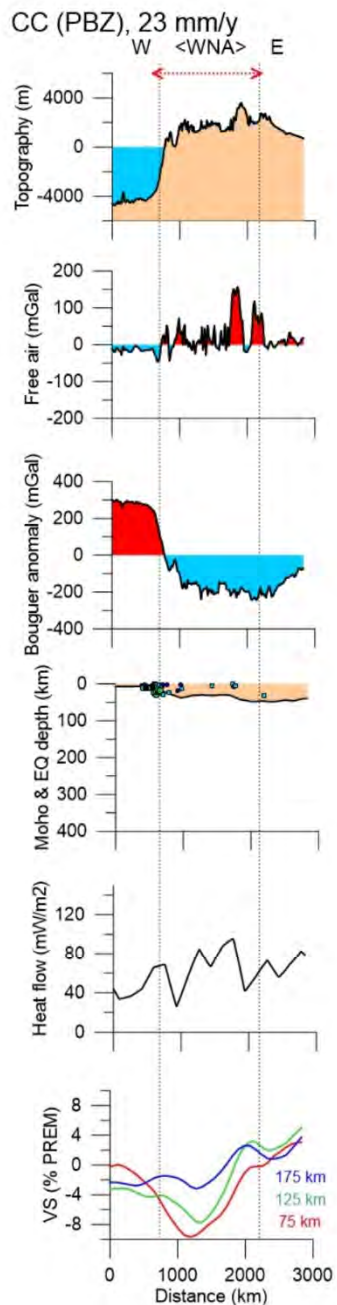
Gibraltar



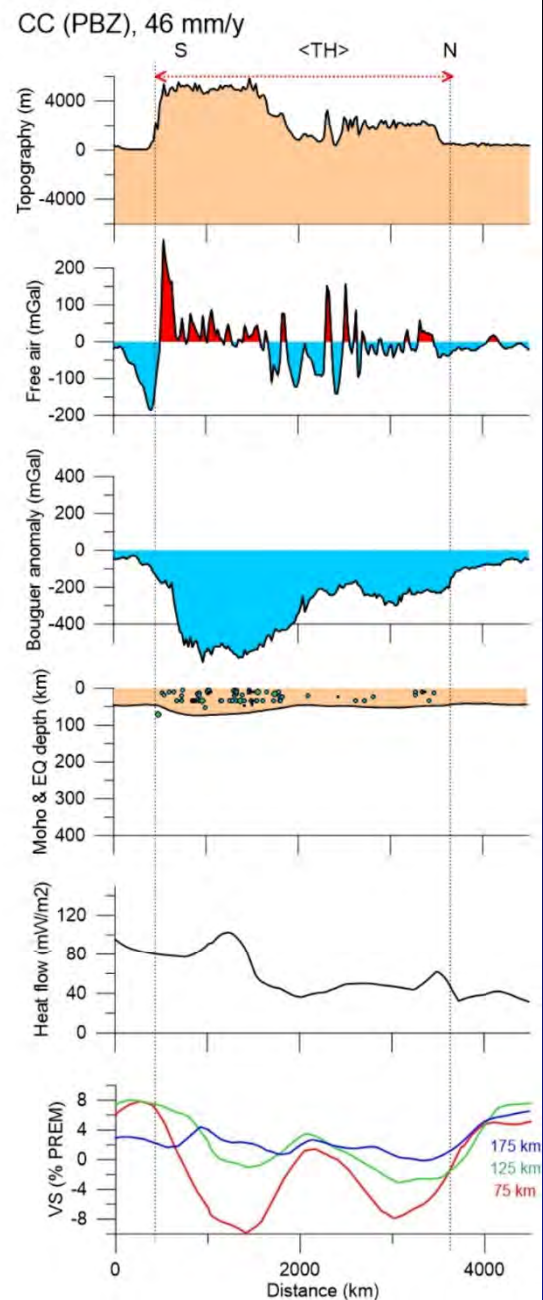
Hellenic arc



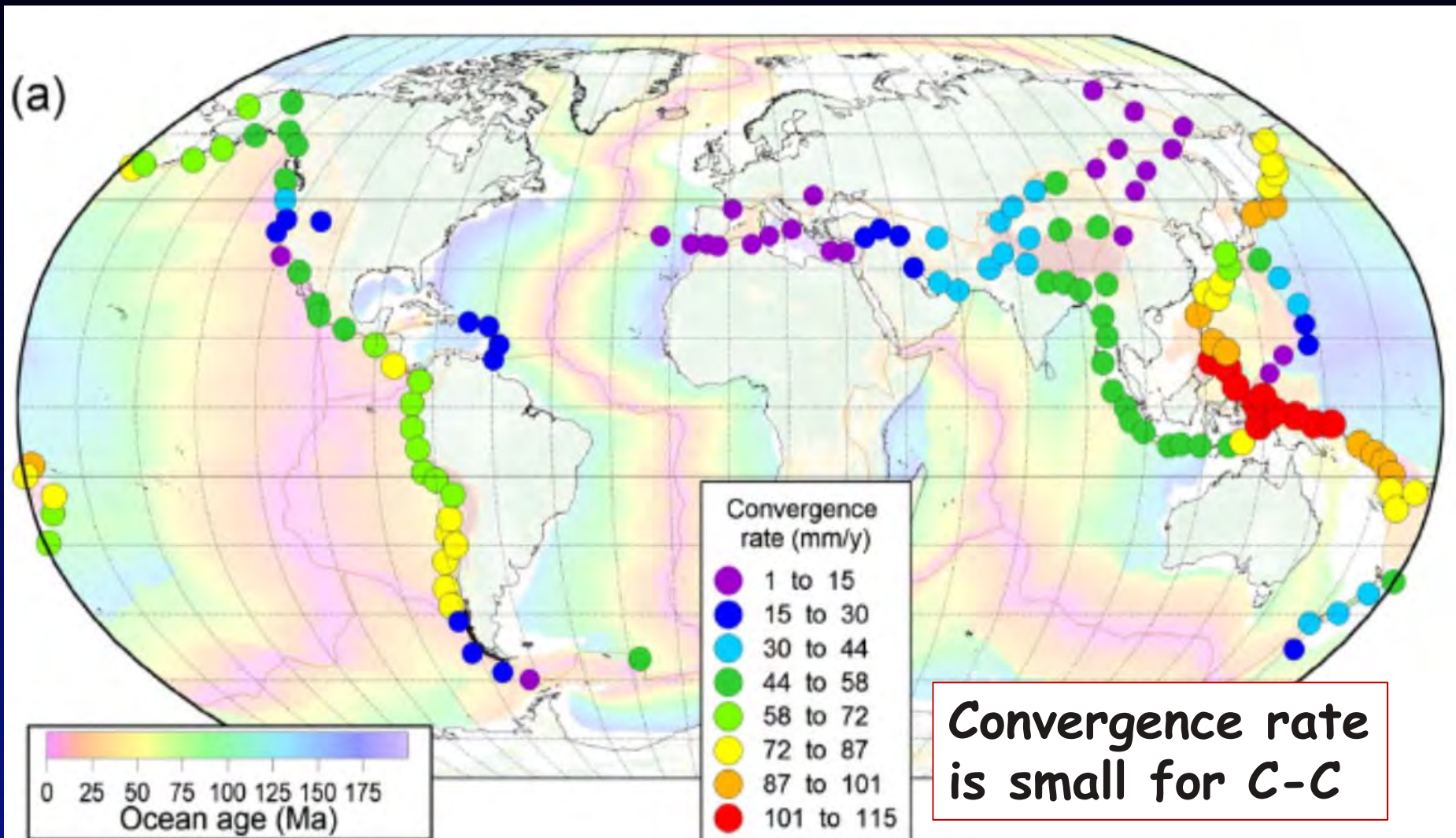
W. N. Am.



