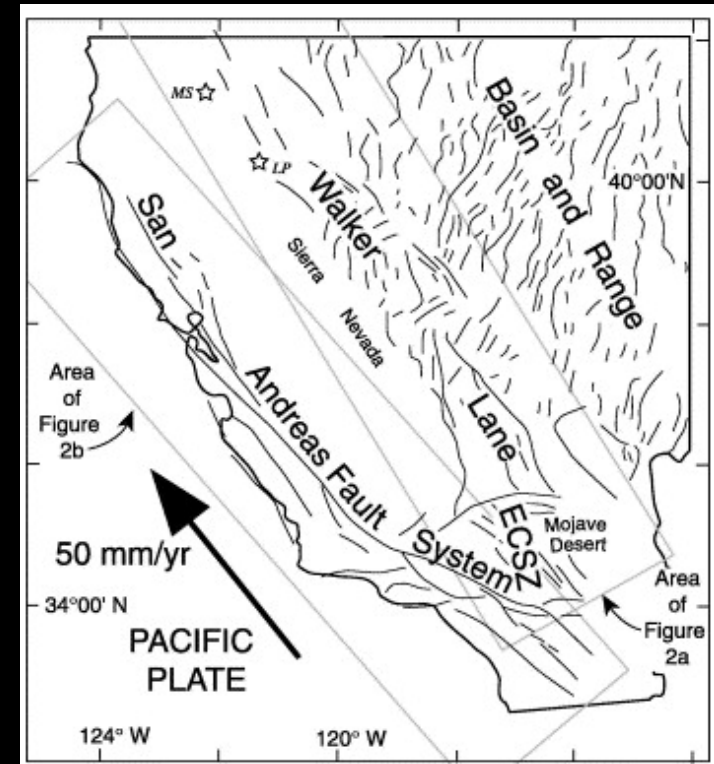


The Walker Lane RIFT SYSTEM: A Natural Laboratory to Study Rift Initiation that Culminated in Seafloor Spreading (in the Gulf of California)



Why am I in charge of this pitch? Because I am representing everyone who was at the Penrose I ran in Walker Lane last August (but heard about this meeting too late to apply).....



As well as the people who are here.....Cathy Busby (UCSB), Graham Kent (UNR), Neal Driscoll (SIO), Chris Henry (UNR), Jim Faulds (UNR) Glenn Biasi (UNR), Ken Smith (UNR), John Louie (UNR), Bill Hammond (UNR), Alistair Harding (SIO), Geoff Blewitt (UNR), Corne Kreemer (UNR), Danny Brothers (USGS), Fred Phillips (NMIT), Jared Kluesner (SIO), Peter Lonsdale (SIO), Keith Putirka (Fresno St.), Debi Kilb (SIO), Pat Cashman (UNR), Paul Umhoefer (NAU), Gary Fuis (USGS), Dan Lizarralde (WHOI), Jeff Babcock (SIO), Kathie Marsaglia (CSUN)

Modern Tectonic Context

Walker Lane rift system - currently accommodates 20-25% of the plate motion between the Pacific and the North American plates

Provides a clear record of rift initiation...

and a time-transgressive view of

processes involved in strain localization.



Sierra Nevada microplate - Walker Lane system:
Global importance because of the many important concepts developed there and exported to other parts of the world (see below).

It remains, however, to integrate disparate geologic, geochemical, and geophysical discoveries into a comprehensive physically based process model for continental rifting.

This requires a long-term program that brings many different types of workers together.

listric normal faults, detachment faults and metamorphic core complexes, “chaos” (large-scale landslide deposits), calderas and geothermal fields (e.g. Long Valley), maar volcanoes and Pluvial lake deposits, vertical axis rotations of crustal blocks, strain partitioning, thermochronologic dating of ancient landscape surfaces and tectonic tilting events, root delamination, and the emplacement of huge

Attributes of Walker Lane-Gulf of California rift system

IMMEDIATE

GLOBALLY APPLICABLE SCIENTIFIC RESULTS

AT RELATIVELY LOW COST,
WITH MAJOR SOCIETAL IMPACT.

Including
an
amphibious
approach

e.g.
Pyramid
Lake

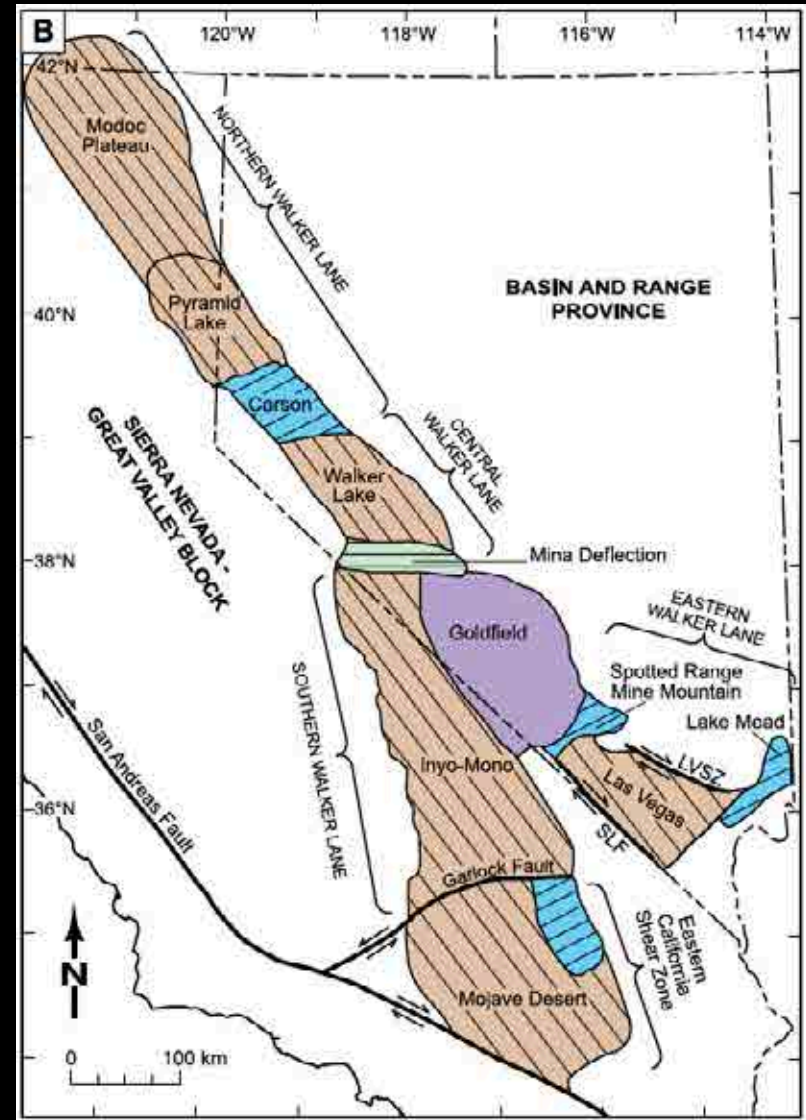


Attributes.....

(1) Data infrastructure (e.g., geologic mapping, digital topography, seismic and geodetic studies)

Integrate disparate geologic maps and analytical data into a single user-friendly database in Google Earth - IMMEDIATE public impact.

Rifting studies can focus immediately on the KEY QUESTIONS.



Attributes, continued:

(2) Ease and safety of physical access - THE POLITICAL REALITY - cost-effective experimentation and data acquisition - Improved scientific return with student and community access.

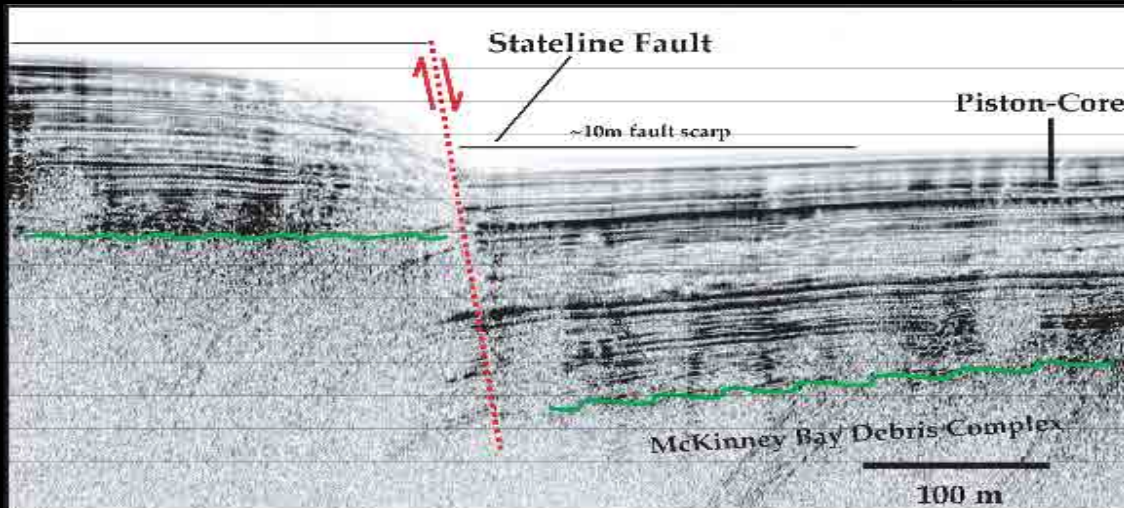
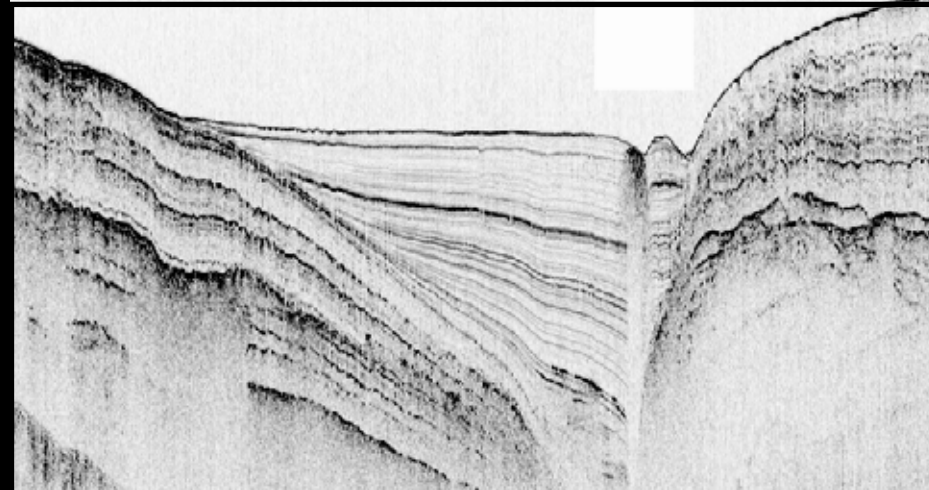
BROADER IMPACT - Students!!!!



Attributes, continued:

(3) Study sites that require access to ocean-going vessels may not begin until 2014 or later, whereas data collection on lakes can begin immediately.

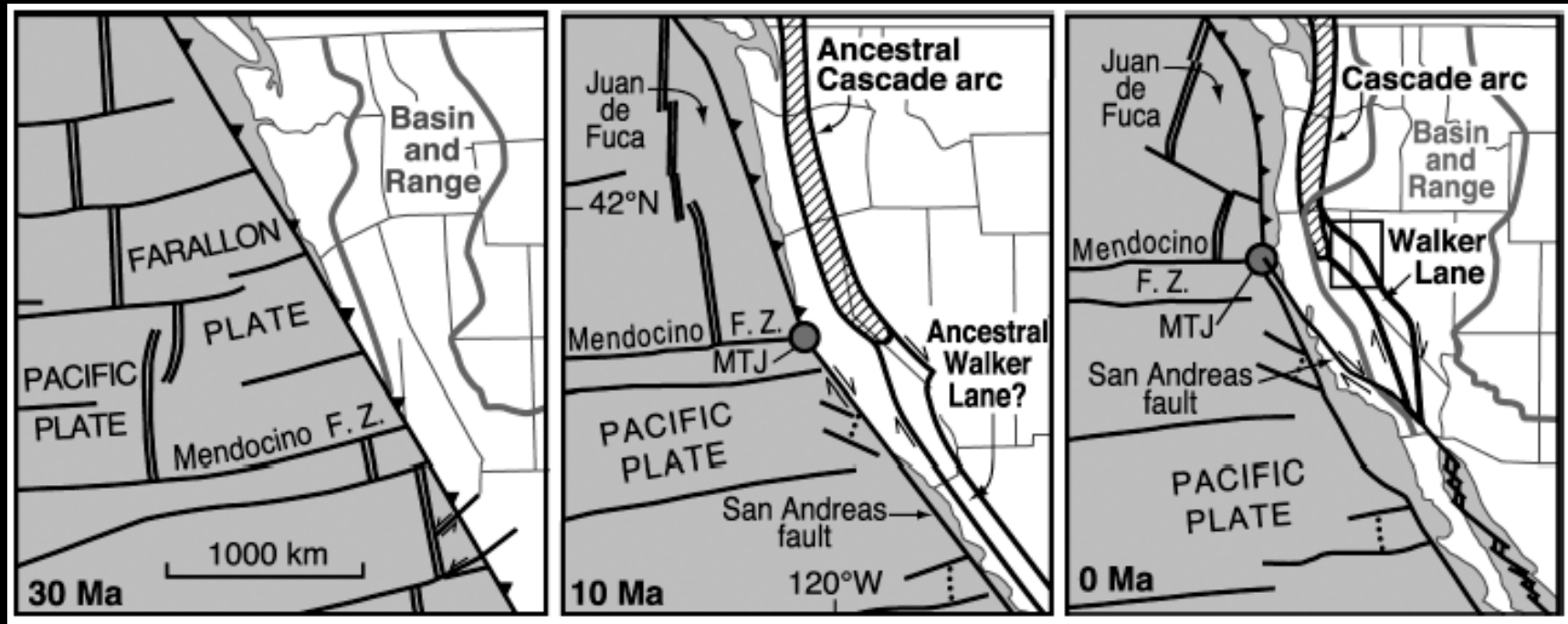
Pyramid Lake, Walker Lane Rift System



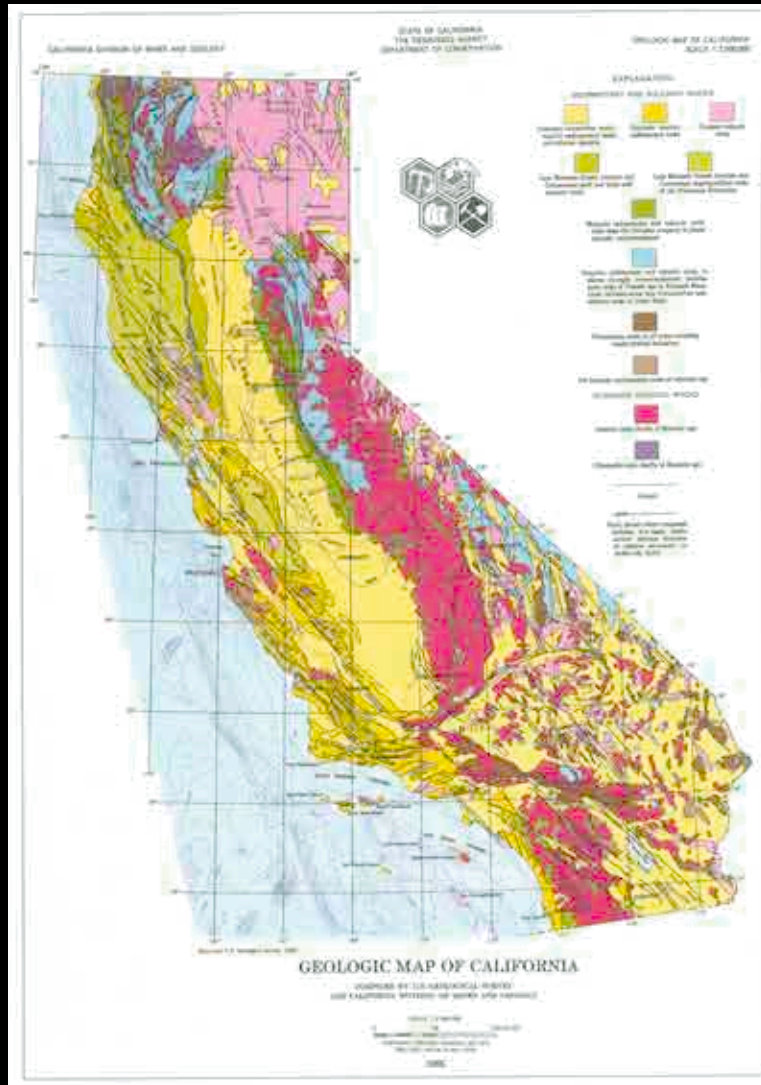
Lake Tahoe,
Walker Lane Rift
System

(4) Due to an on-going history of active rifting from southern Gulf of Ca to northern Walker Lane, **time and space transgressive processes can be observed -**

Post breakup (Gulf) to new rift (Salton Trough) to rift initiation (Walker Lane).....



.....across an array of continental crustal types with a variety of prior magmatic-tectonic histories.



Including Precambrian craton, Paleozoic passive margin, Paleozoic accreted terranes, Mesozoic to Cenozoic subduction margin.

(5) Outstanding exposure and preservation at the FULL RANGE OF STRUCTURAL LEVELS promoting interaction of a...