Tibet-Himalayas

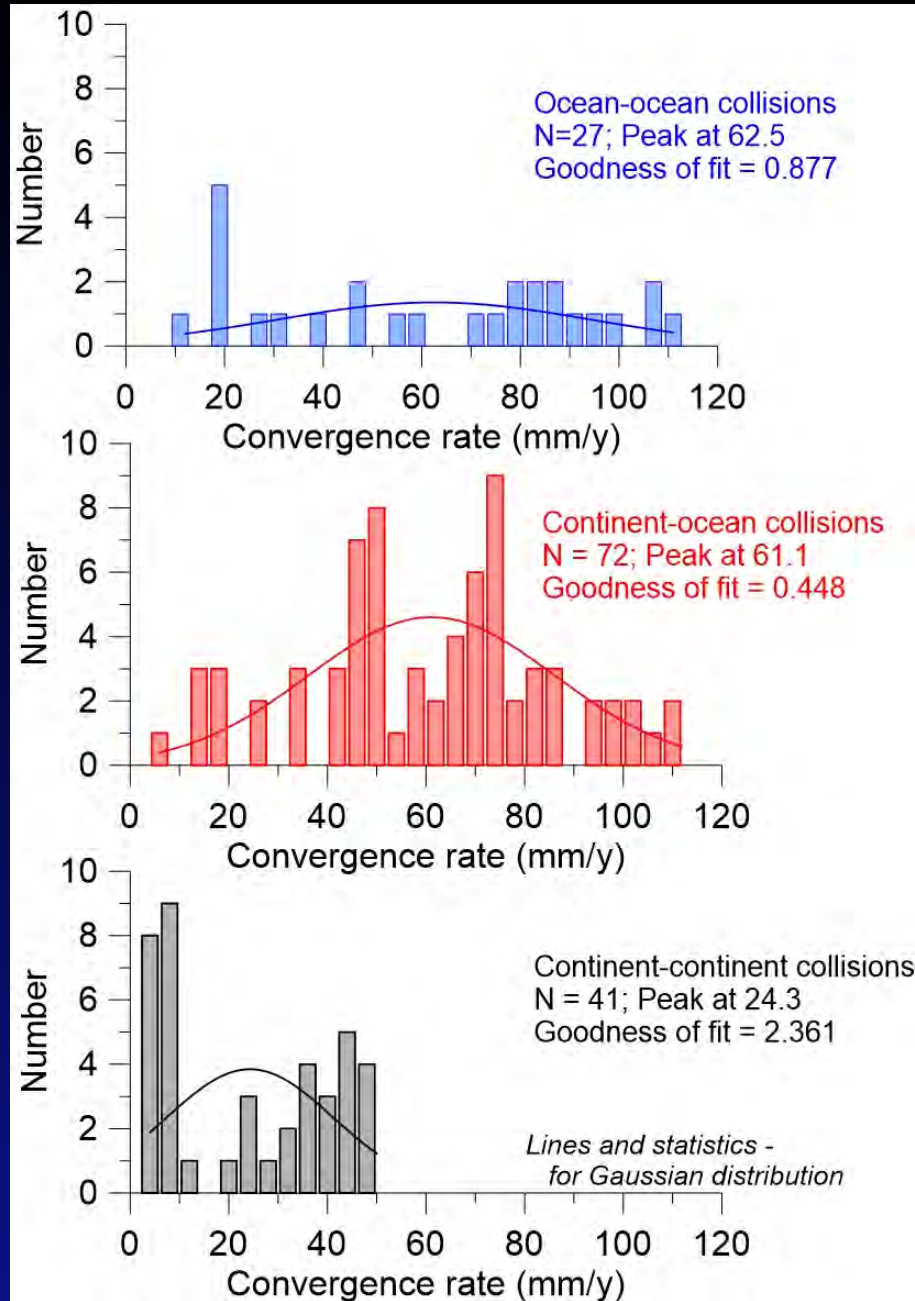


Convergence rate & Age of ocean floor



Age of the ocean floor based on Mueller et al., 2008.

Rates based on the MORVEL plate velocity model, DeMets et al., 2010



Convergence rate

Is larger when overriding plate is oceanic

O-O:

- Broad range
- uniform distribution

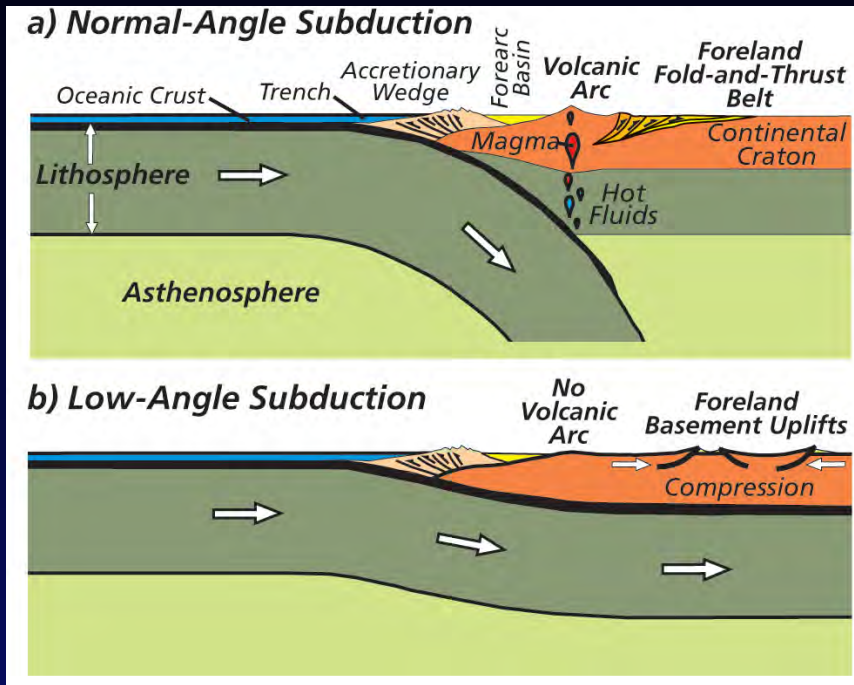
C-O:

- Broad range
- Two peaks at ca. 45 and 75 mm/y

C-C:

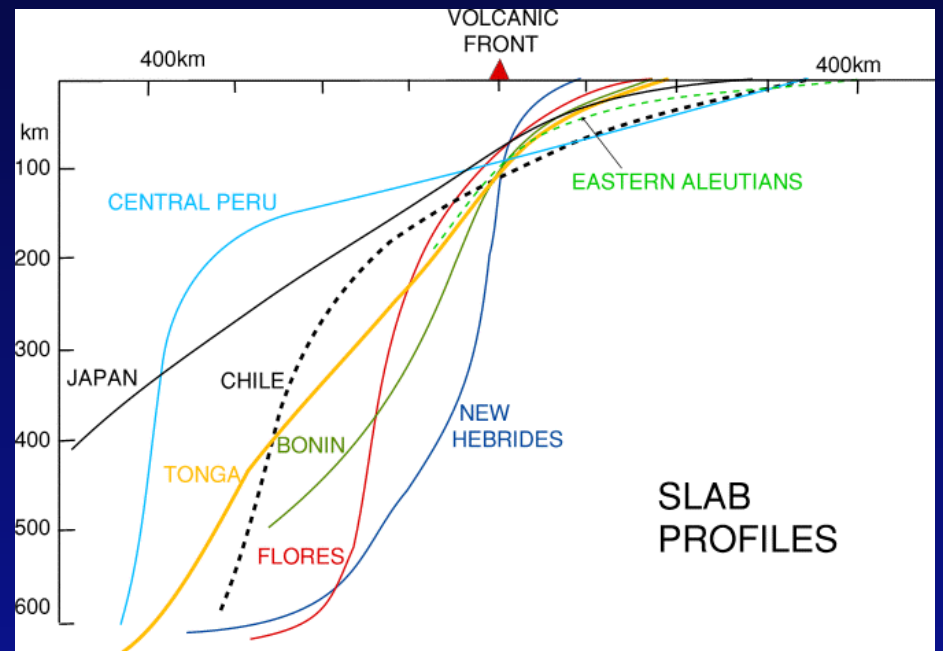
- Small values
- Two peaks: sharp at ca. 5 mm/y and broad at 30-50 mm/y

Subduction angle



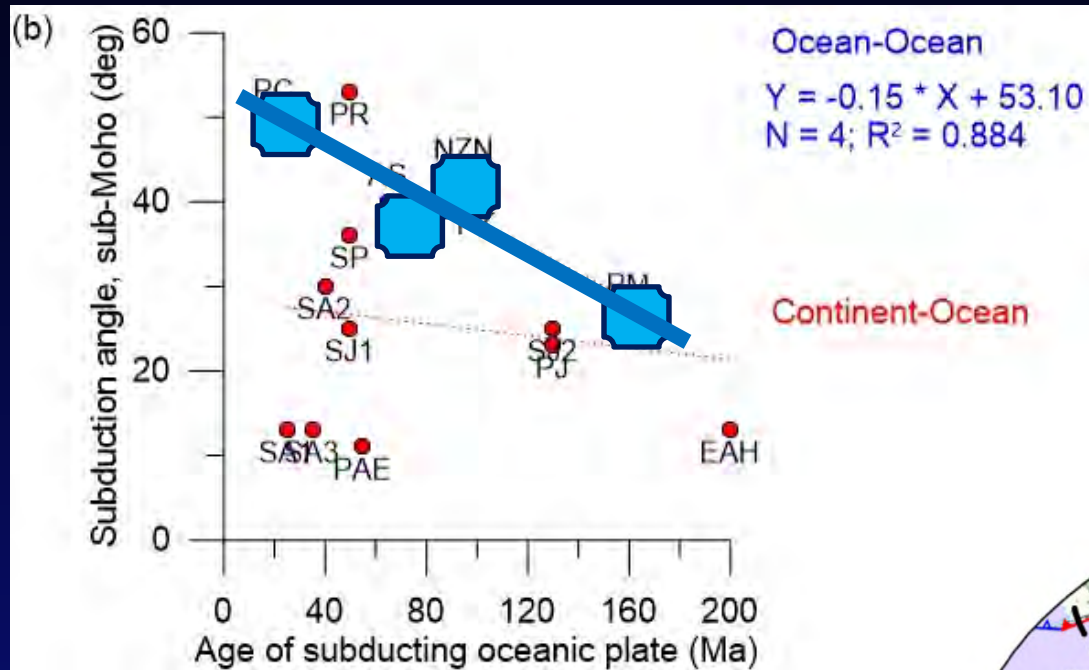
Mariana Type

Chilean Type



Subduction angle: sub-Moho

vs age of incoming plate



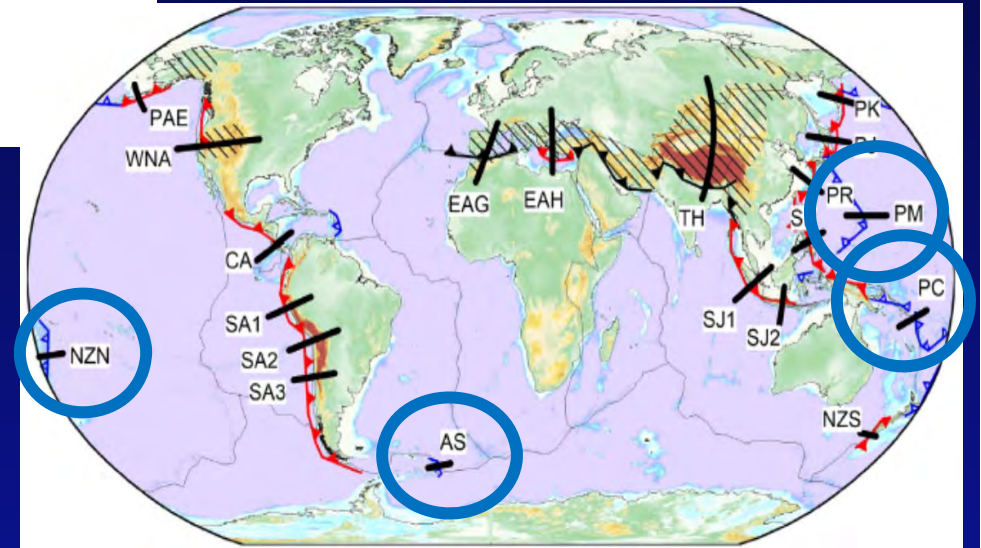
O-O:

Depends on the age of incoming plate (smaller for older plates)