*****Multidisciplinary team*****

that will be assembled and interact closely
in an

INTEGRATED AMPHIBIOUS APPROACH

e.g. volcanic and sedimentary basins, structural analysis, igneous and metamorphic petrology, deep-level magmatic processes, geochronology and isotope geochemistry, geophysical surveys of crust and mantle features, seismology, geodetics, active tectonics and surface dating, hyperspectral mapping of geology and alteration, economic geology, geothermal exploration, etc

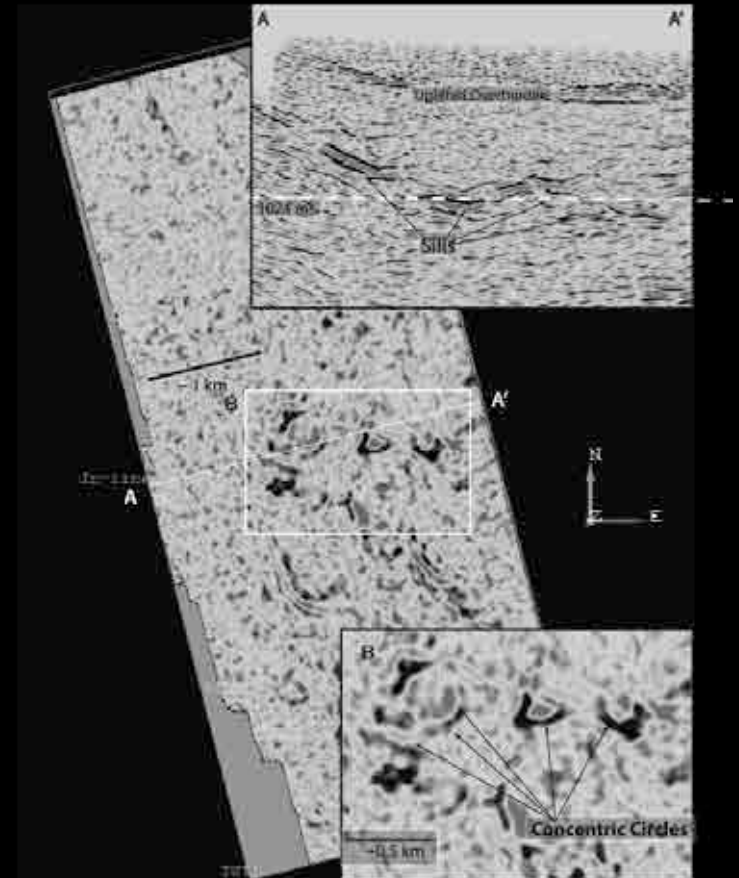
****THIS is what distinguishes us from core NSF*****

Attributes, continued:

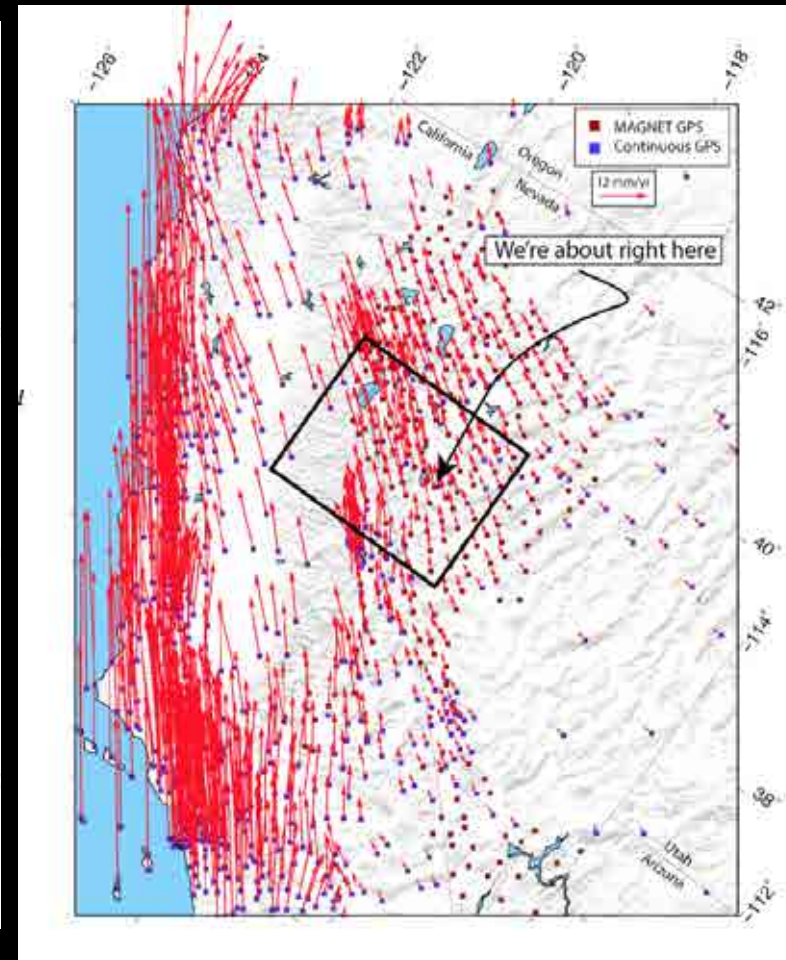
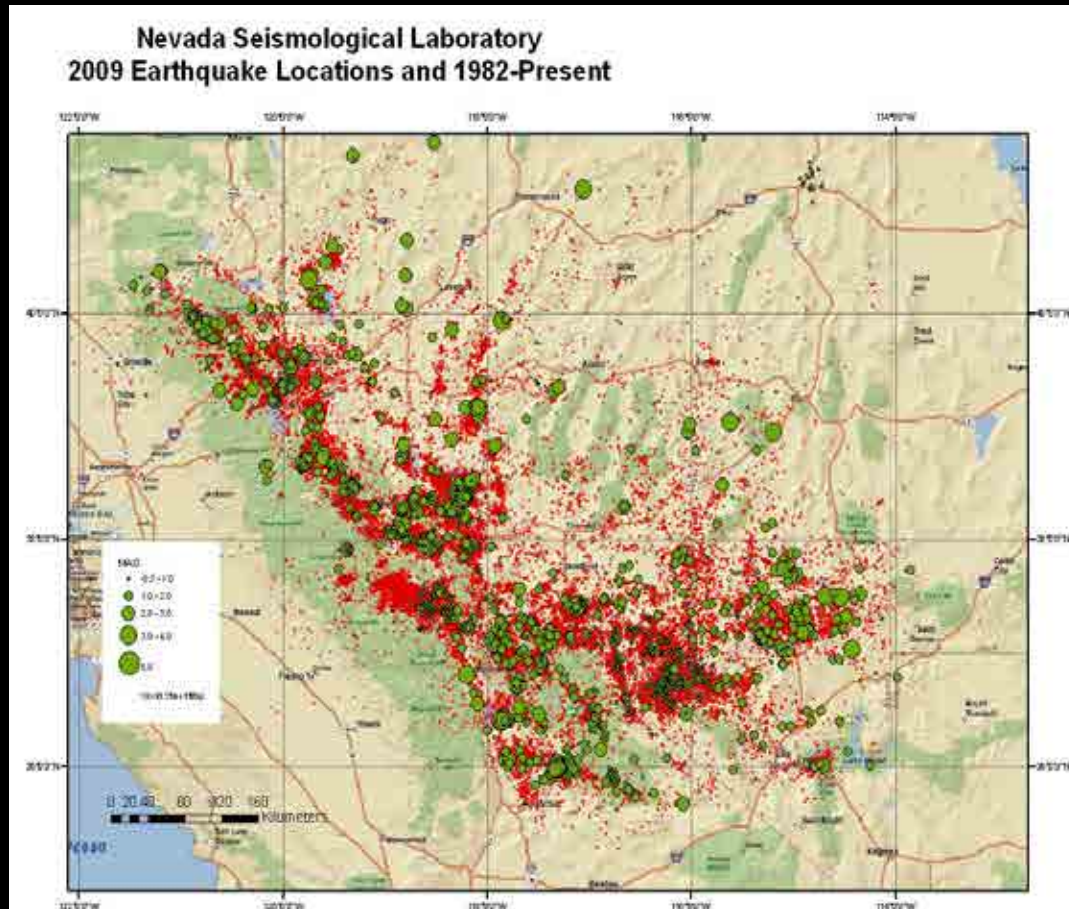
(6) Active magmatism - evaluate **volatile fluxes and sources**; use new geochronologic techniques to **date very young volcanic rocks and associated structures**.

e.g. seismic reflection imaging
of mafic sills

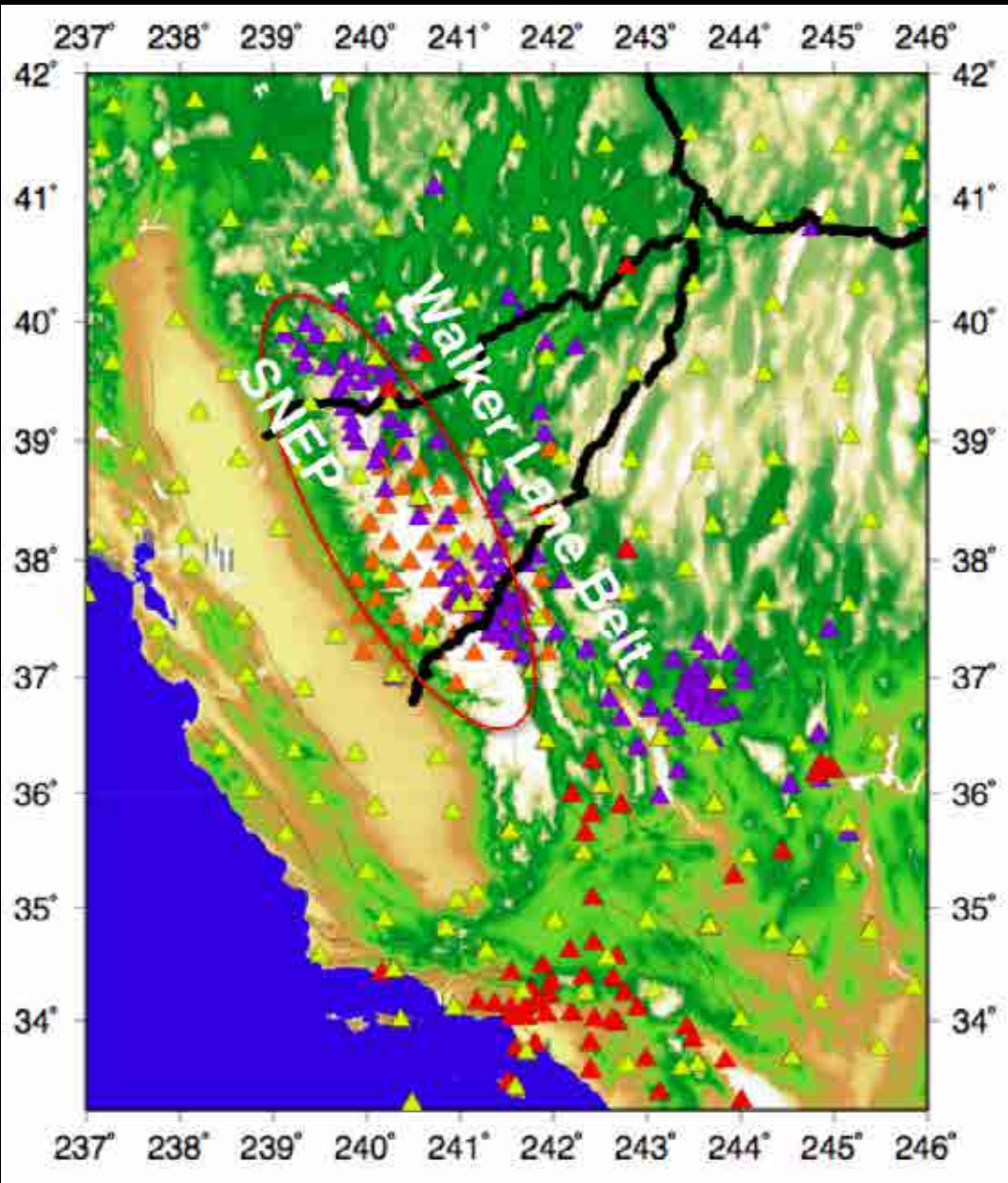
(7) Magmatism over past
>26 Ma - **can determine
longer-term fault motion
rates with dateable
volcanic stratigraphy.**



(8) Active tectonism - **real-time measurements** with GPS and seismic monitoring, **discovery of previously unmapped faults** with LIDAR, **dating of very young fault surfaces** with new cosmogenic techniques.



Attributes, continued:



(9) Portable array data already available from Sierra Nevada - need to

EXTEND THIS TO WALKER LANE RIFT SYSTEM

To determine the structure of the lower crust and upper mantle.

Attributes, continued:

(10) Intense geothermal exploration - **\$300 MILLION** from DOE and Industry in Nevada over next 5 years - **\$20 MILLION** in seismic alone -

societal impact (**green energy**) and leveraging from other funding agencies.



Coso
Geothermal
Field

MATRIX ITEM 1: Why does rifting occur - specifically, how do extensional forces exceed the yield strength of the lithosphere?

Determine relative importance of:

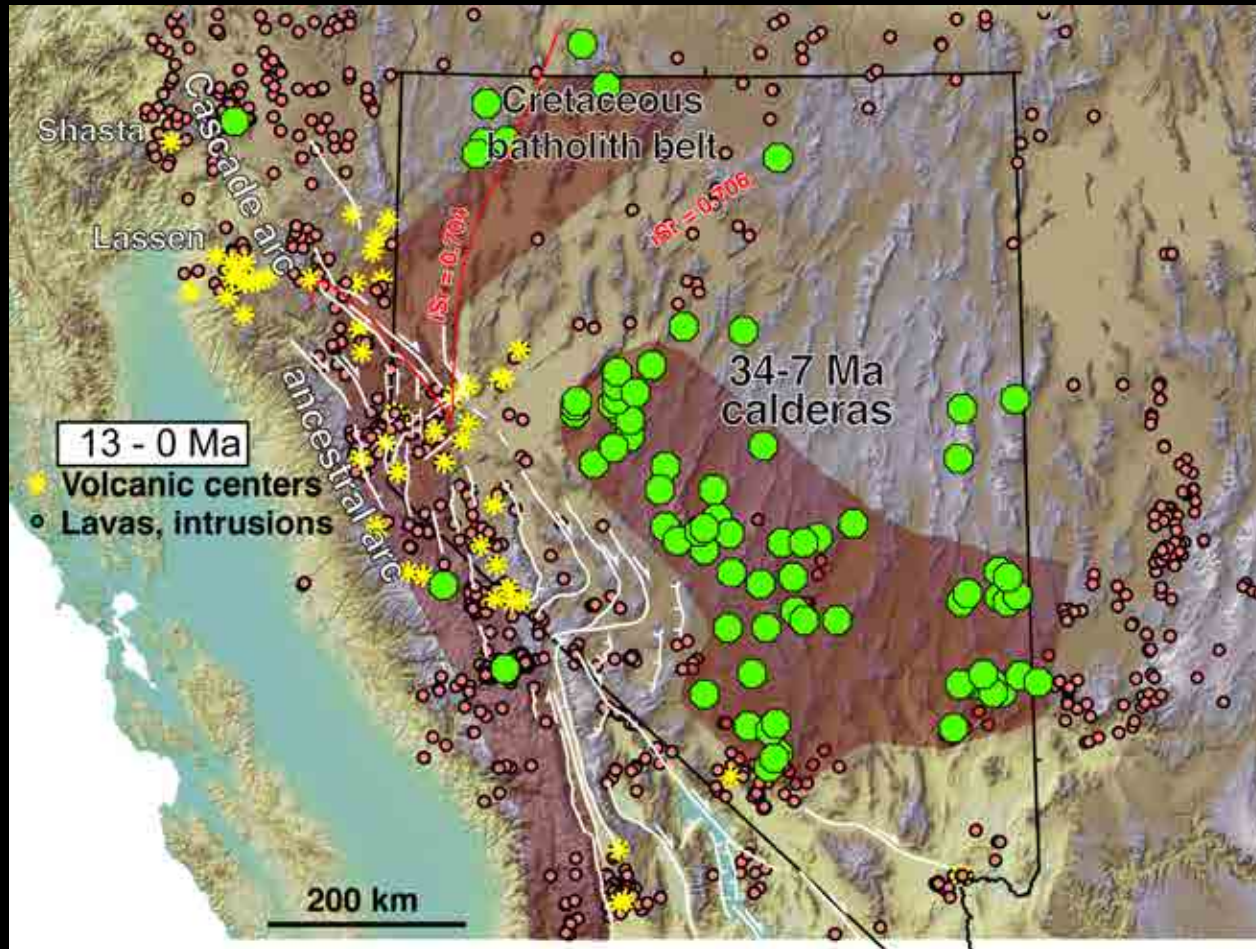
Regional forces:

1. Gravitational collapse of crust thickened in Sevier and Laramide orogenies (Paleocene)
2. Eocene to Miocene extension caused by slab rollback (Eocene to Miocene)
3. Change from subduction margin to transform margin
4. Basin and Range extension due to growth of the San Andreas fault boundary (16 Ma-present)
5. Onset of transtension at ~12 Ma (due to change in Pacific plate motion)

Local body forces:

1. Positive and negative contributions of mantle buoyancy (e.g. root delamination, asthenospheric upwelling)

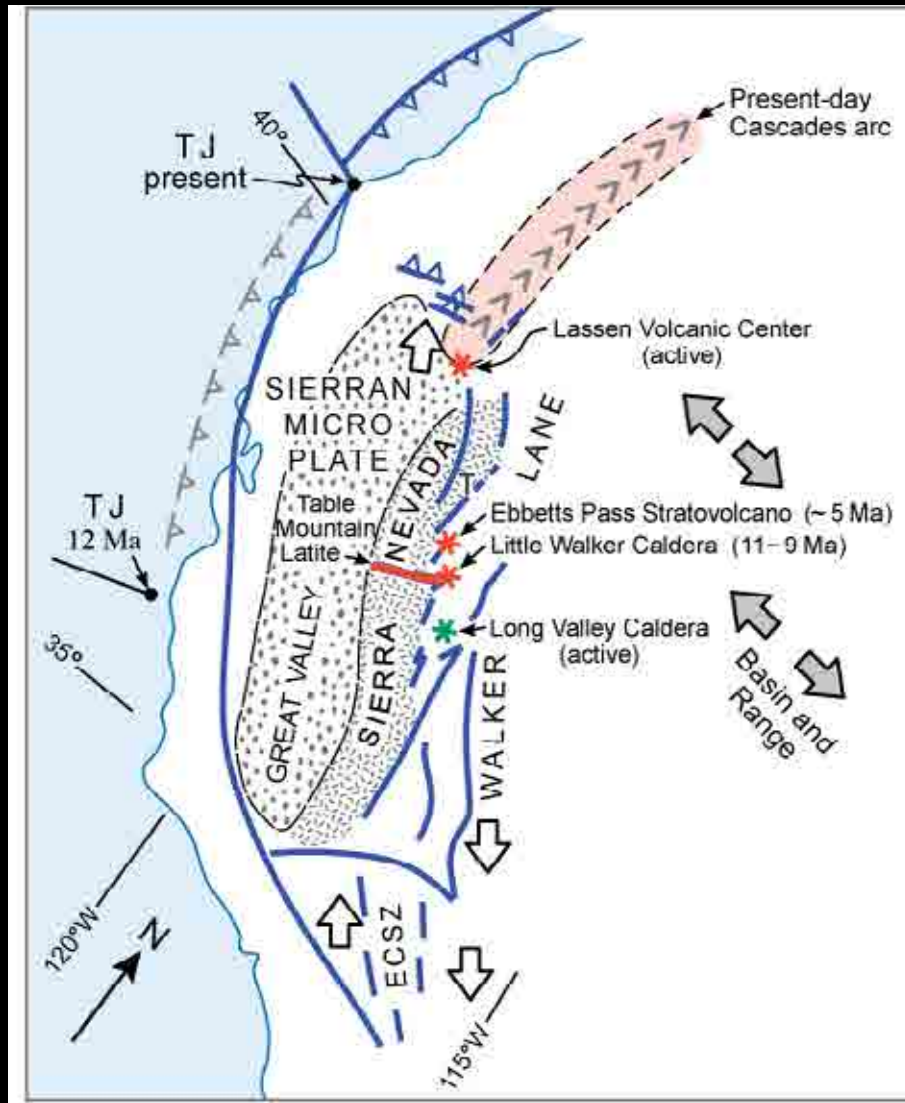
HYPOTHESIS: Walker Lane rift system formed between
Cretaceous and mid-Cenozoic caldera batholithic belts (browns)
- and cut through the Cretaceous batholith belt along the
Ancestral Cascade Arc axis (orange dots).



Link with
Cascadia

We are still
discovering
major volcanic
centers and
determining
their tectonic
controls.....

...but large volcanic centers appear to be located on transtensional stepovers:



ACTIVE:

Lassen arc stratovolcanic center

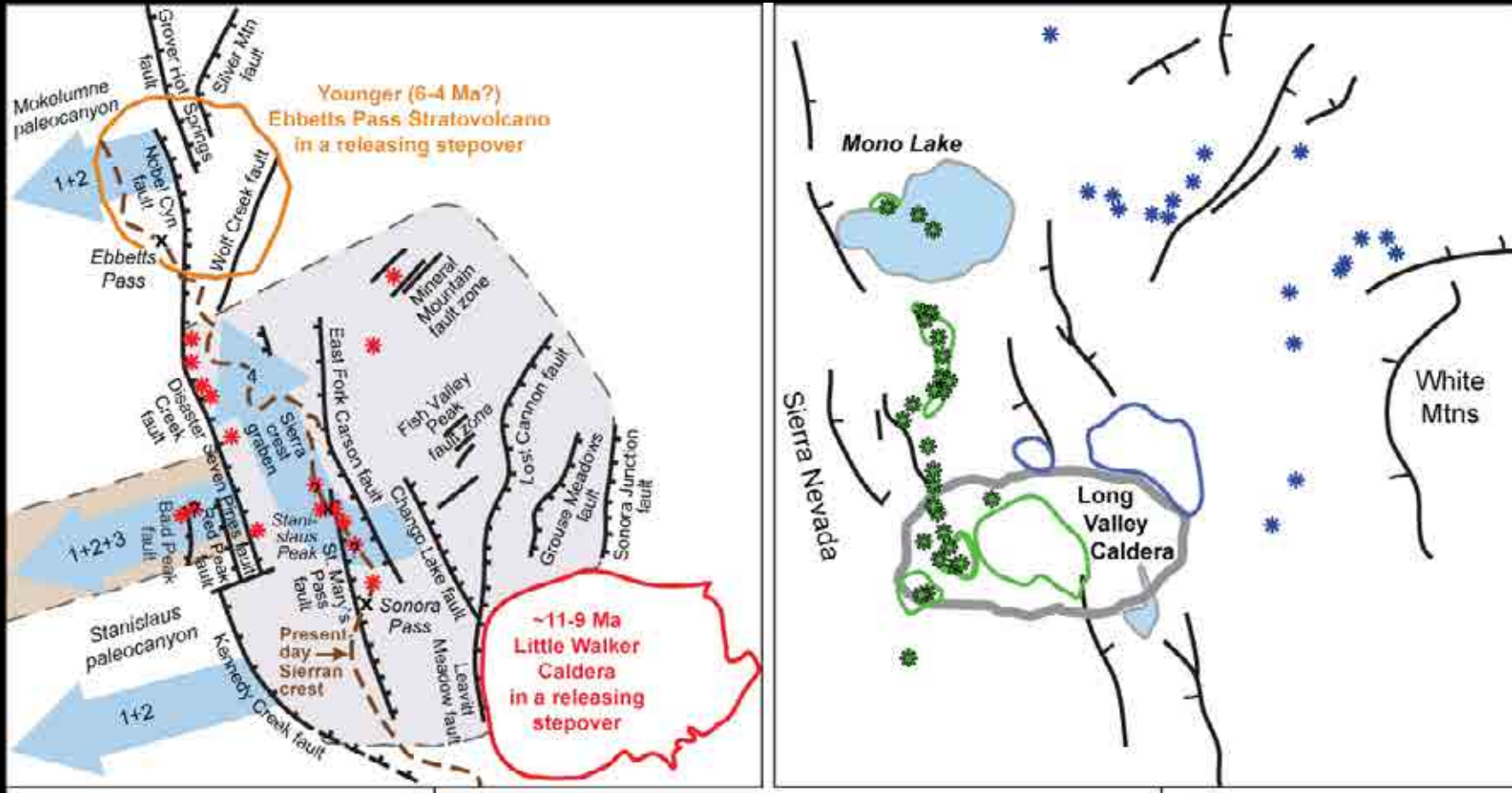
Long Valley rift caldera

CENTERS NEWLY DISCOVERED:

Sierra Crest Graben Andesite Flood Lava Complex beneath Little Walker arc caldera (11 - 9 Ma)

Ebbetts Pass arc stratovolcanic center (5 Ma)

Miocene Little Walker Caldera (left) built atop grabens filled with fissure-fed flood andesite (lavender)
 shown at same scale as the active Long Valley caldera (right)



NNW-SSE faults right transtensional, ENE-WSW faults left transtensional. Stars = vents.

STRATIGRAPHIC EXPRESSION OF RIFT INITIATION

at ~ 12 Ma:

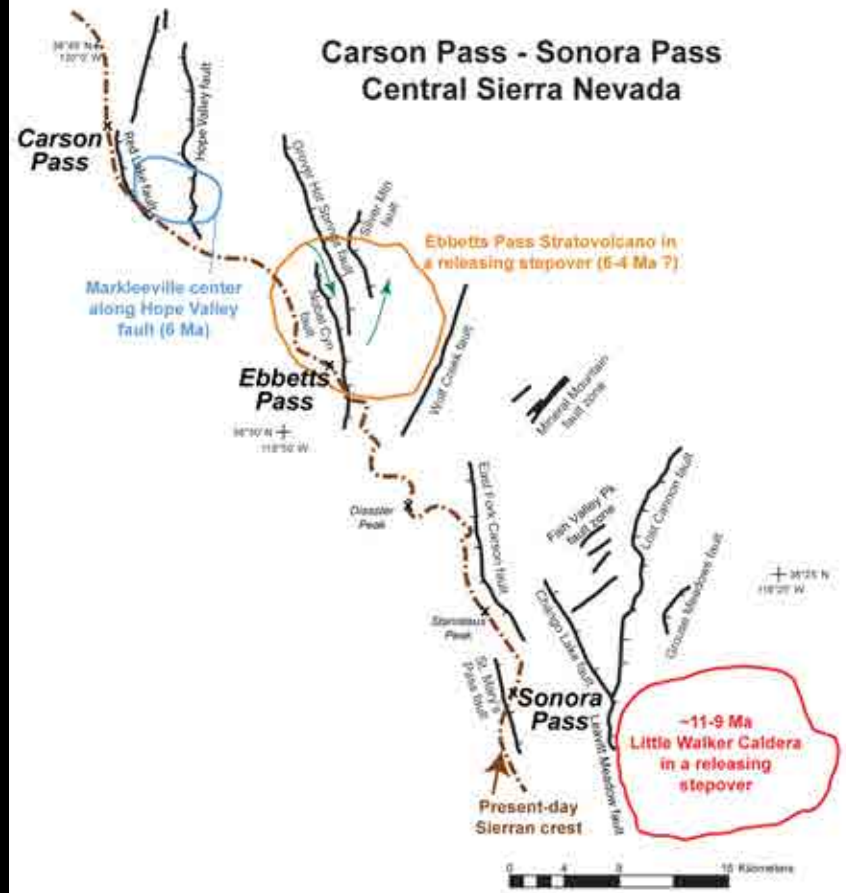
Extreme effusive
eruptions of “flood
andesite”

and

unusually large-
volume, widespread
landslide deposits

(Busby et al. in review)





Another major volcanic center sited on transtensional stepover:

Newly-discovered

5 Ma

Ebbetts Pass Stratovolcano

(Busby et al., unpublished data)



MATRIX ITEM 3: What controls the architecture of rifts and rift margins? **SEDIMENTARY BASIN Outcrop and Lake Studies**

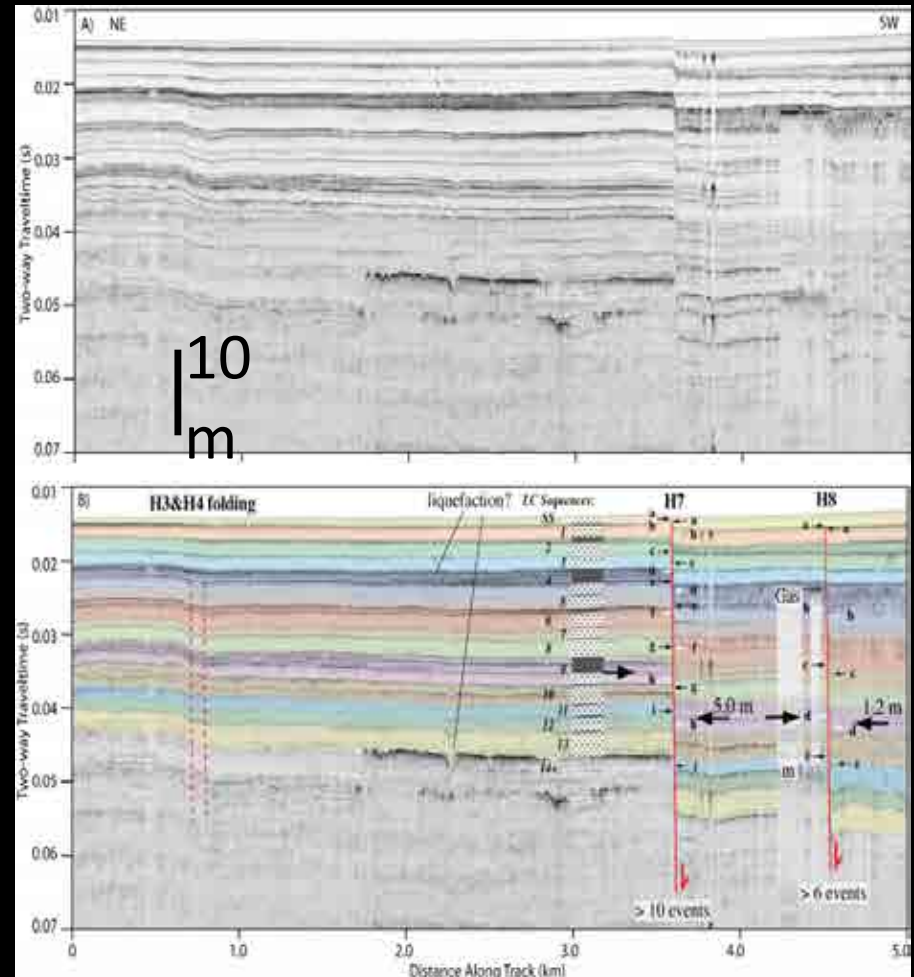
Great variety in:

drainage area sizes/bedrock geology/erosional processes

basin ages, sizes and structures, sedimentation rates, fault movement histories and paleoseismicity records

hazards record from landslides and seiches

paleoclimate record from outcrop and drill core



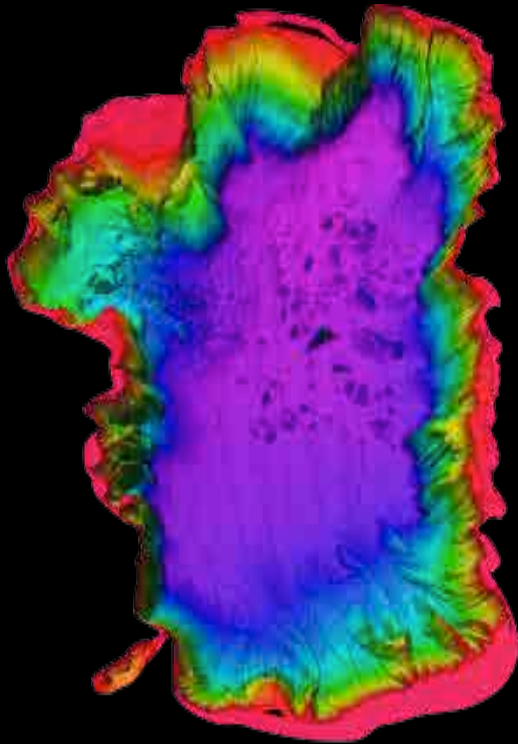
Salton Sea

Lake Tahoe/Pyramid Lake to Salton Sea to Gulf of CA:

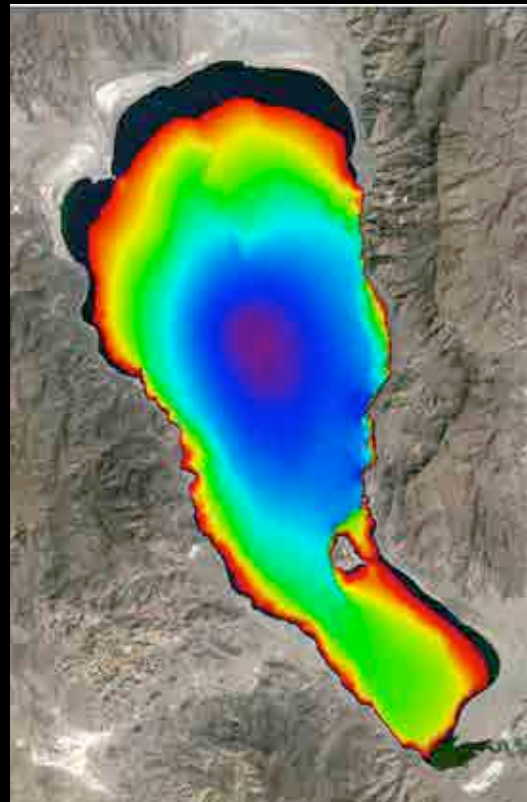
High resolution studies of **modern lakes** and the **dissected deposits of Pleistocene pluvial lakes**

Influence of Sedimentation on Rift Development

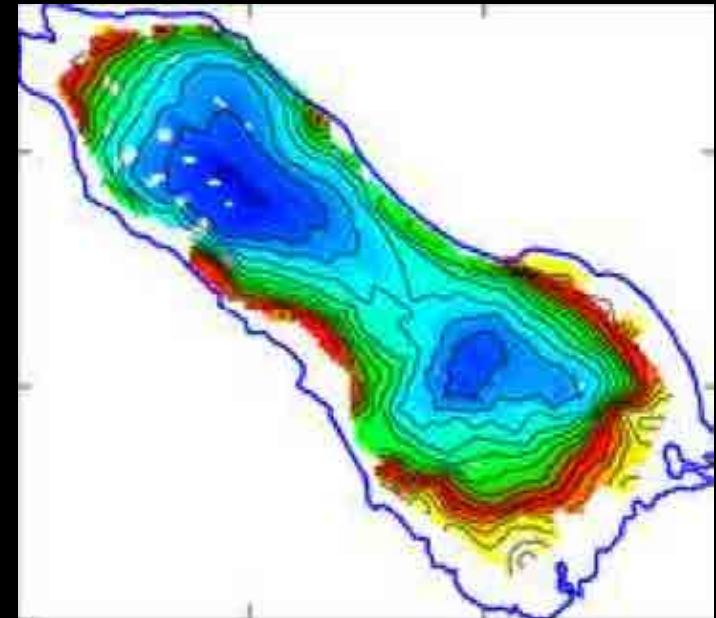
Lake Tahoe



Pyramid Lake



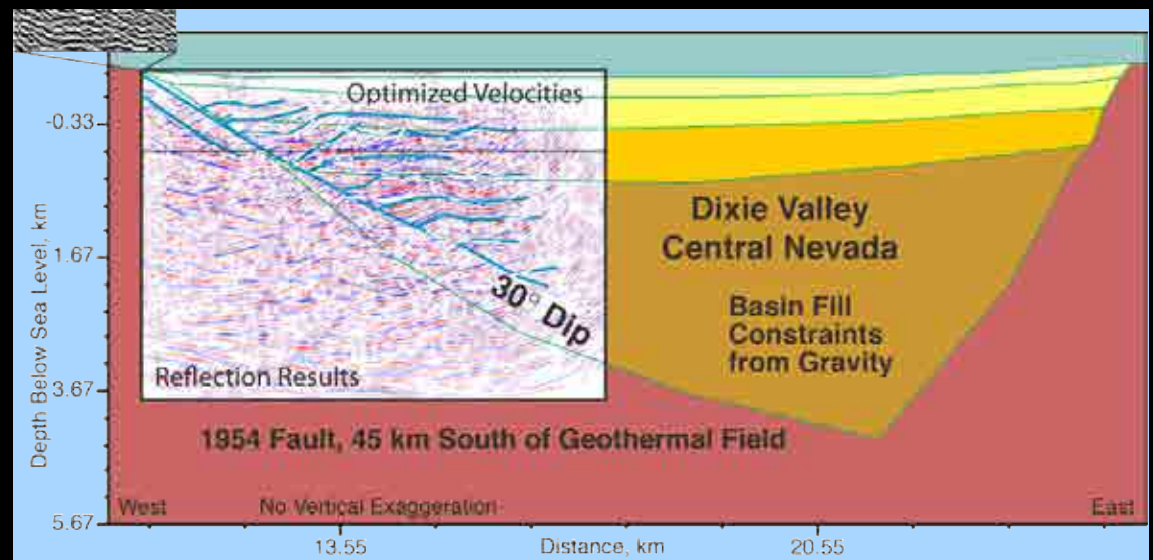
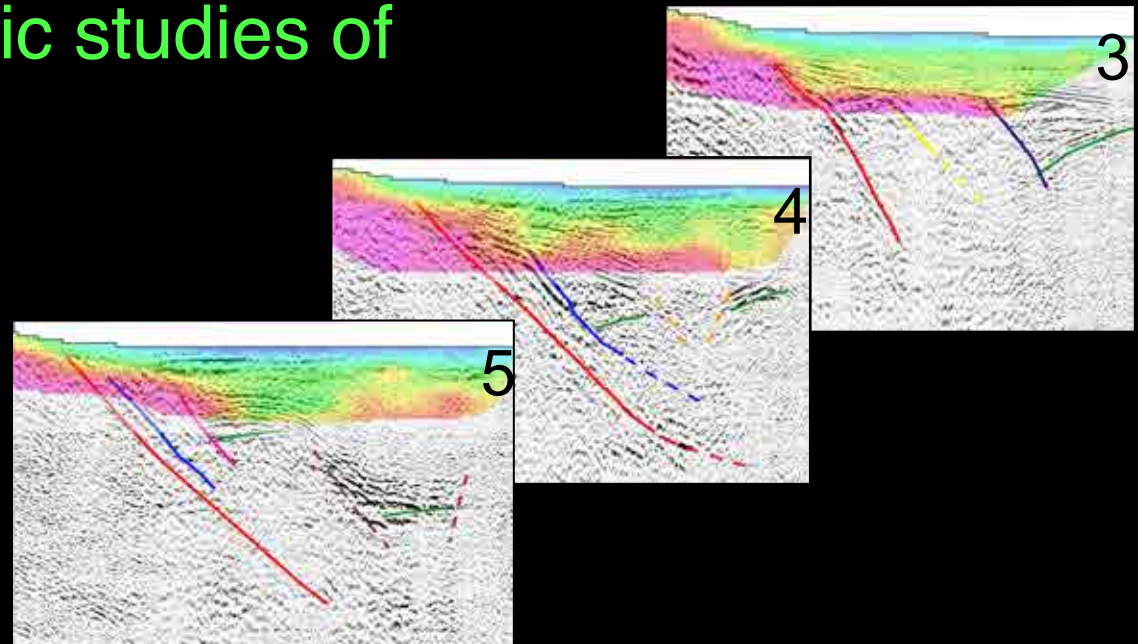
Salton Sea