C-O: No correlation

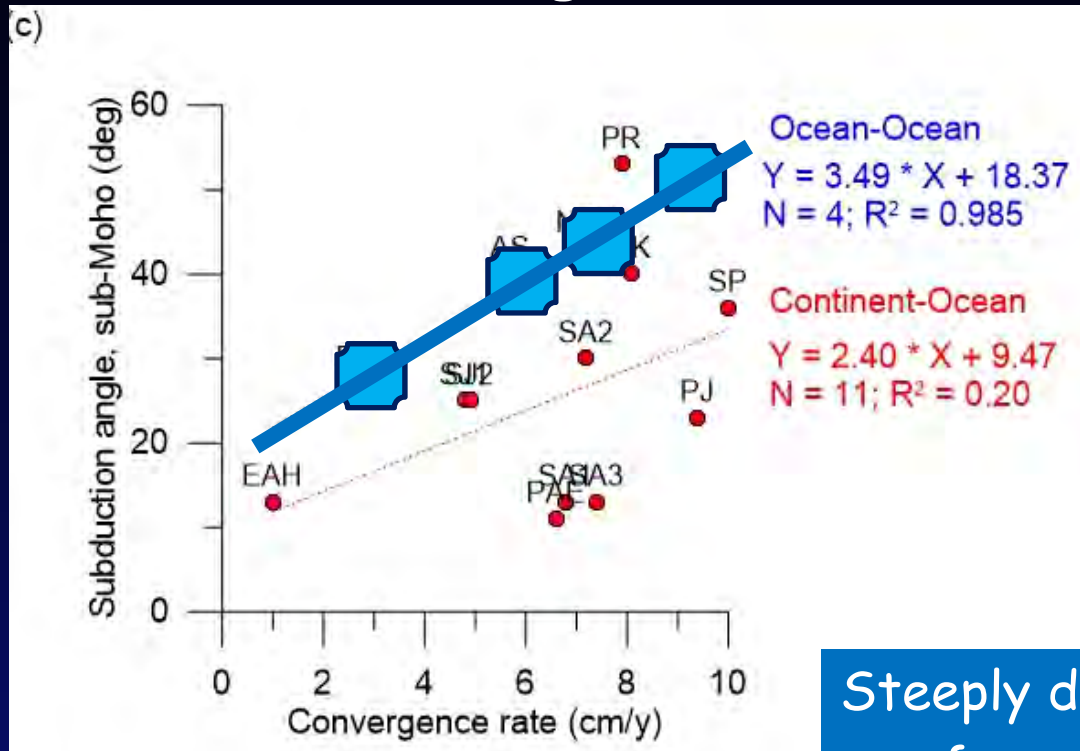
O-O:

Marianna New Zeal. N
 Sandwich New Caled.



Subduction angle: sub-Moho

vs convergence rate



O-O:

Depends on convergence rate
(higher for higher rates)

C-O: Similar trend with
weak correlation

Steeply dipping slabs are characteristic:

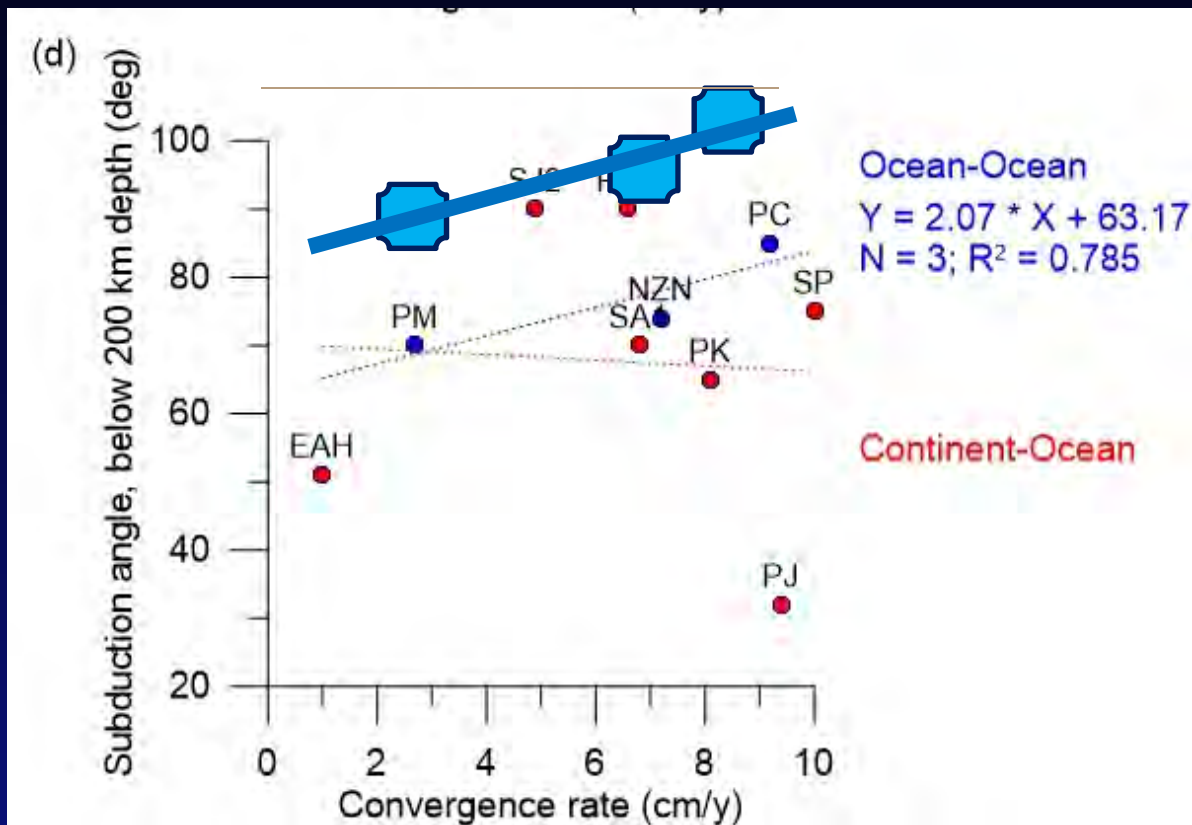
- of young oceanic subducting plates;
- of oceanic plates with high convergence rate,

with slab rotation towards a near-vertical dip angle at depths below ca. 500 km at very high convergence rate.

Marianna and Chile belong to different types (O-O and C-O)
→ comparing apples and pears?