Examples of Seismic studies of basin architecture

**Only the
range-front
fault breaks
the surface**

A fault that has
been active
at low angle
since 23 Ma



WE PROPOSE A U.S. RIFT PROGRAM

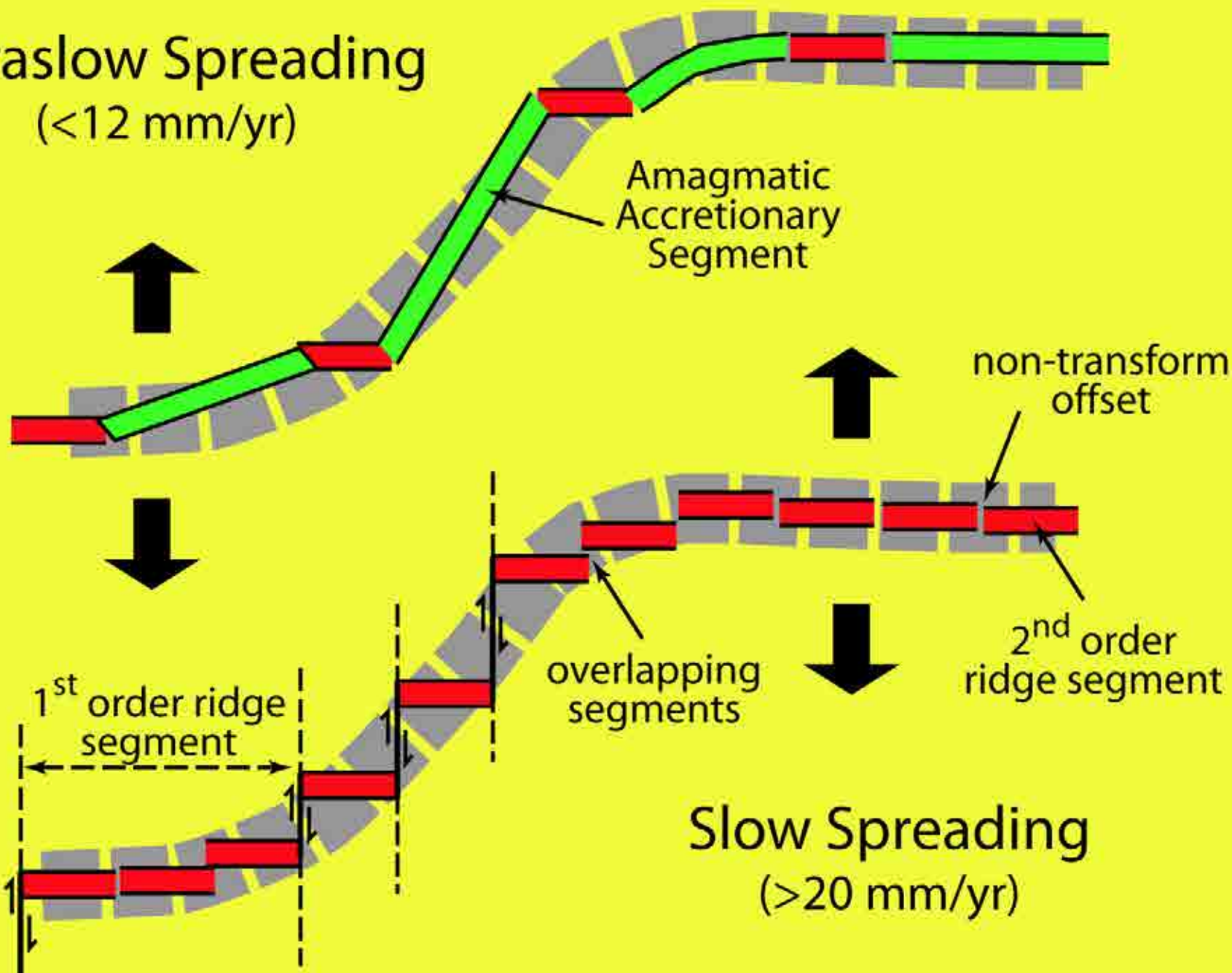
to interface with Cascadia

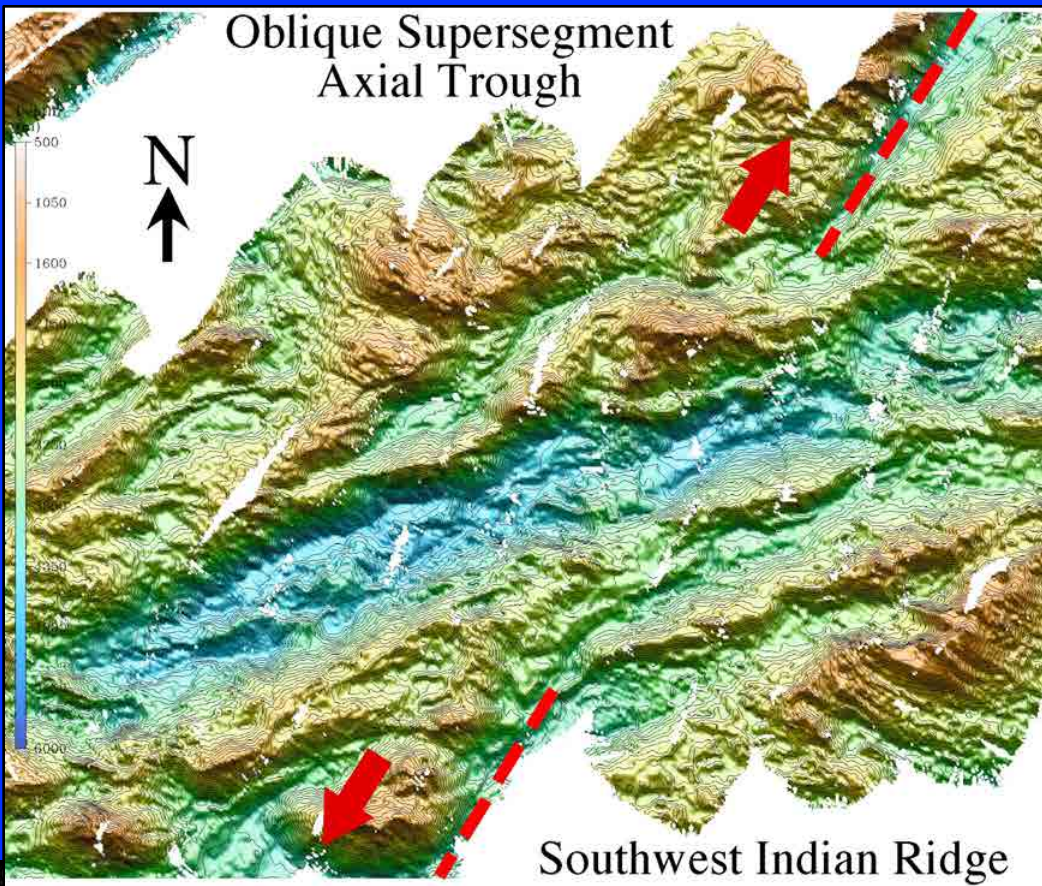
By comparing rift initiation in the Walker Lane
and Rio Grande rift systems

with the successful Gulf of CA rift

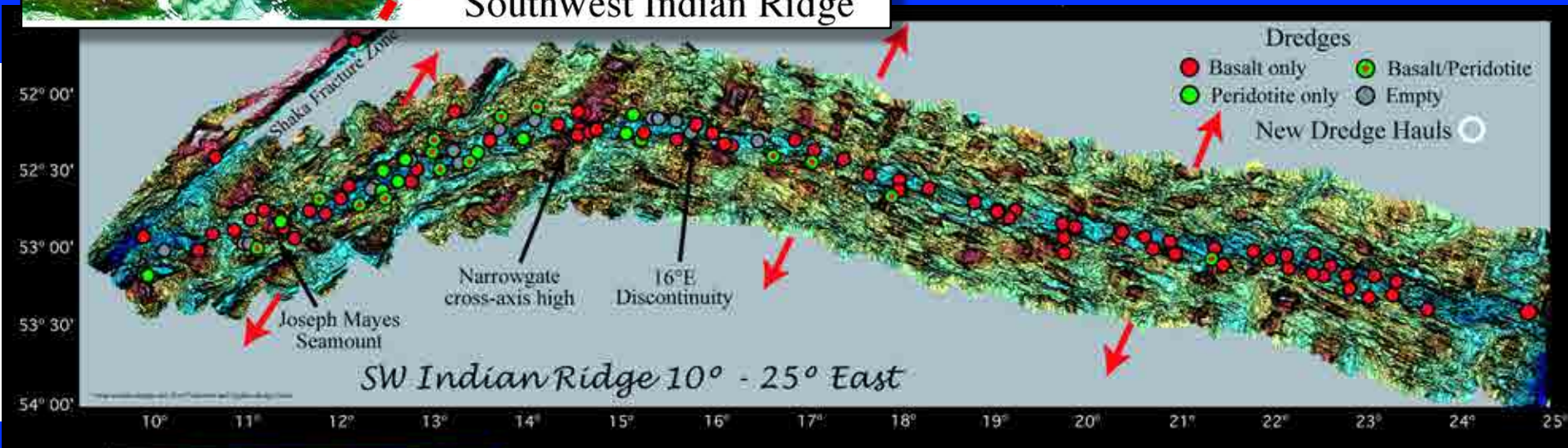
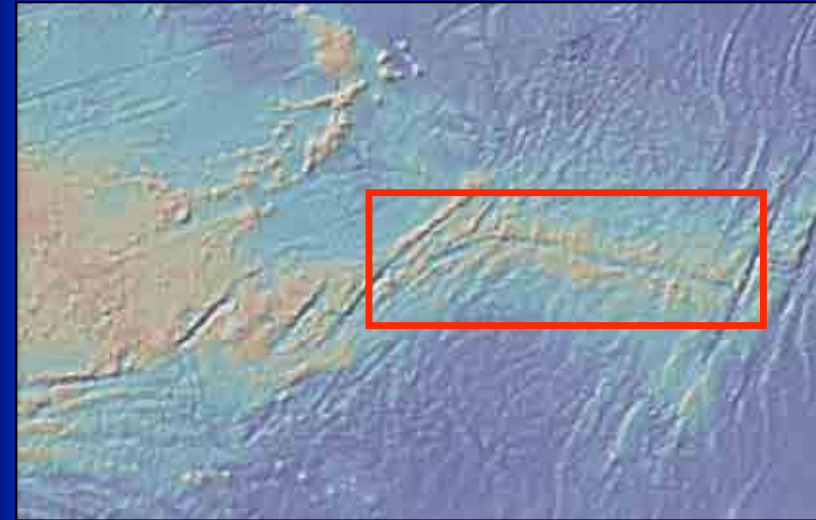
and the Mississippi Valley failed rift and the
“mature” East Coast rift

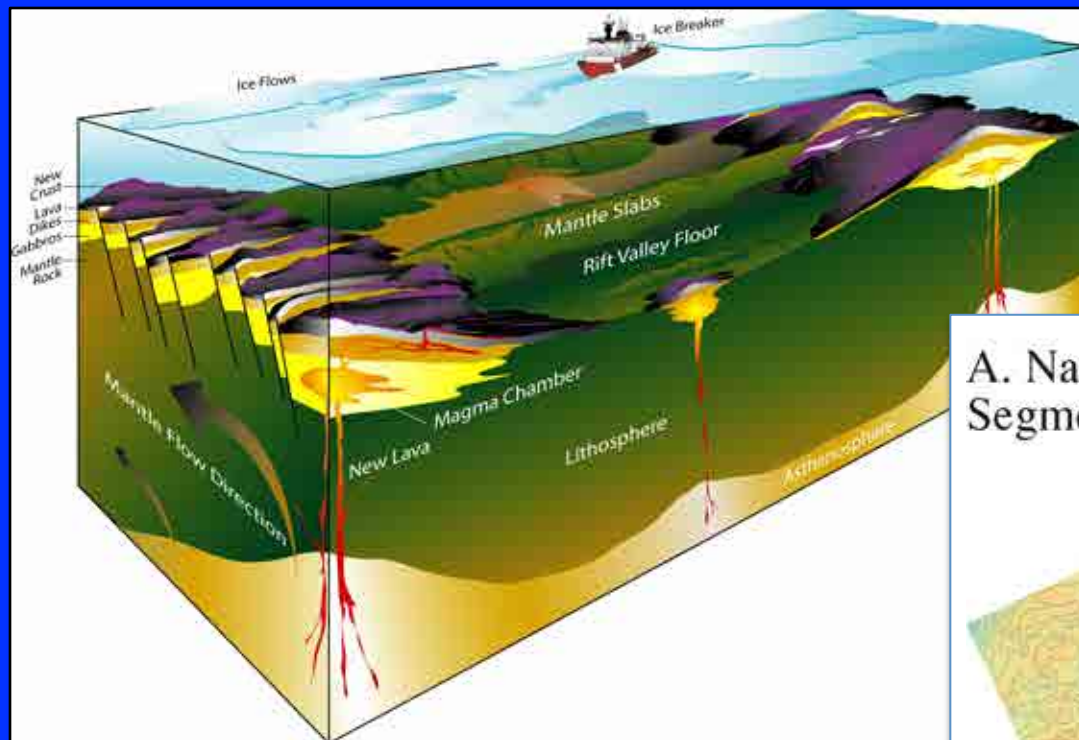
Ultraslow Spreading (<12 mm/yr)