Subduction angle: below 200 km

vs convergence rate

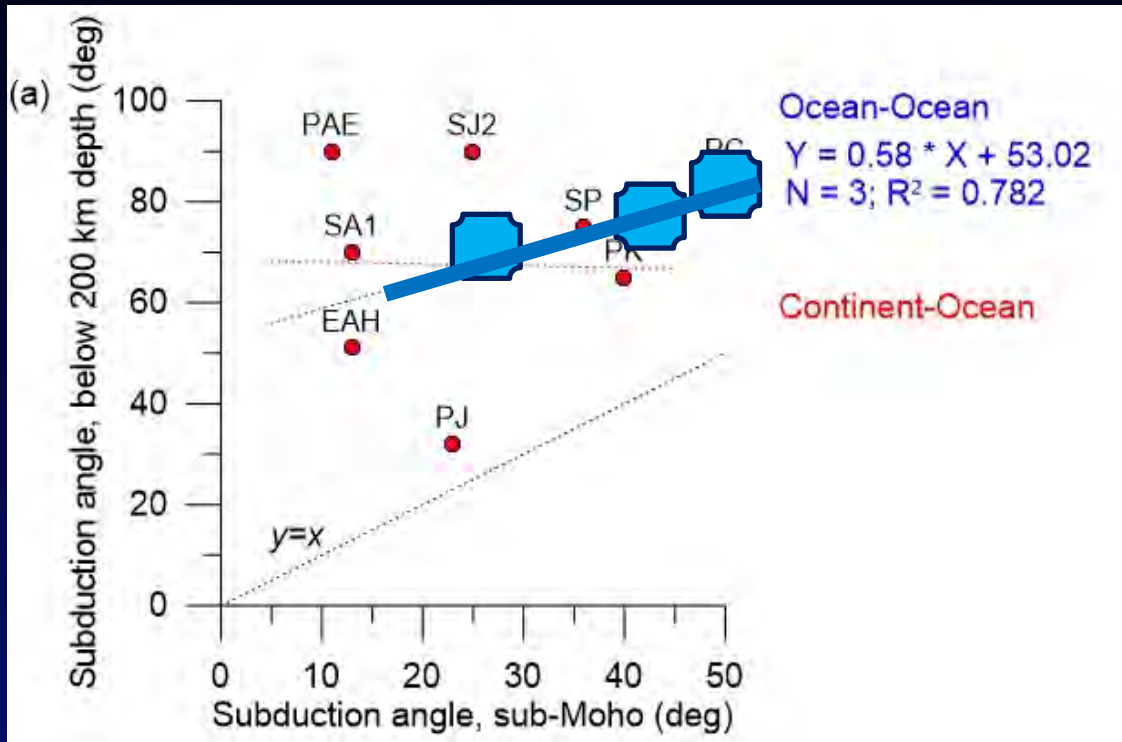


O-O:

Depends on convergence rate (higher for higher rates)

C-O: No correlation

Subduction angle: below Moho and >200 km

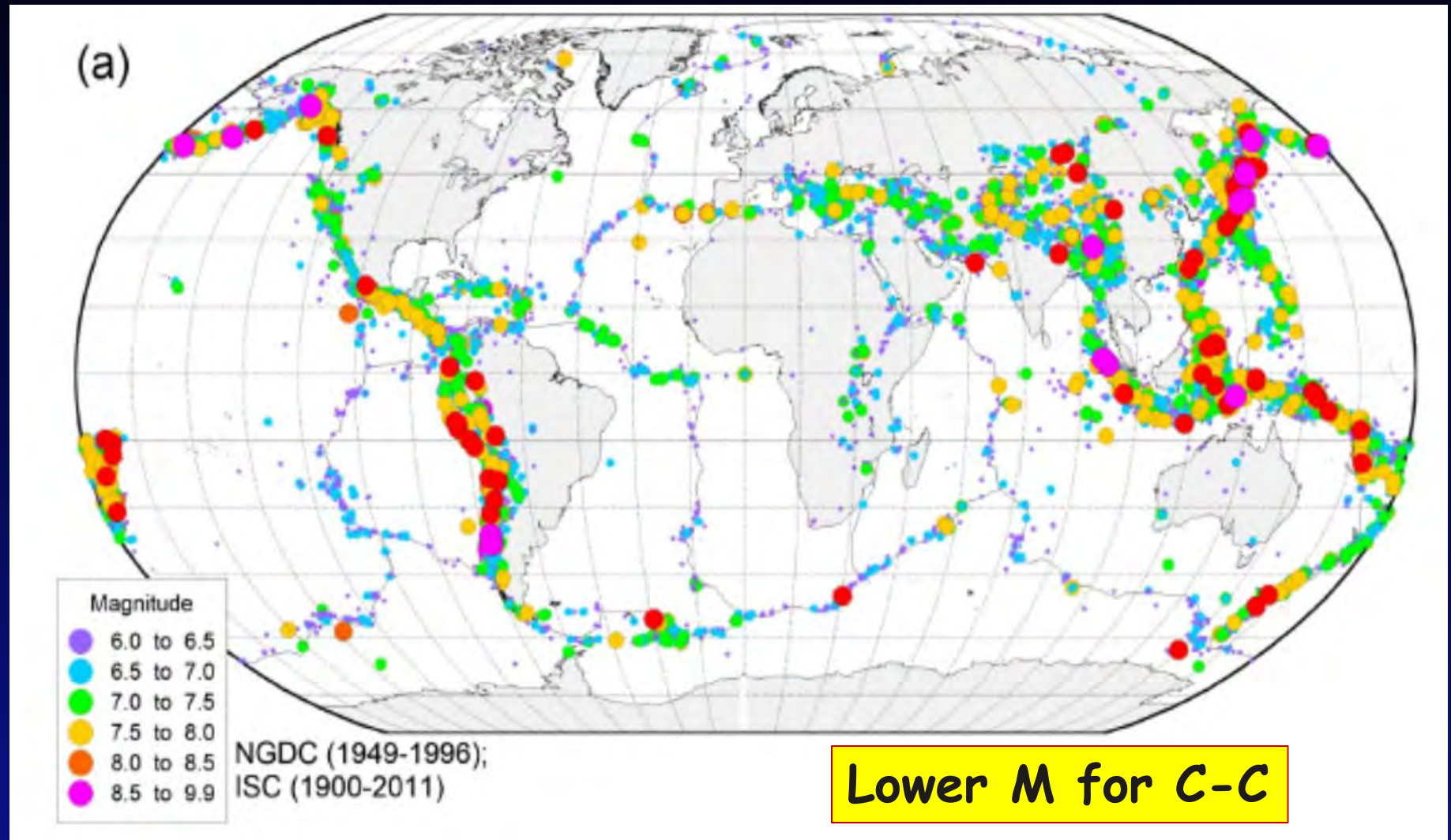


O-O:

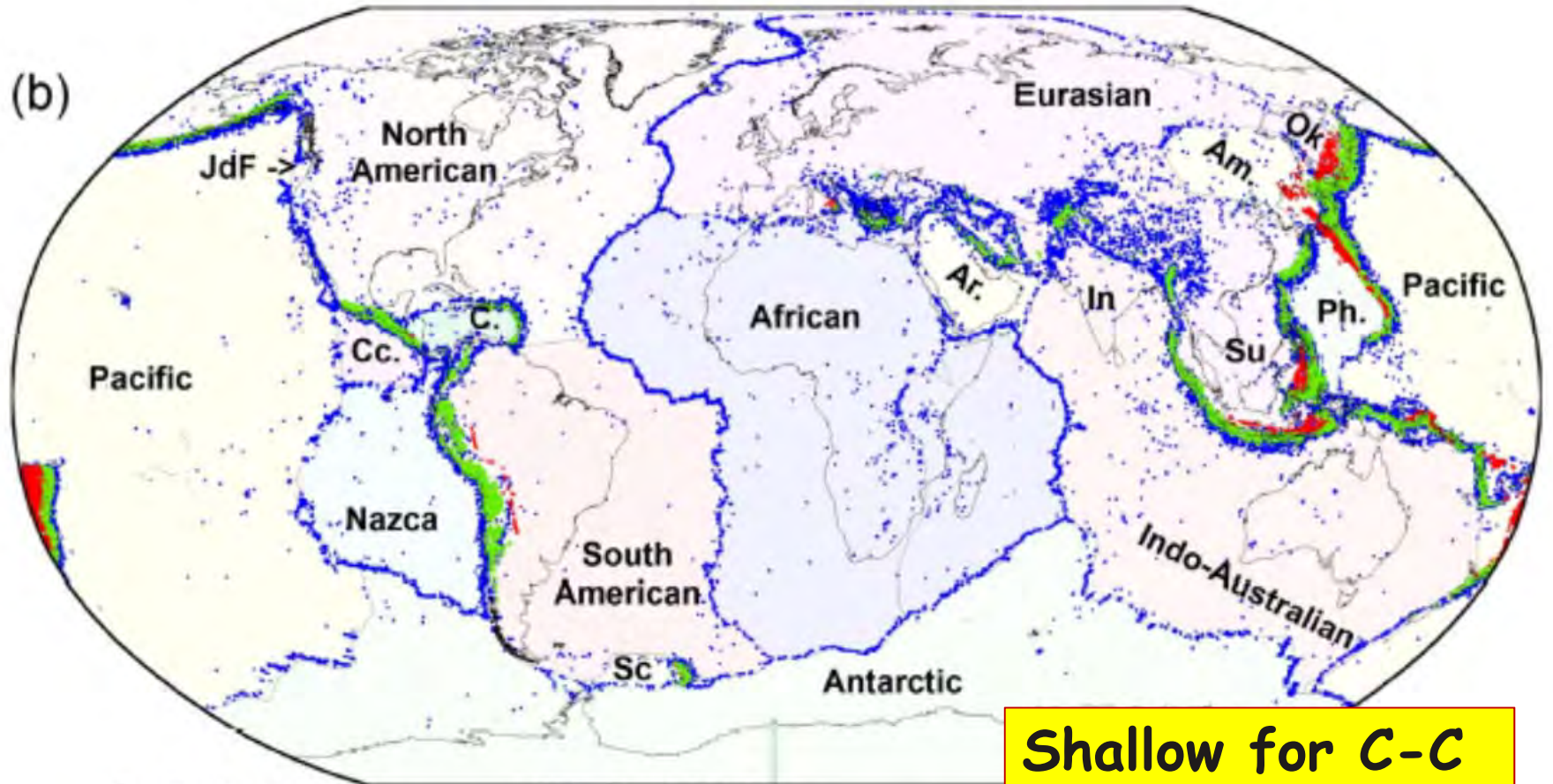
Dip angles for sub-Moho and deep portions are different, but correlated

C-O: No correlation

Seismicity: Magnitude



Seismicity: Depth



Depth of seismicity

- 0-70 km
- 70-300 km
- 300-700 km

JdF - Juan de Fuca

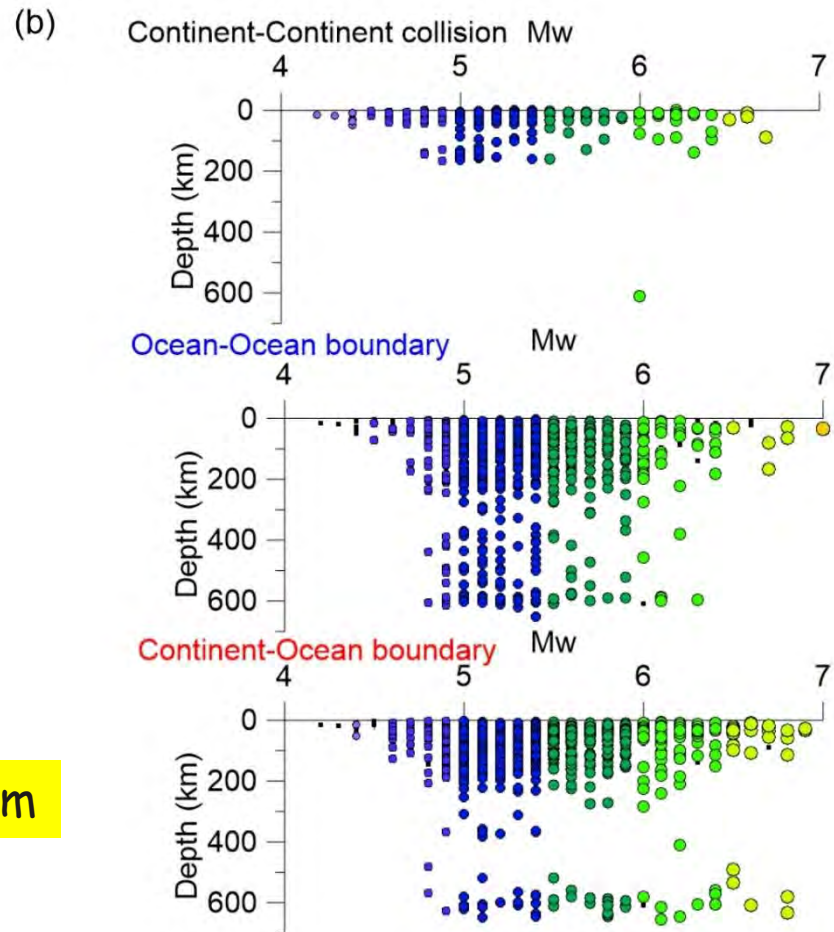
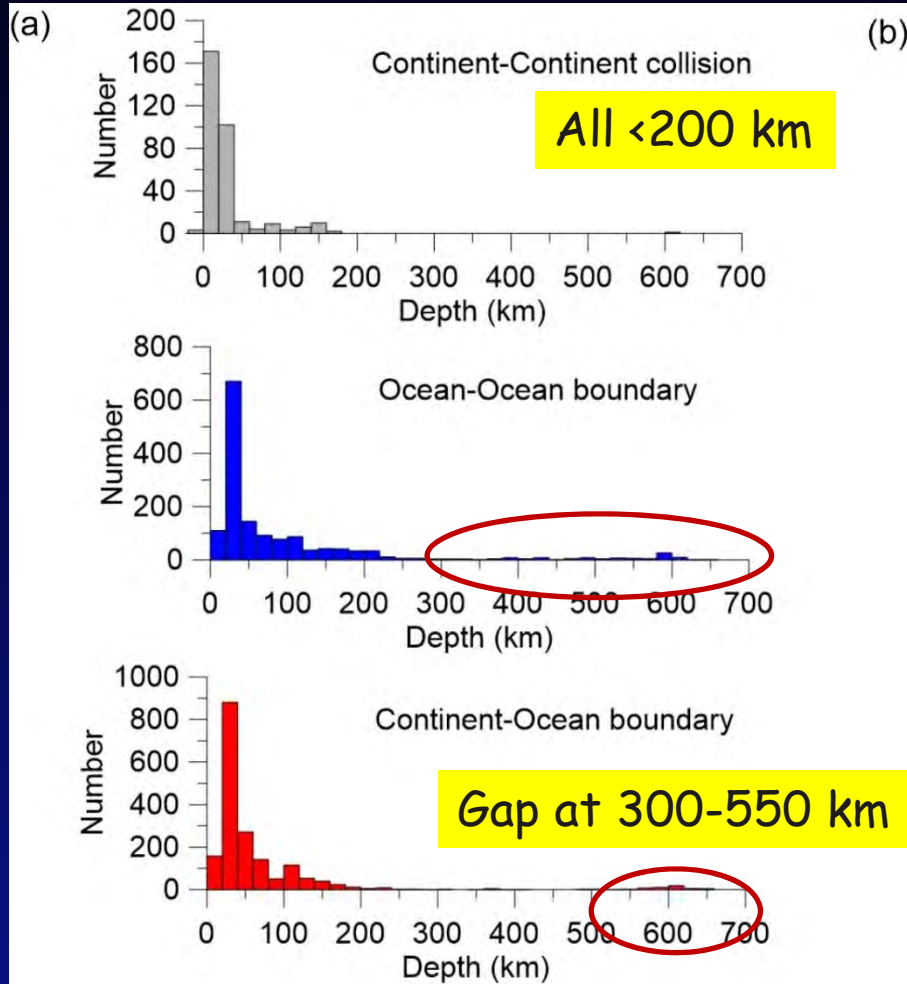
- Cc. - Cocos
- C. - Caribbean
- Sc. - Scotia