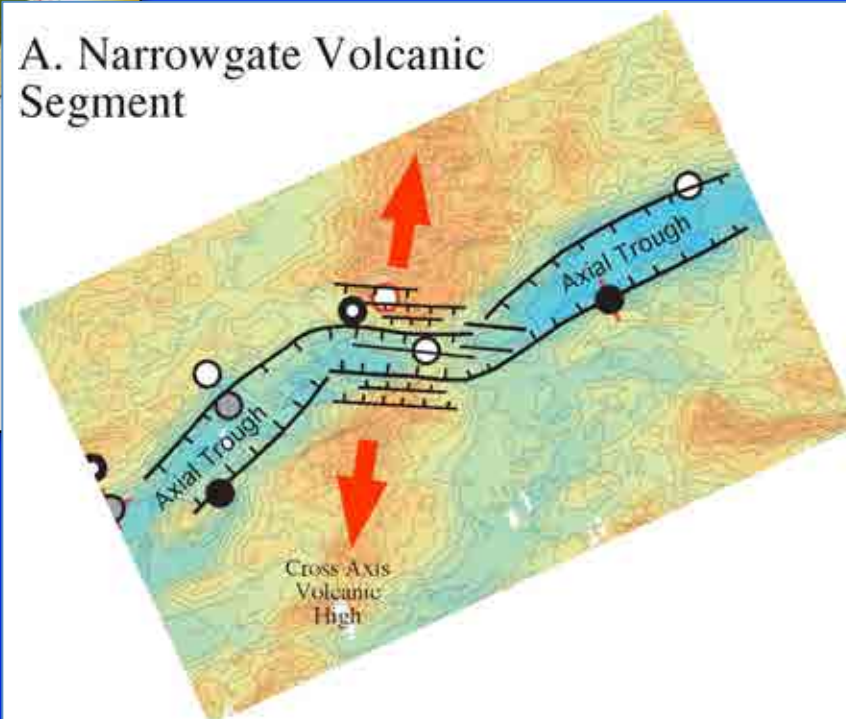


Ultralow Spreading SW Indian Ridge

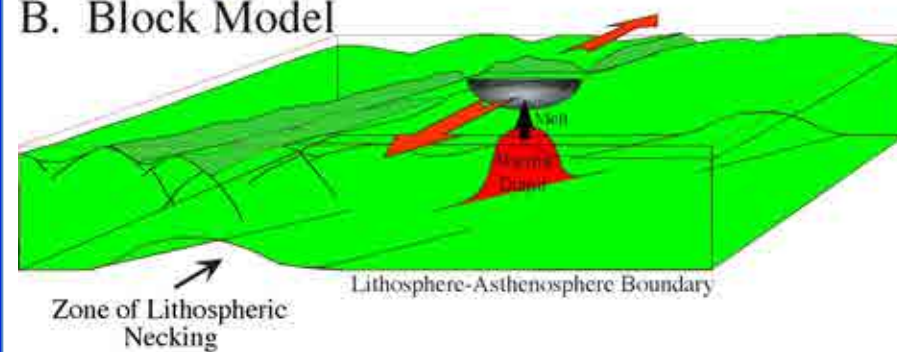




A. Narrowgate Volcanic Segment



B. Block Model

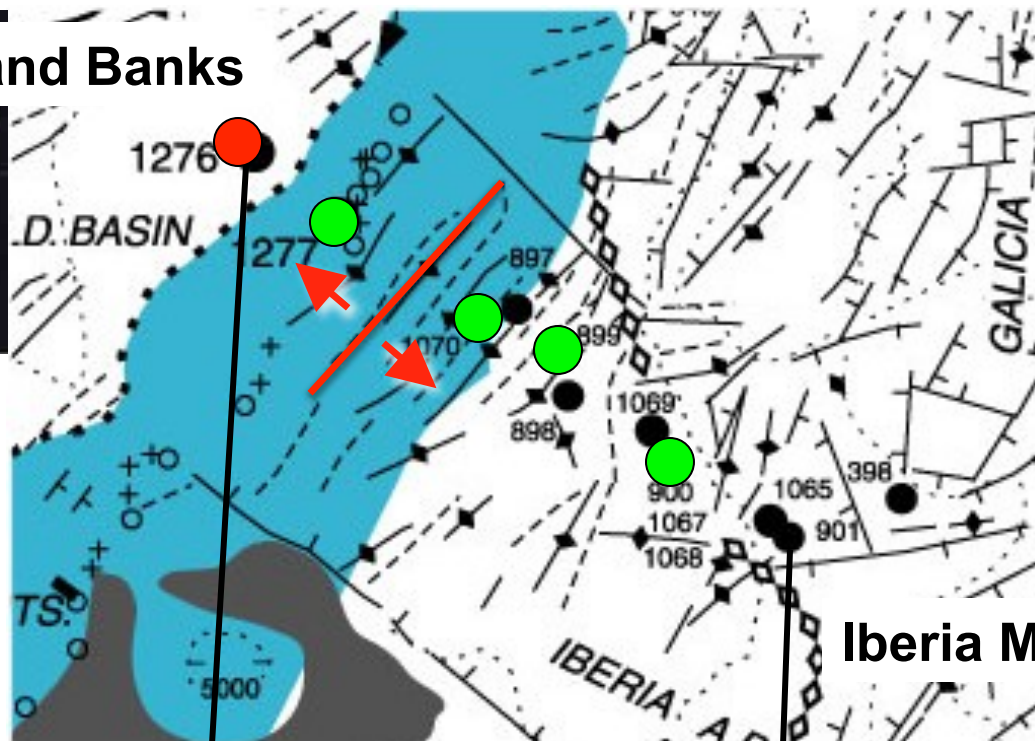




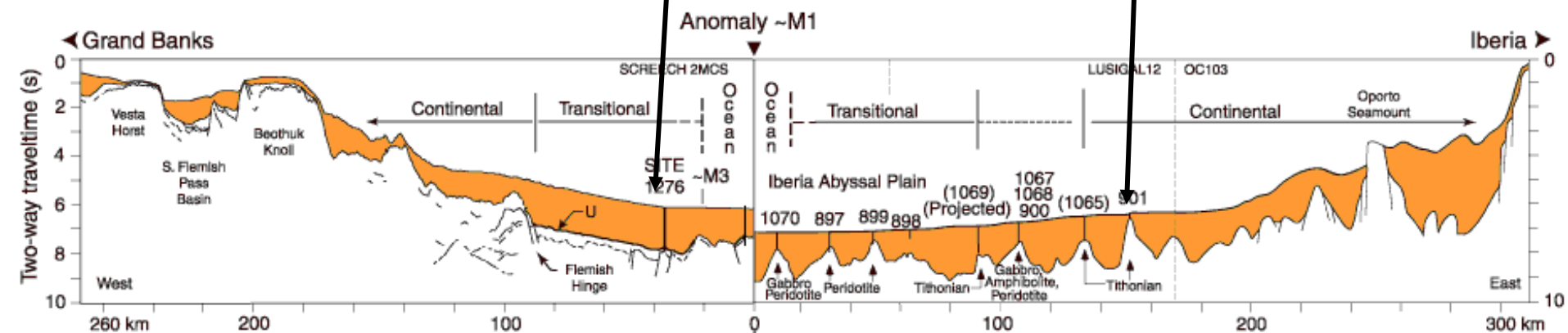
Grand Banks

Drilling Legs:
149, 173, 210

120 Ma,
Mid-Cretaceous



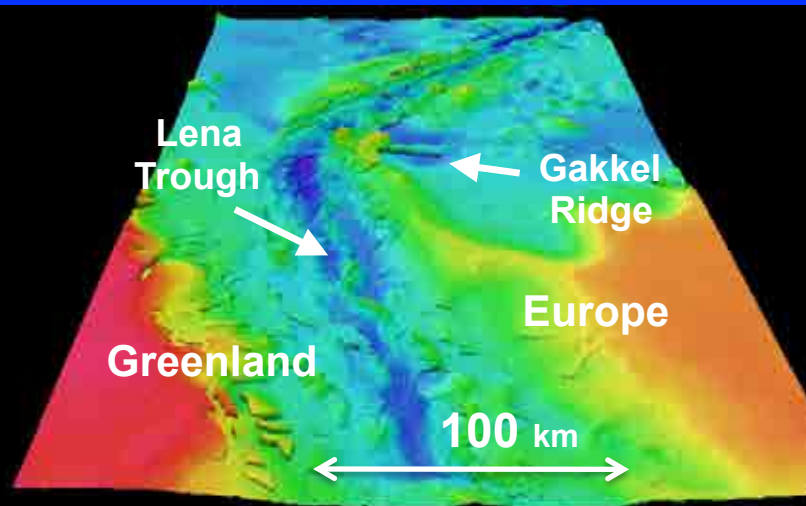
Iberia Margin



Gabbroic-Ultramafic!
"Amagmatic" extension

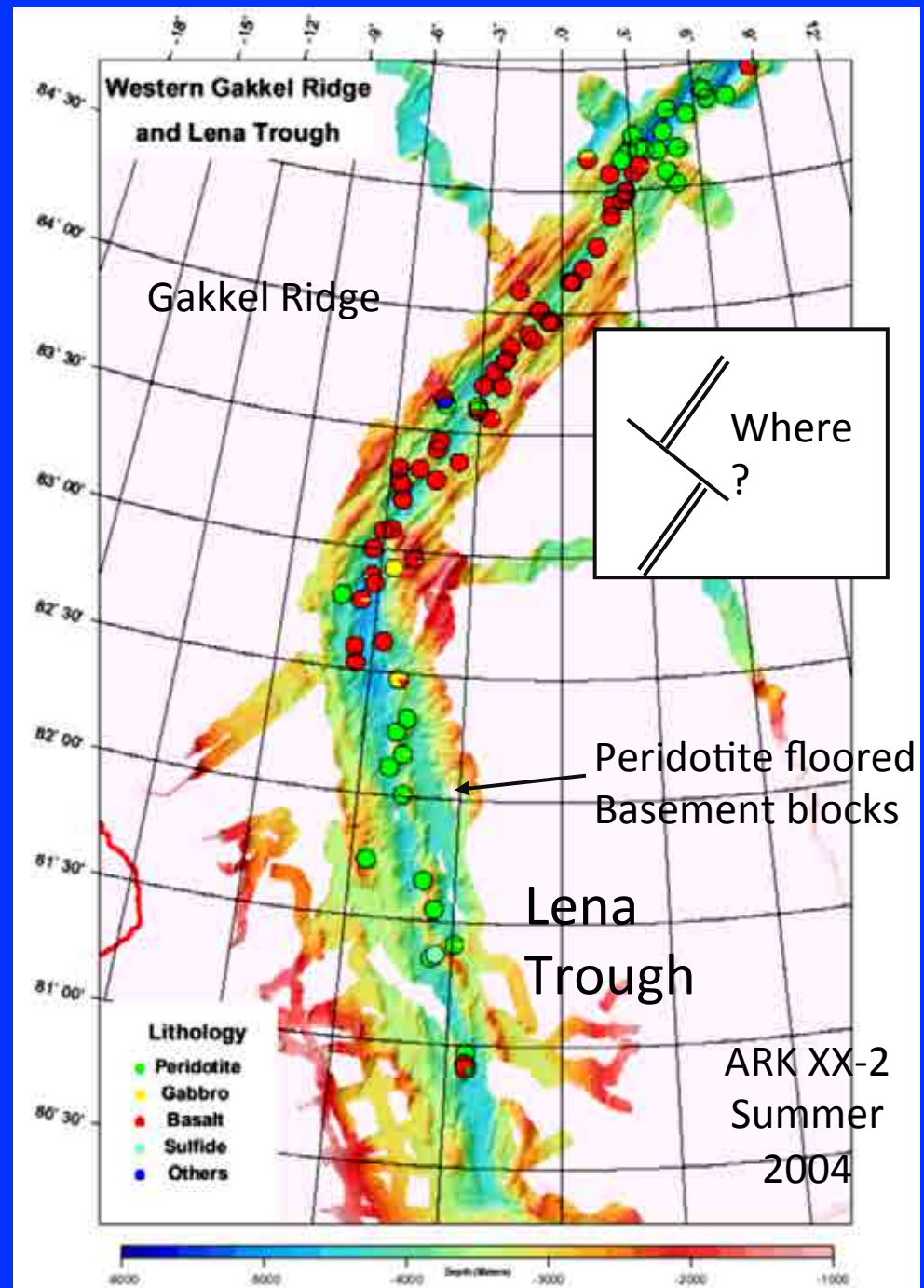
Grand Banks

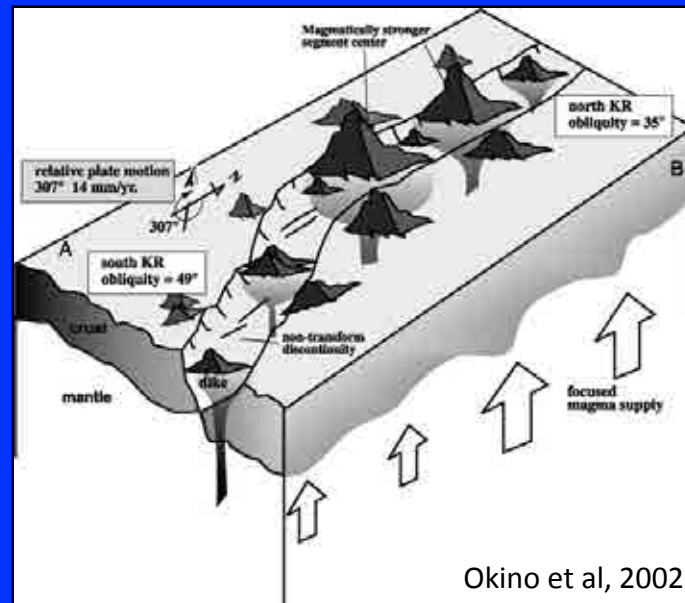
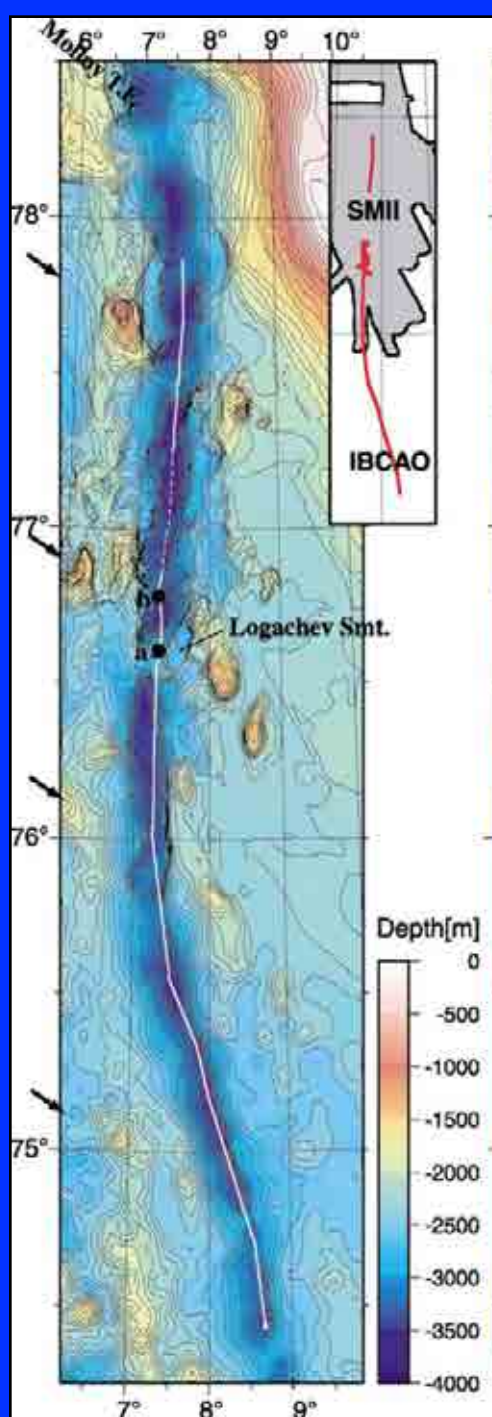
Iberia



Lena Trough – a modern example of the ocean – continent transition

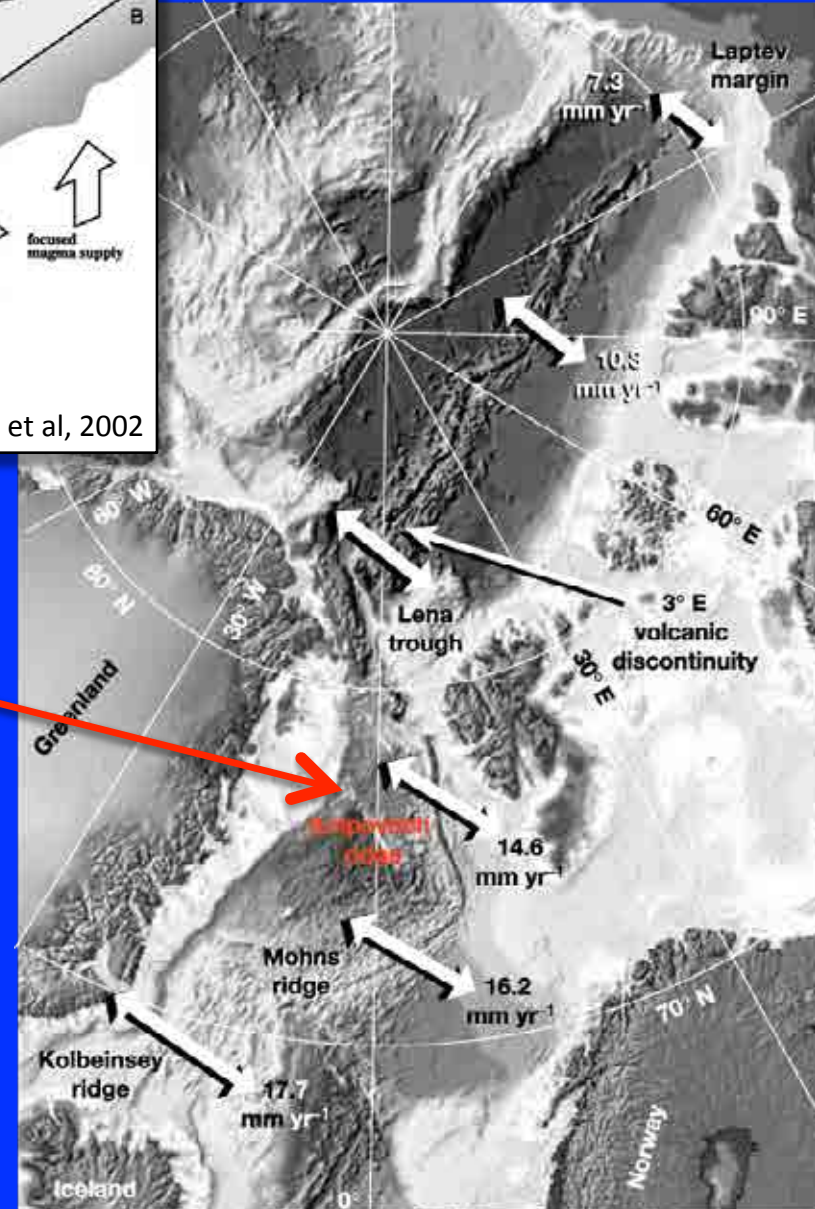
Amagmatic accretionary segments seen at ultra slow ridges spreading at an ESR < 12 mm/yr are likely the characteristic plate boundary structure of non-volcanic rifted margins.

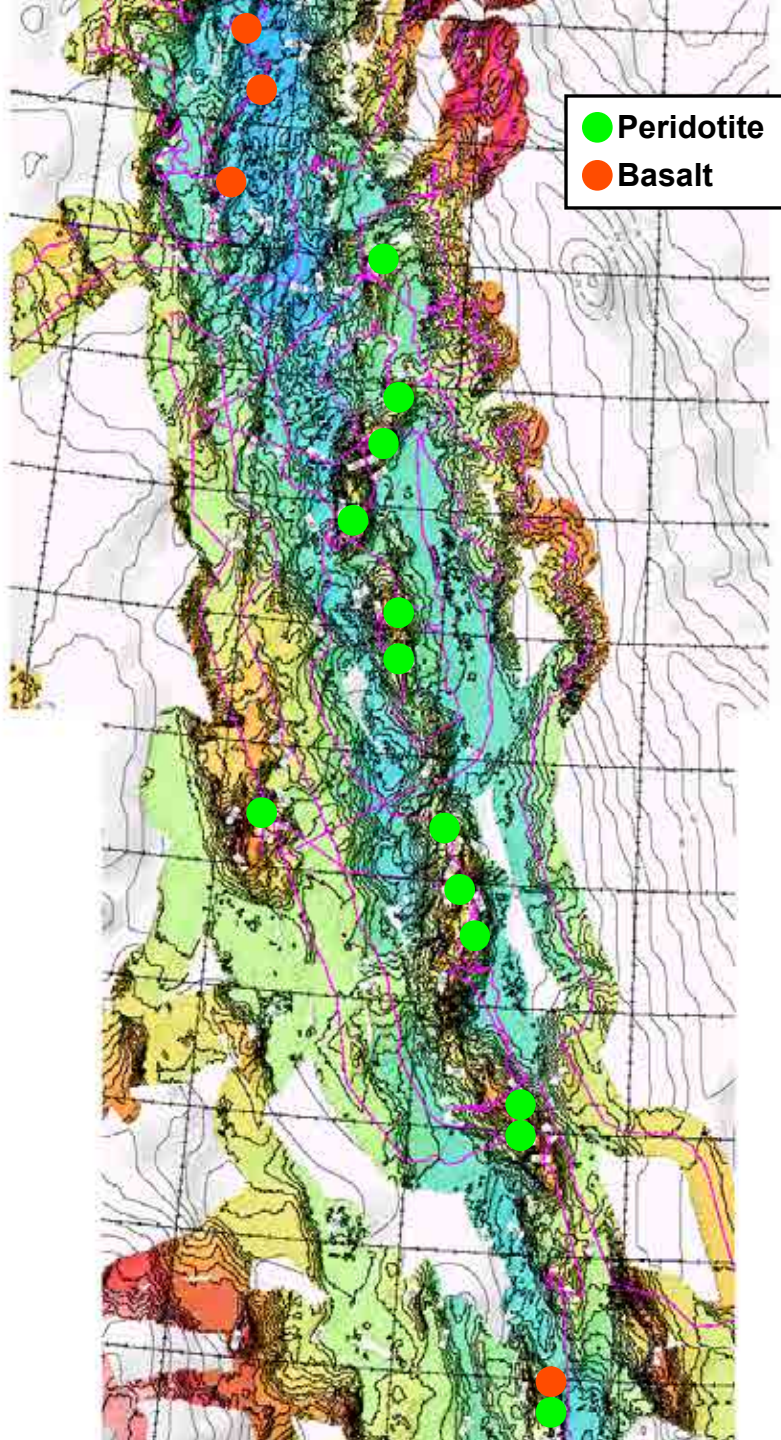




Okino et al, 2002

Knipovitch Ridge

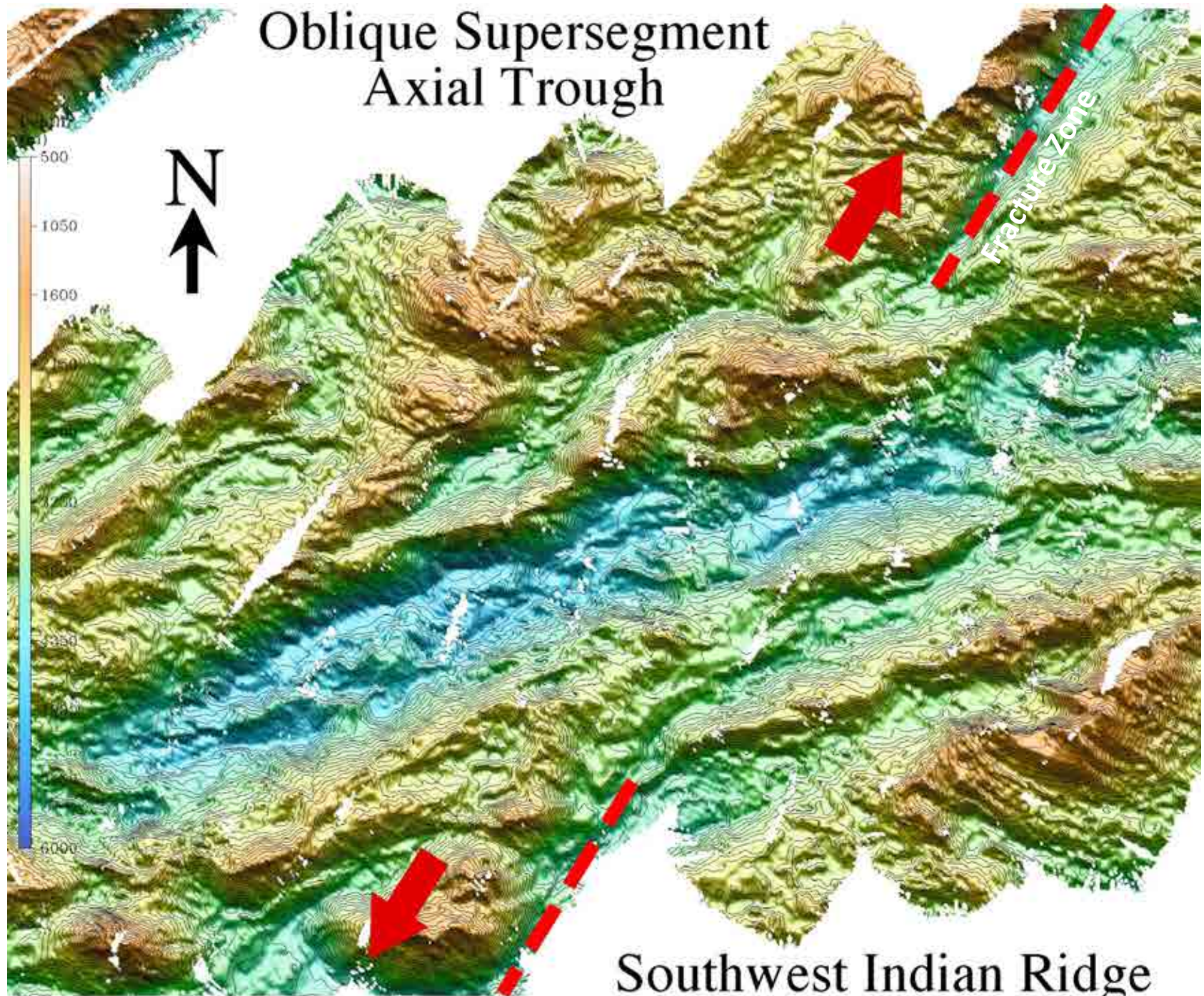




“Amagmatic”!

- Mostly ultramafic outcrop
- Little or no volcanism
- Volcanism at segment ends, not center (!)
- Very fertile mantle compositions
- Melting very variable
- Melt production, stagnation, etc....

Oblique Supersegment Axial Trough



Southwest Indian Ridge

Rapid Response efforts in rifts – volcanic eruptions, dike intrusions, and large earthquakes – Tijuana EQ, EAR

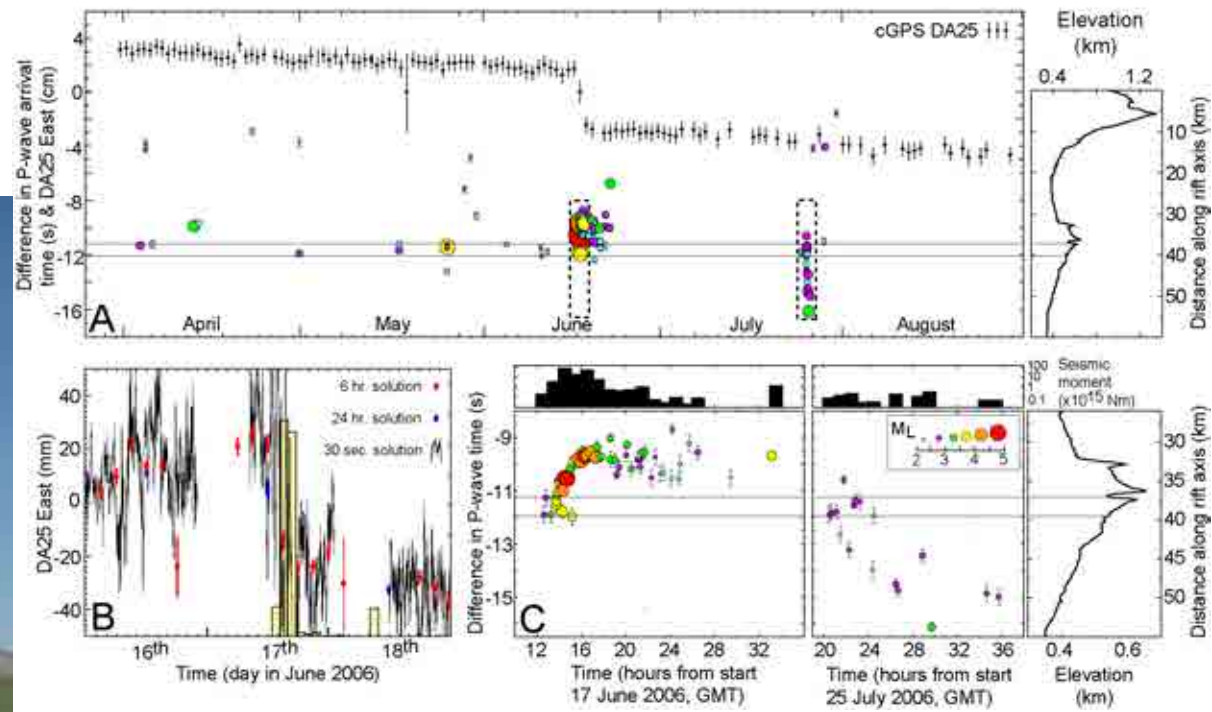
Dike intrusions in particular characterized by $M_L < 5$ EQ's – generally not considered for rapid response– dike dimensions; track magma movement; aftershocks illuminate fault/dike dimensions

Space-based geodesy (GPS, InSAR) – measure stress diffusion rates and vertical/horizontal crustal movements – determine mantle viscosity to rapid opening .

CONSTRAIN RHEOLOGY, DIMENSIONS, ASEISMIC/SEISMIC STRAIN, MAGMA-VOLATILE FLUXES



Keir et al., figure 3



Submarine Landslides

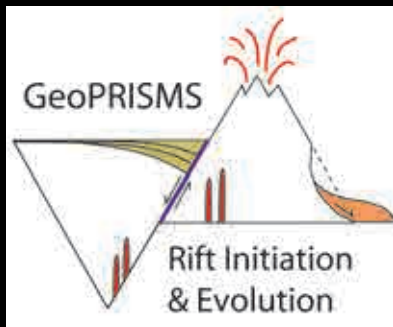
Rapid Response Opportunities

(Best data would be pre-, during-, post-failure but...)

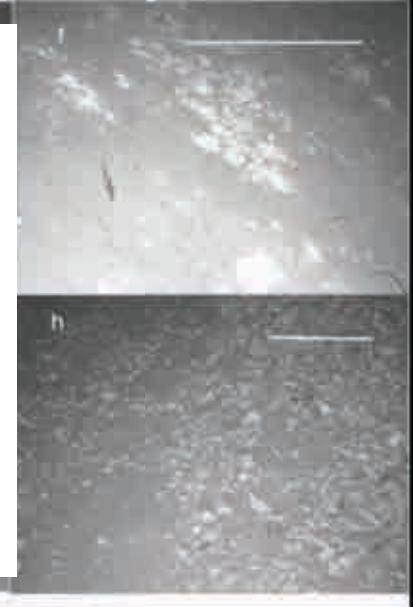
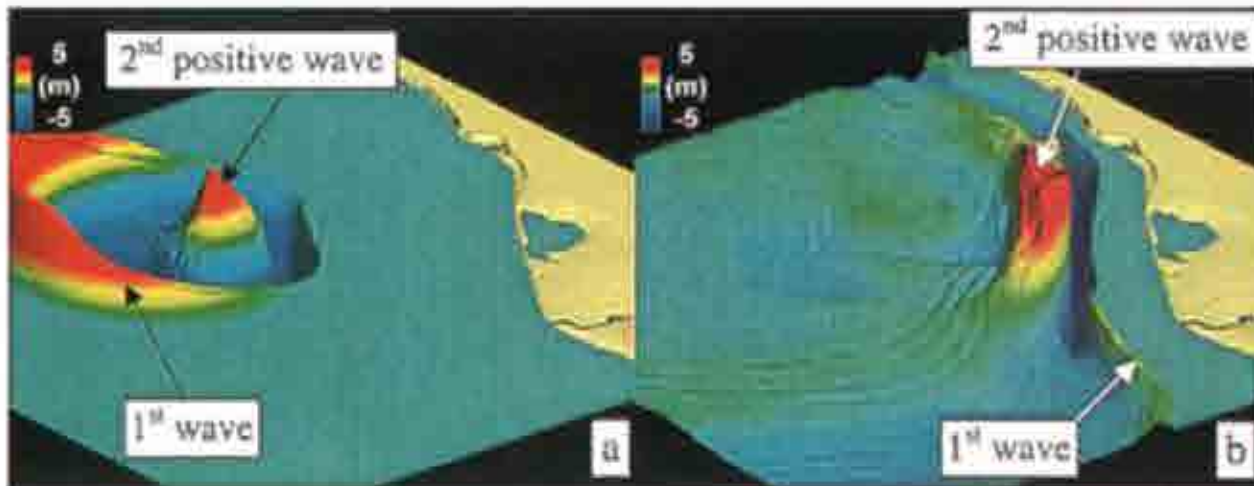
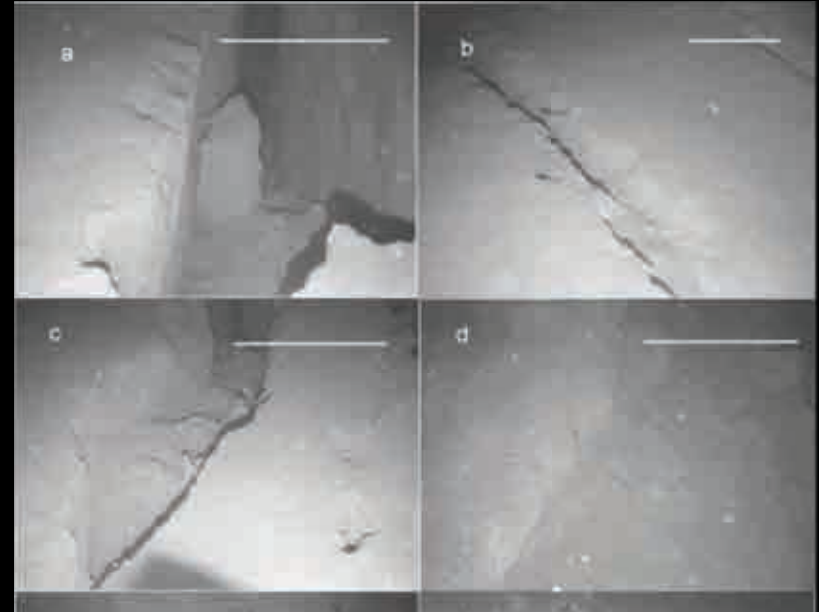
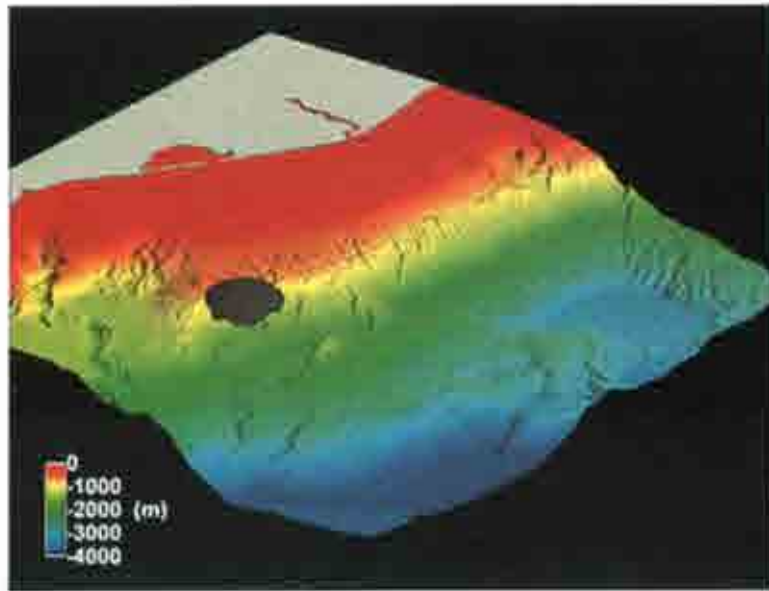
M. Hornbach, P. Flemings (Jackson School of Geosciences)

R. Harris (Oregon State University)

B. Dugan (Rice University)



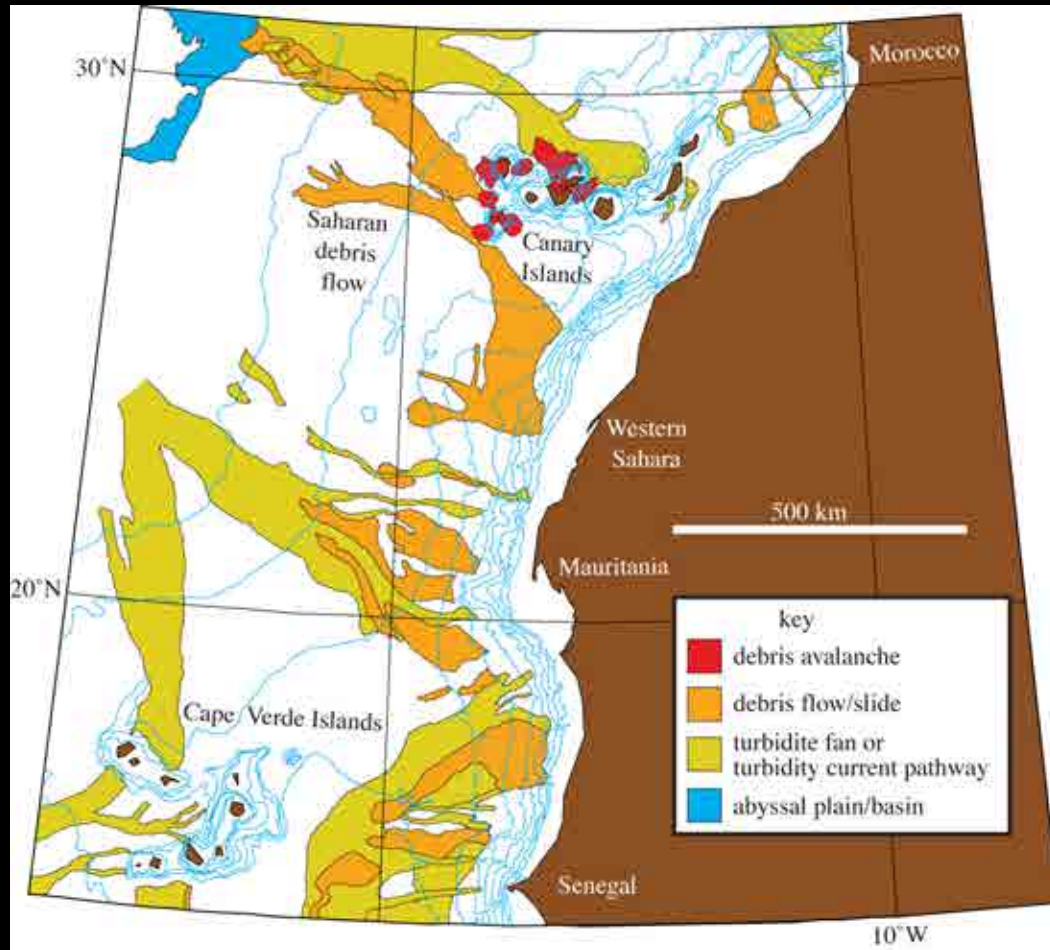
17 July 1998 Papua New Guinea Tsunami



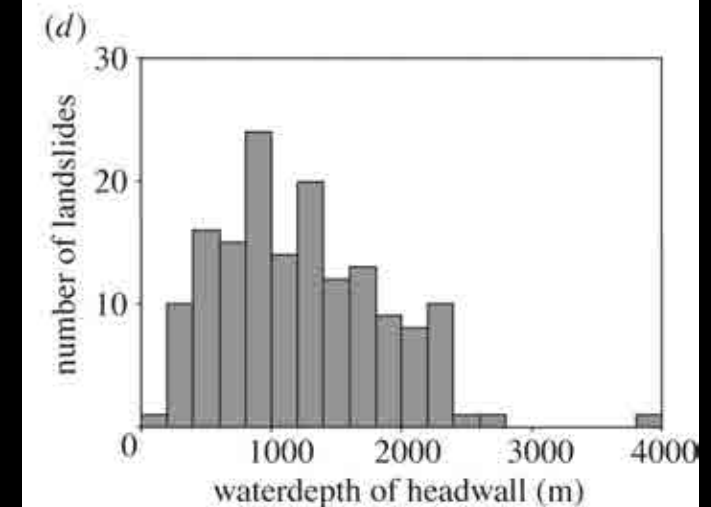
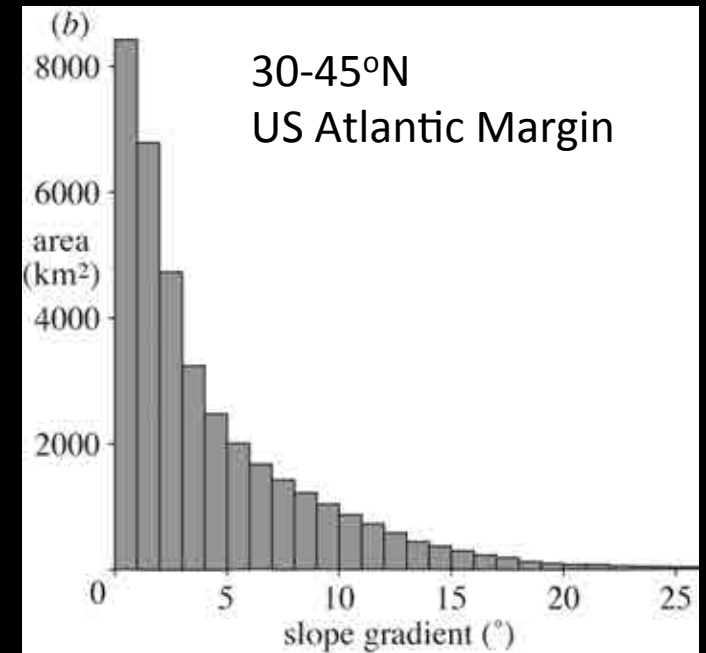
Heinrich et al., 2000

Tappin et al., 2002

Single Events, Margin-Scale Impacts



Masson et al., 2006

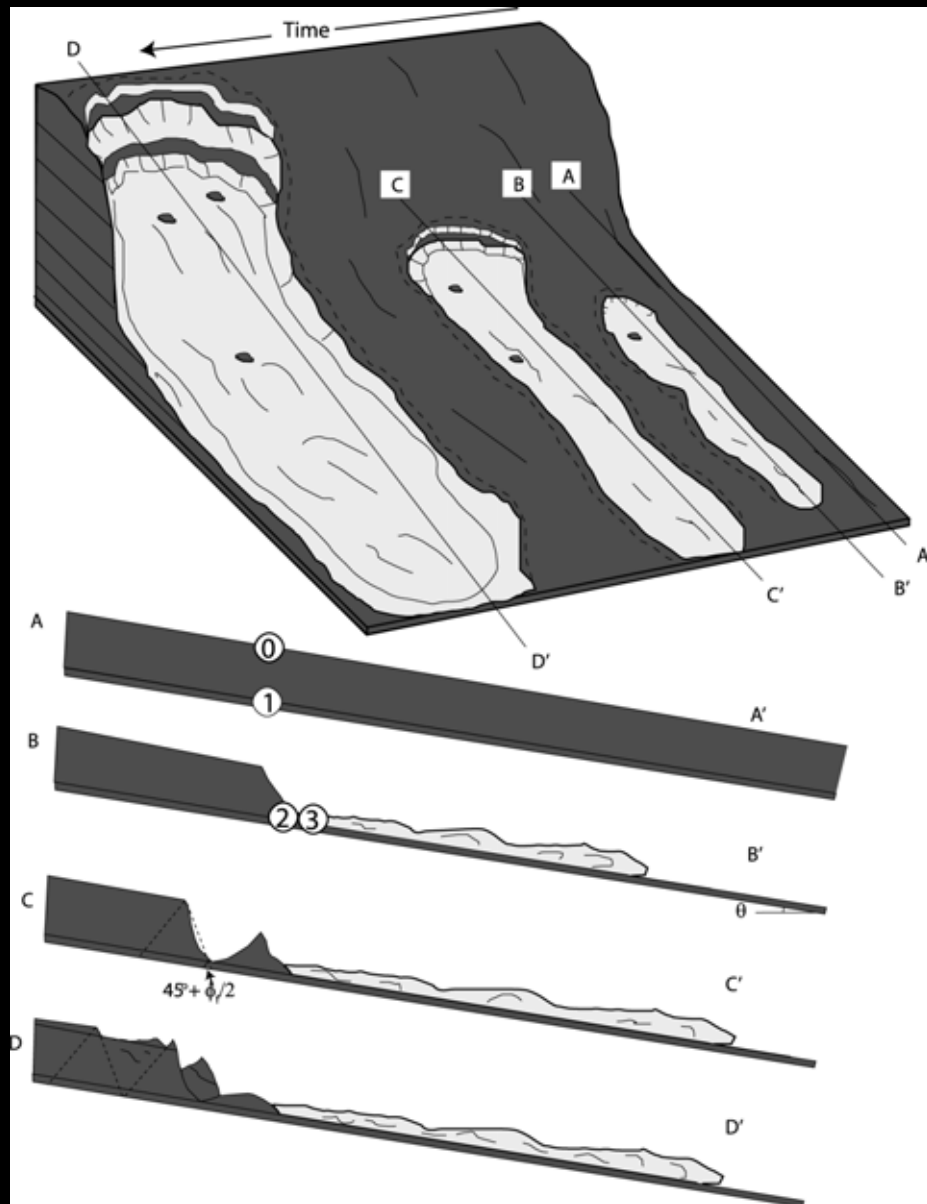


Why submarine landslides?