- Ar. - Arabian
- Ph. - Phillipine
- Su - Sunda

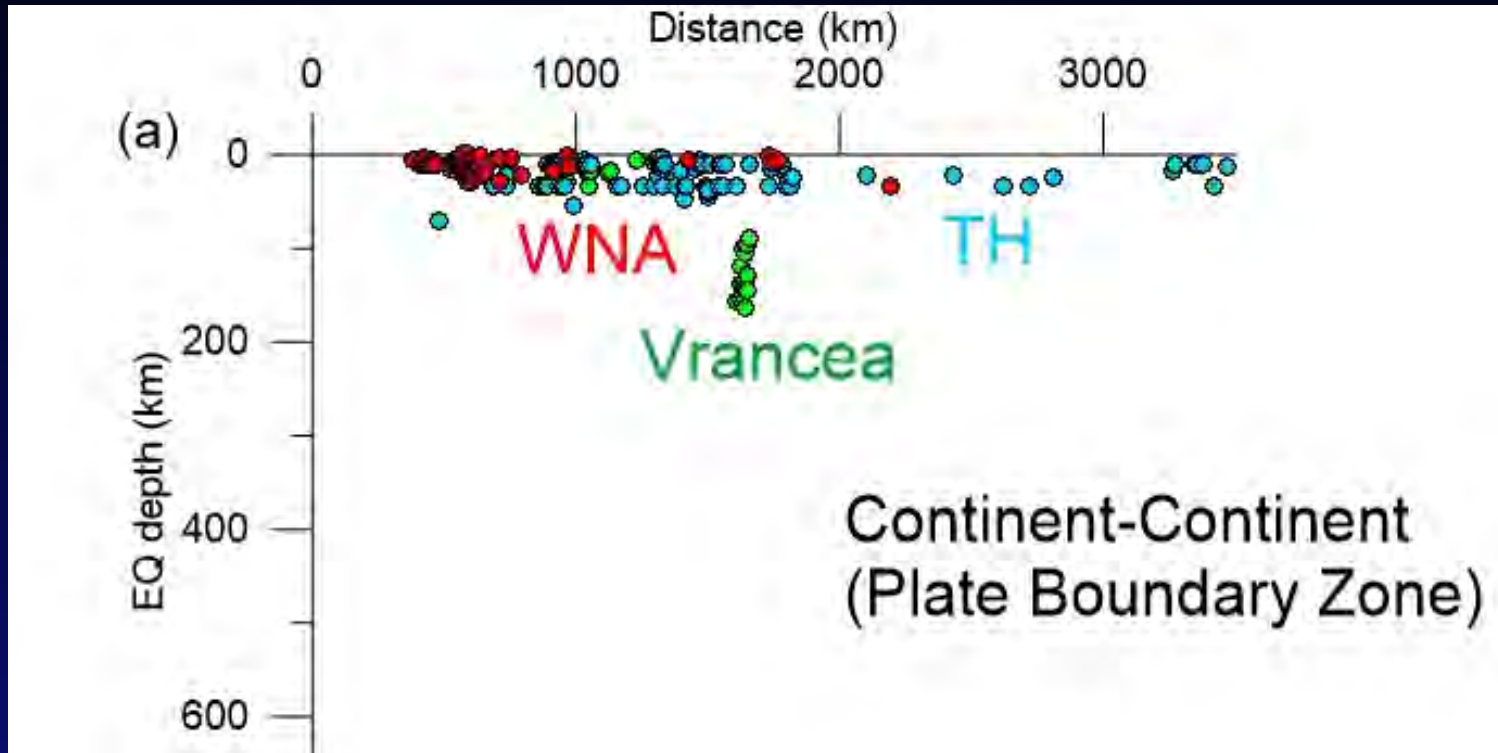
- Am. - Amurian
- Ok. - Okhotsk
- In. - Indian

Seismicity vs Depth

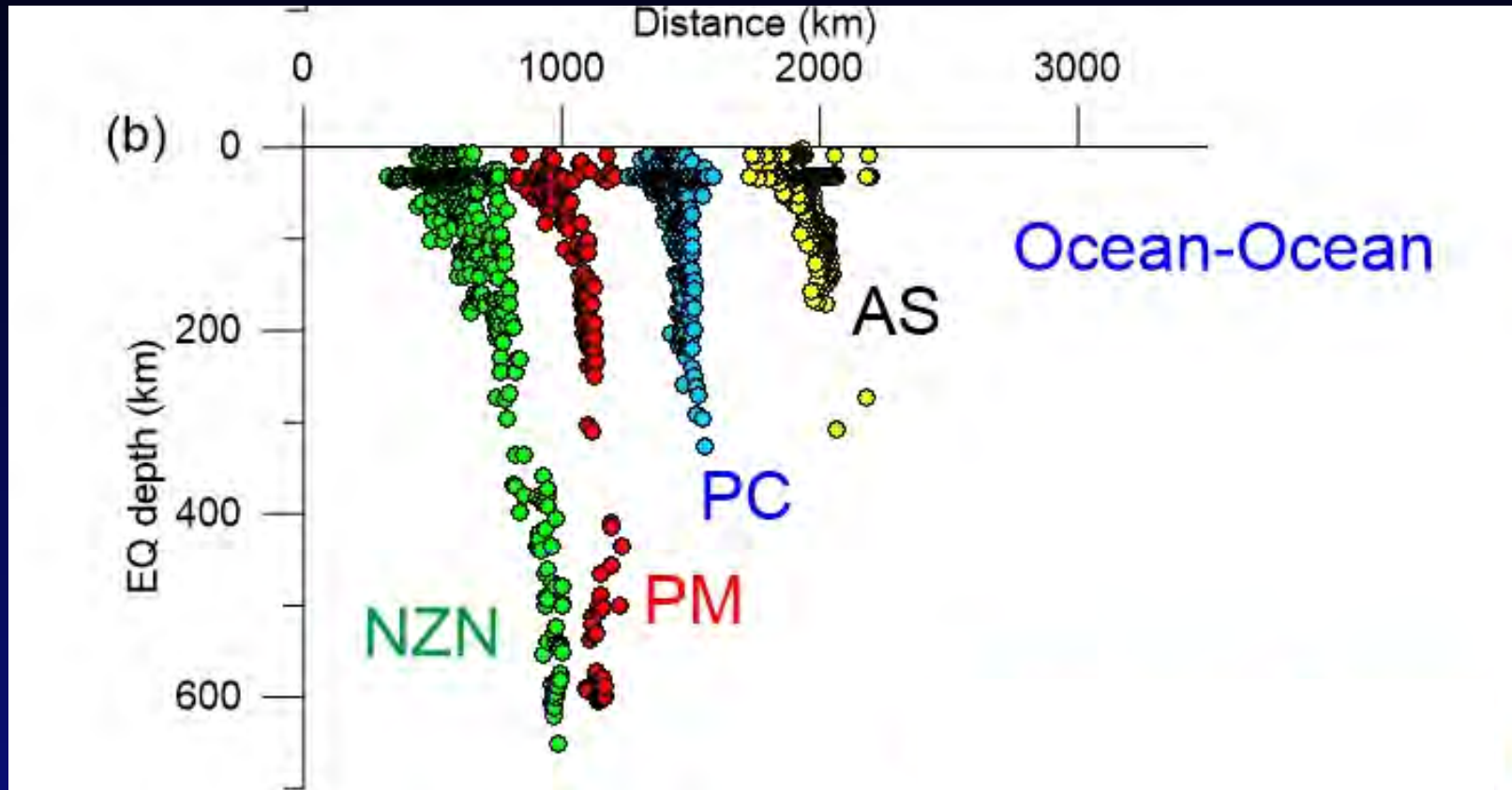
$M > 8$ occur only when the overriding plate is oceanic



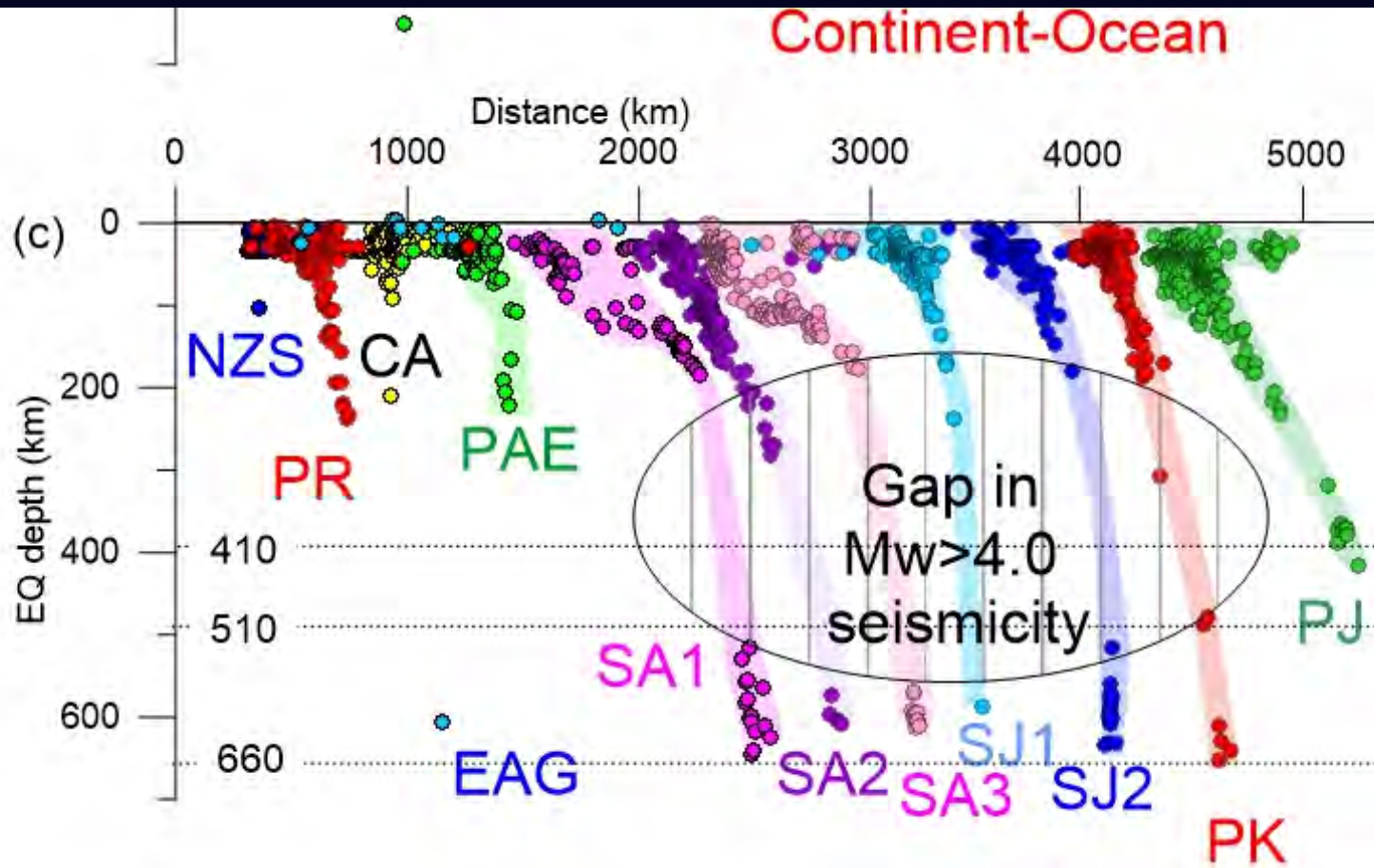
Seismicity vs Depth



Seismicity vs Depth



Seismicity vs Depth



Conclusions

1. Systematic patterns across subduction zones:

- Exist for Free Air and Bouguer;
- Do not exist for heat flow and Vs in global models
- -> large variability and "individuality" of subd. Zones

2. Subduction angle correlates only for O-O subd.:

- Sub-Moho:
 - smaller for older plates;
 - higher for higher conv. rates;
- Below 200 km:
 - weakly correlates with sub-Moho dip;
 - higher for higher conv. rates

3. Seismicity

- Gap at ca. 300-520 km at C-O margins