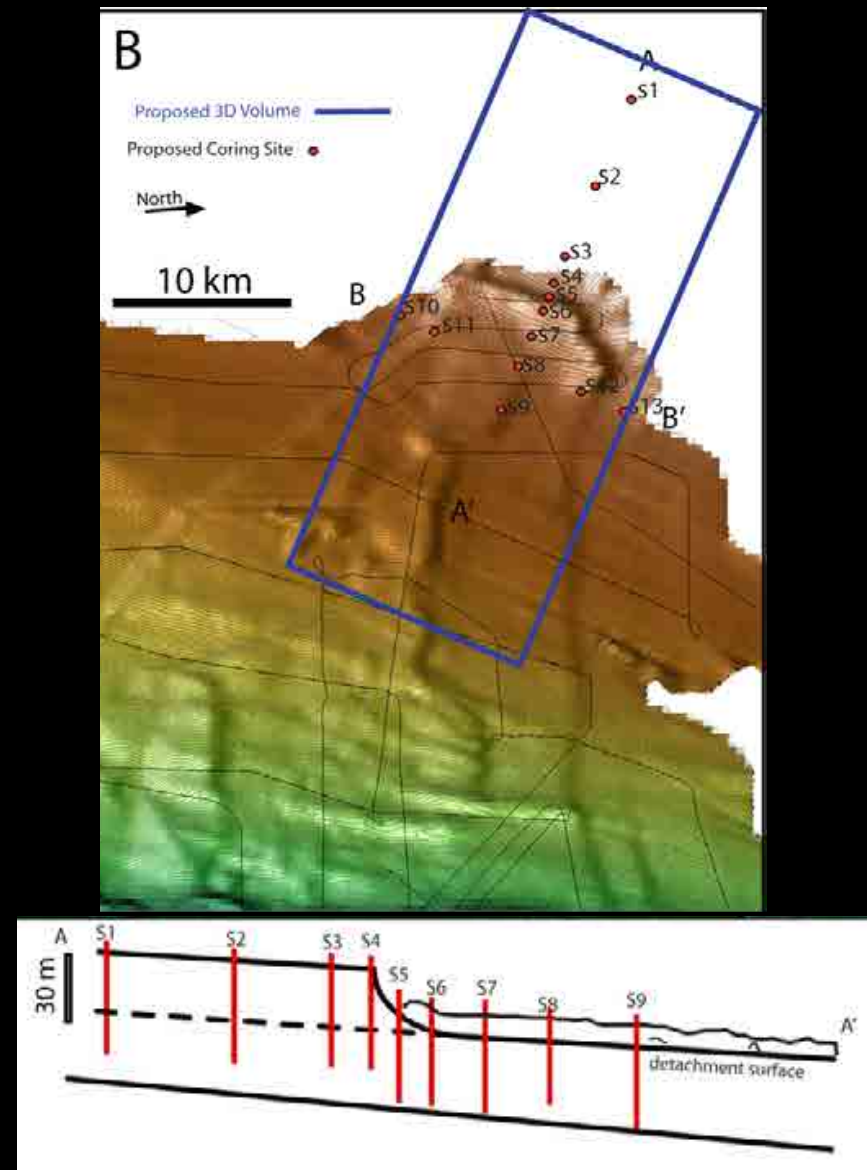
- 1) Widespread, large-scale margin evolution
- 2) Rapid, m-km but affect overall margin shape
- 3) Geohazard potential (process and society)
- 4) Integrates surface processes and feedbacks
- 5) Multi-disciplinary
 - geology, geophysics, geotechnics, fluids...

Key Data/Observations – Slide and Adjacent

- 1) Morphology and distribution
- 2) In situ pressure, temperature, and strength
- 3) Tilt and strain data
- 4) Multiple (4d) images or measurements



Sawyer et al., 2009



Hornbach et al. White Paper

GeoPRISMS-Addressable Questions

What dictates size, runout, and recurrence?
- margin shape and evolution

What drives failure along different margins?

What controls hazard potential?

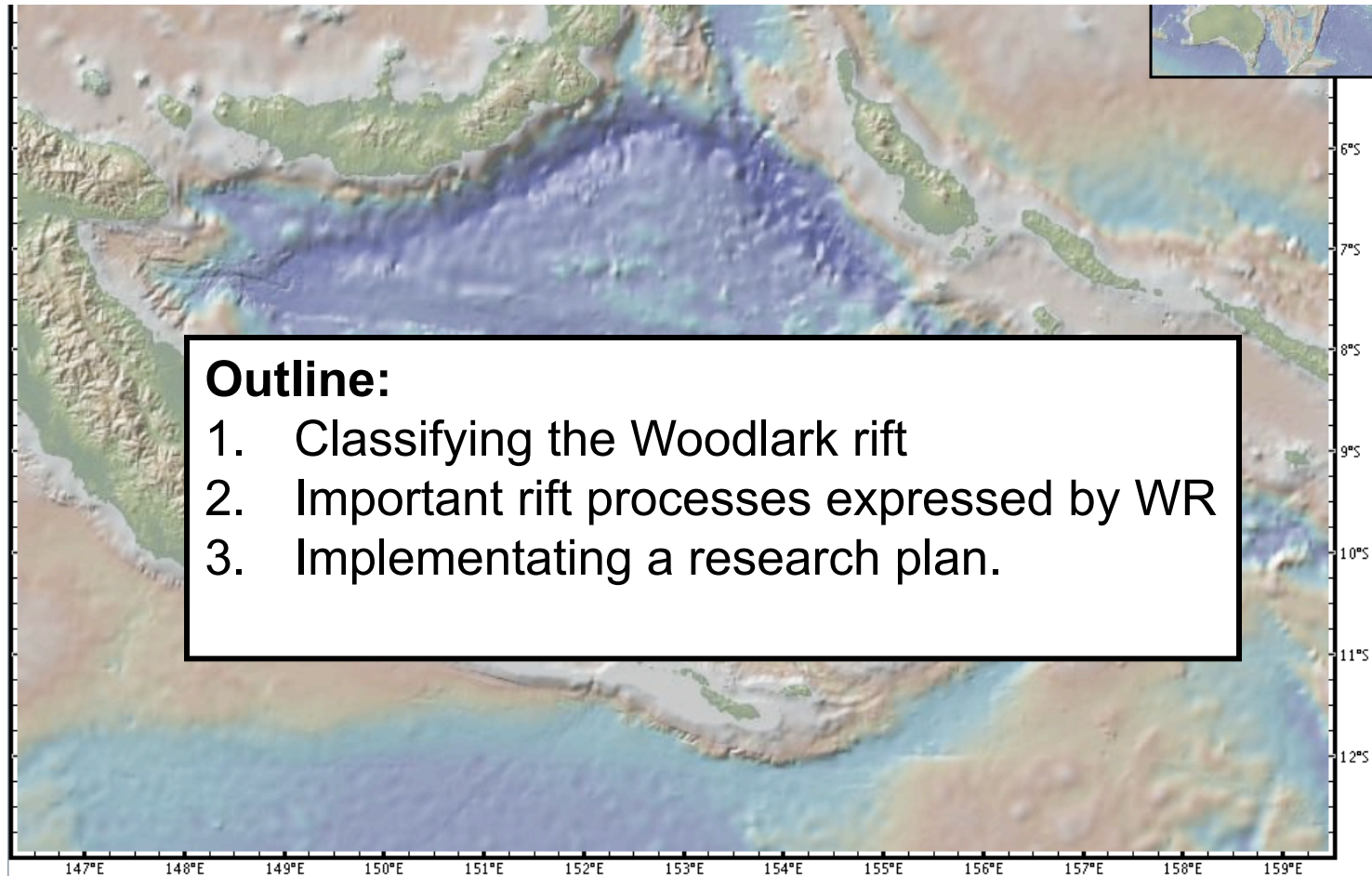
Is the failure evolving or stopping? Why?

How are different failures recorded in the strata?

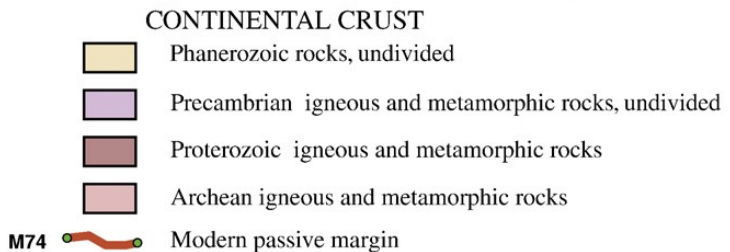
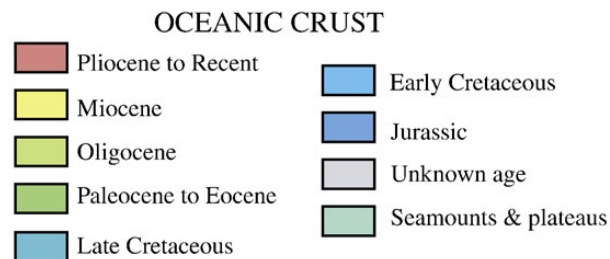
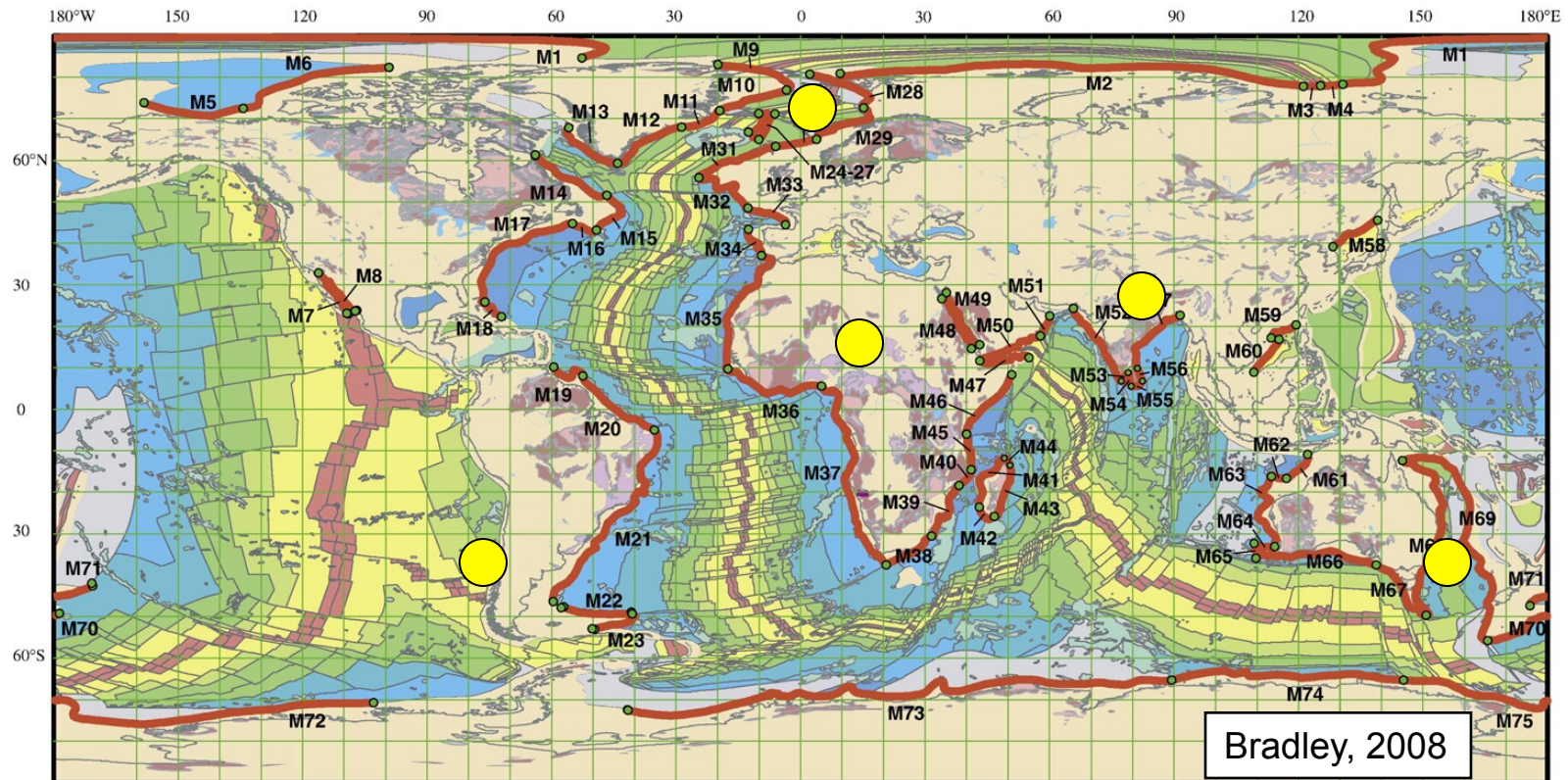
Woodlark Rifting White Paper

Important processes and implementation

Paul Mann, Suzanne Baldwin, Paul Fitzgerald, Geoff Abers, Jim Gaherty,
Laura Wallace, Guy Fitz, Nathan Daczko and others



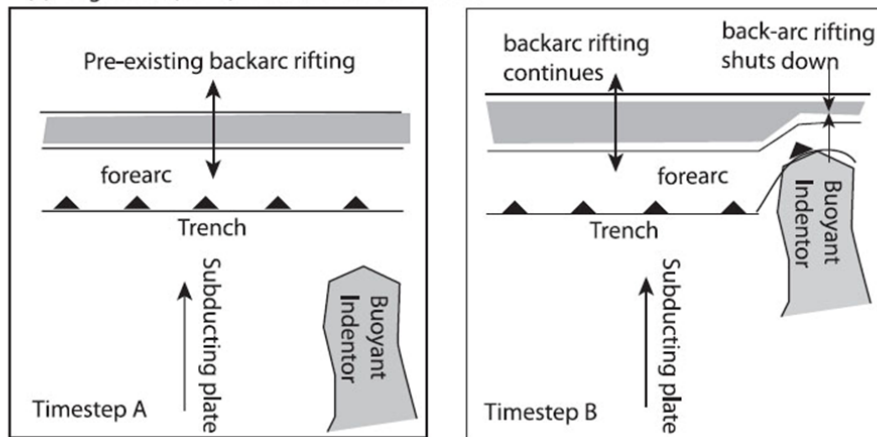
“Dead rifts” front major ocean basins and are deeply buried by 75 modern passive margins (brown lines); intracratonic rifts are either failed rifts (aulacogens) or very slowly evolving (>10 my) and can lack clear driving tectonic forces



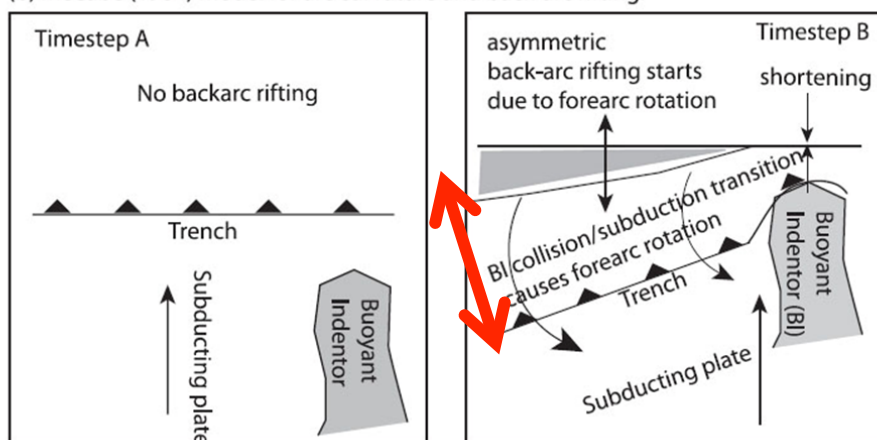
Rifts forming today on active plate margins at fast rates form by 3 processes: ***indent-related, rollback-related, or strike-slip-related***. All three types are more accessible than “dead rifts” beneath passive margins and evolve at faster plate boundary rates than intracratonic rifts

Laura Wallace papers:
crustal-driven indent process

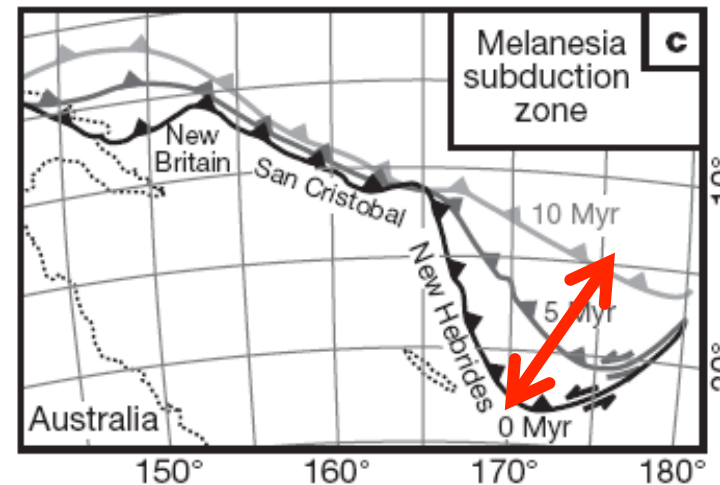
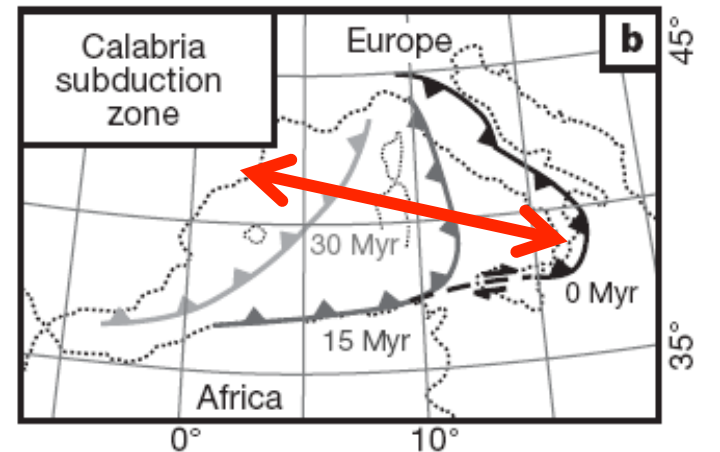
(a) Vogt et al. (1976) model for arc curvature:

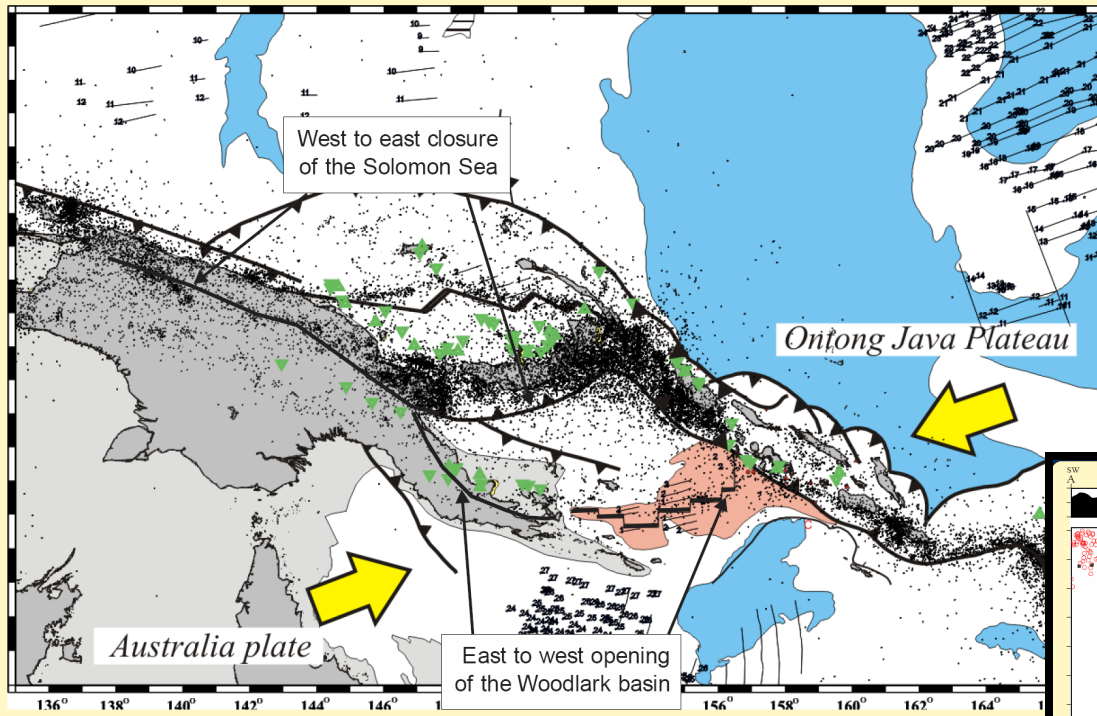


(b) McCabe (1984) model for arc curvature and back-arc rifting:

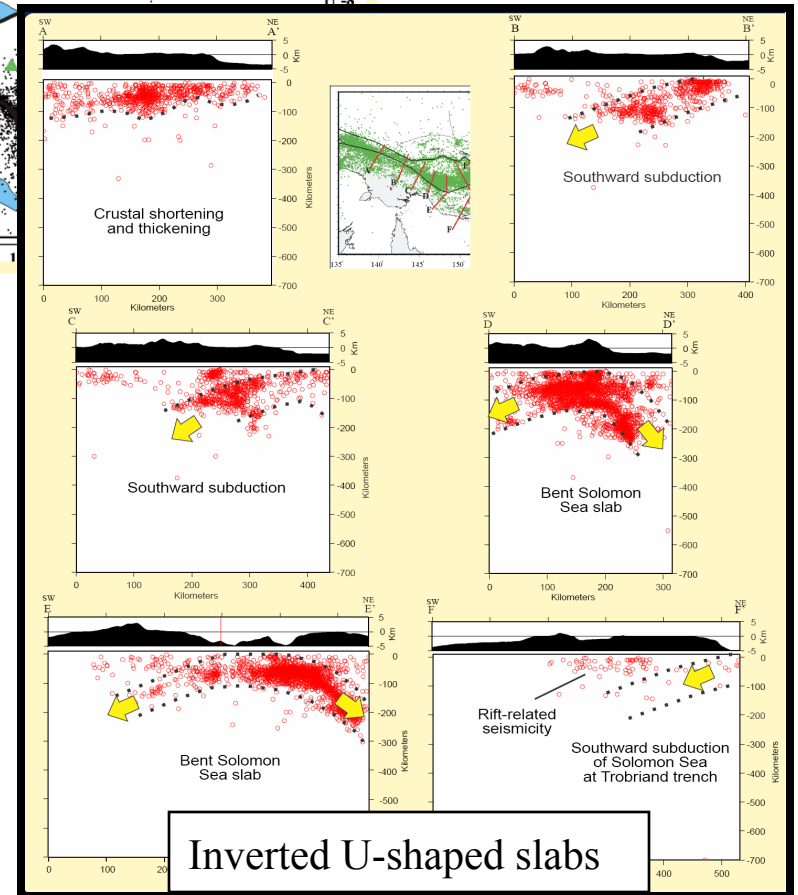


W.P. Schellart papers:
mantle-based process

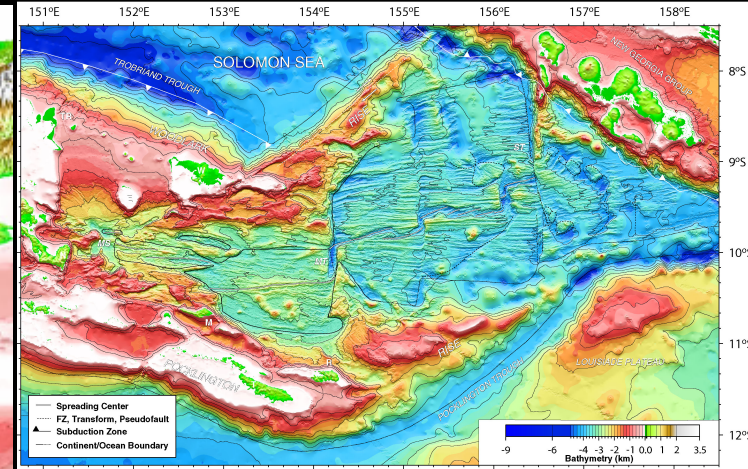
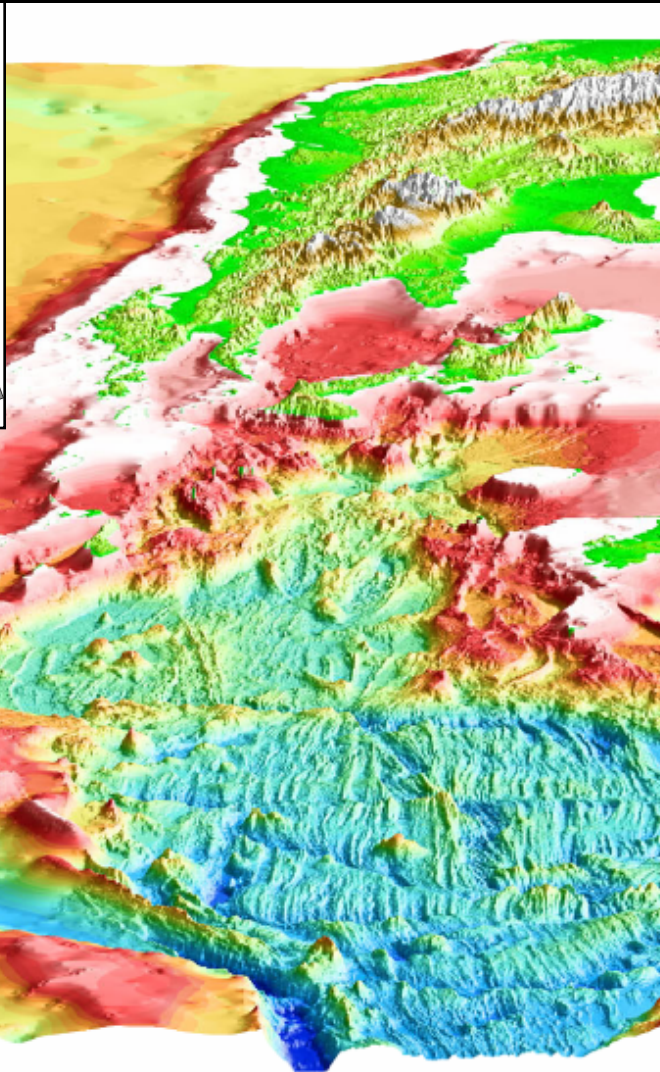
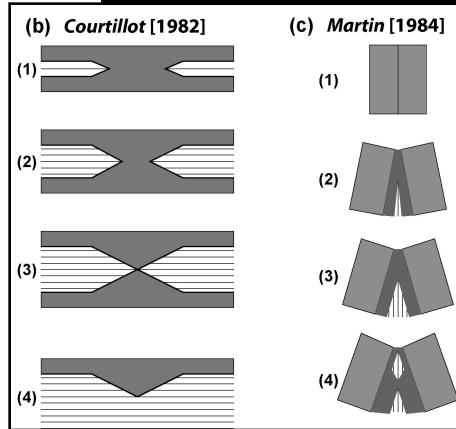




Rifting in Papua New Guinea area is either indent or rollback-related and very rapidly evolving (Australia began colliding with PNG about 30 Ma – middle Oligocene; OJP convergence and Woodlark basin started ~6 Ma, rate of westward propagation of rift tip is 140 km/my.

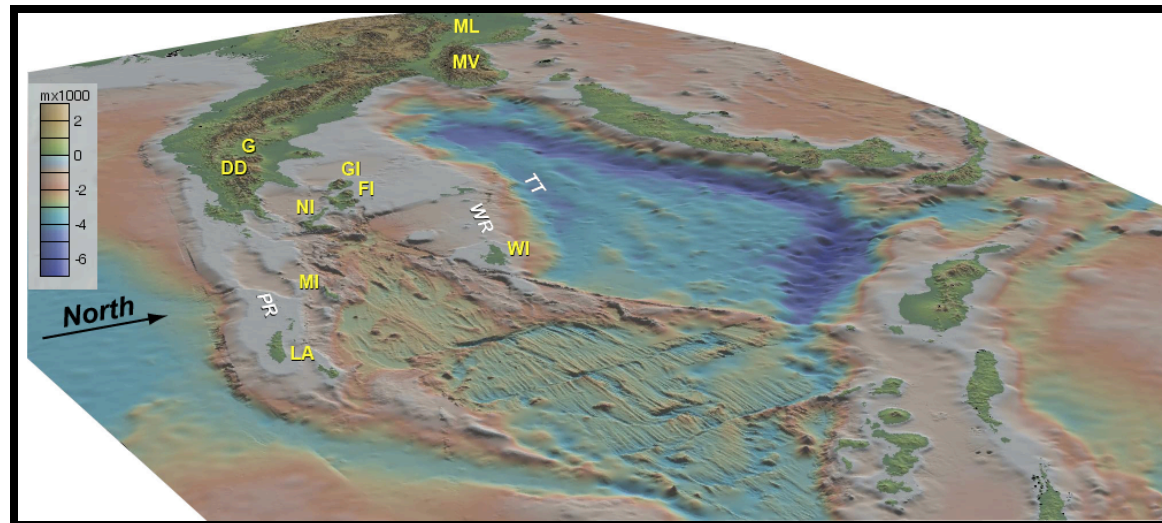
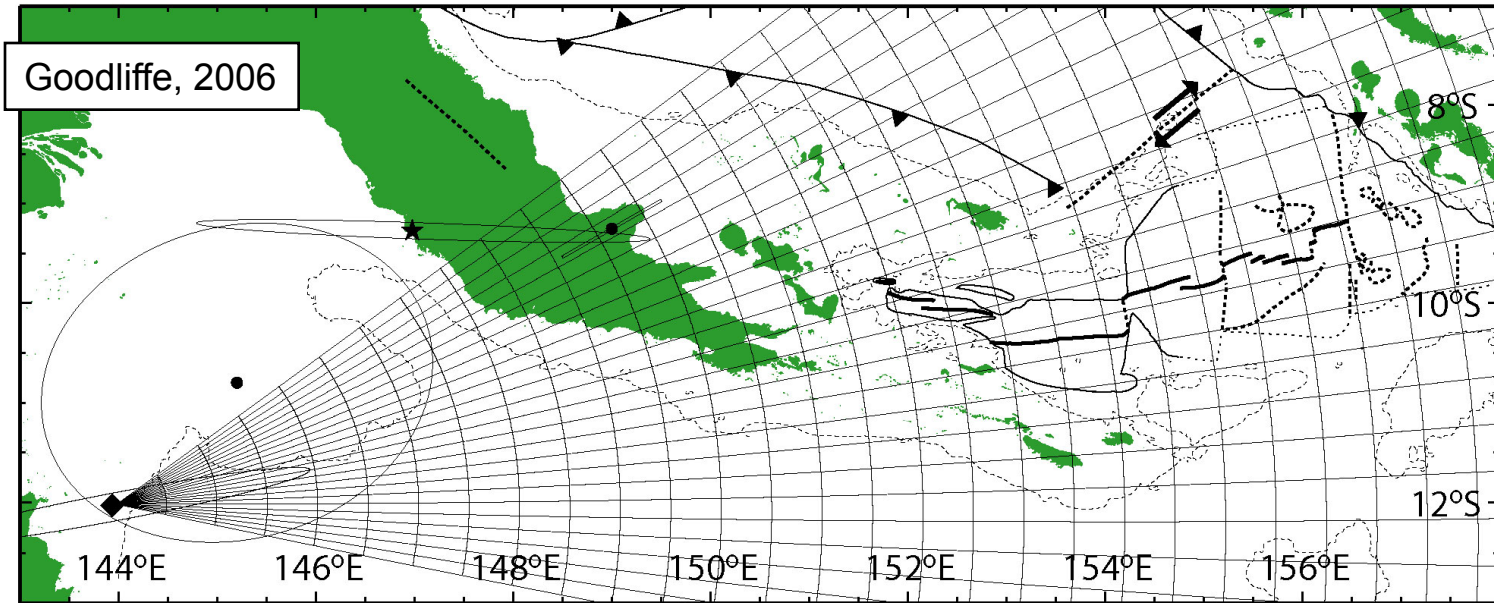


Active oceanic spreading centers propagating into continents are rare and present valuable analogs to the larger ocean basins because of their, diachronous time-space progression: Woodlark, Lena, Afar

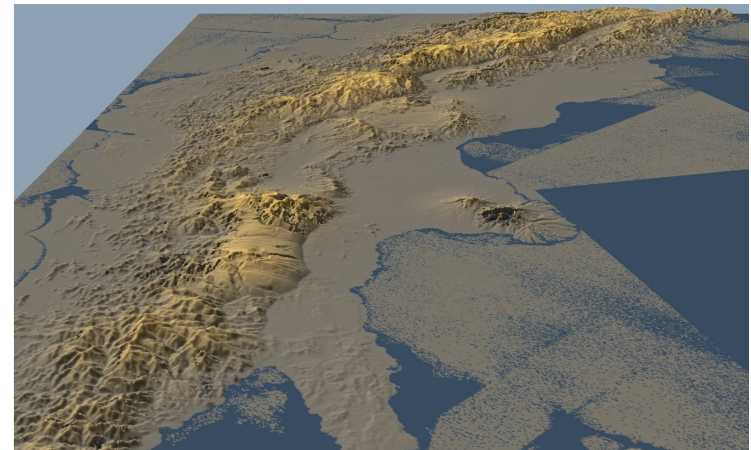
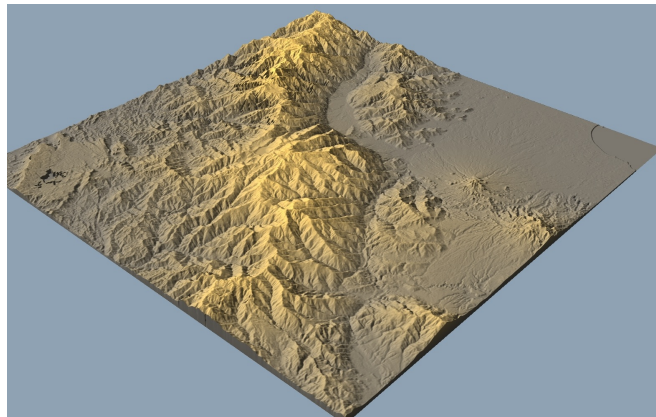
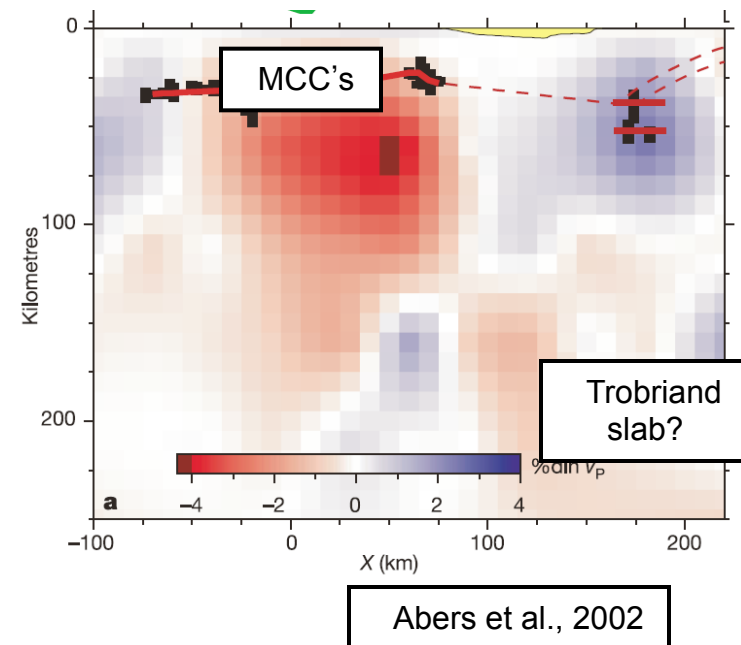
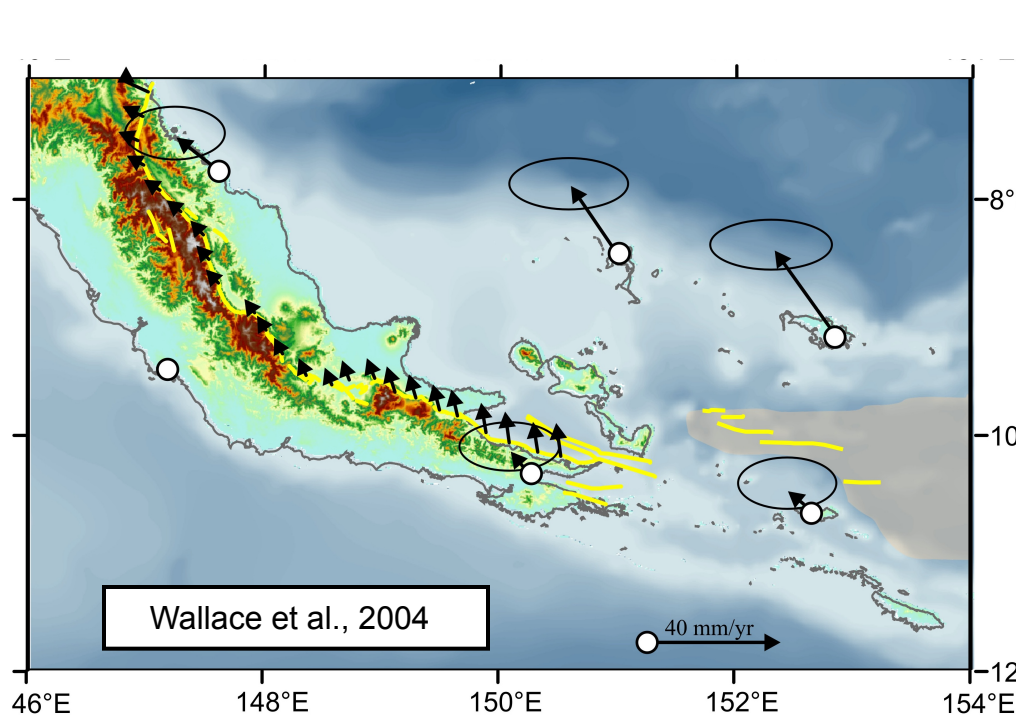


Andrew Goodliffe et al., 2005

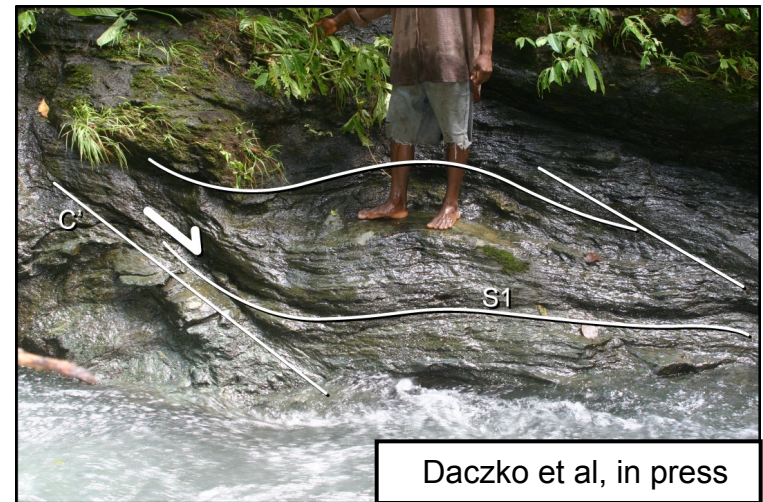
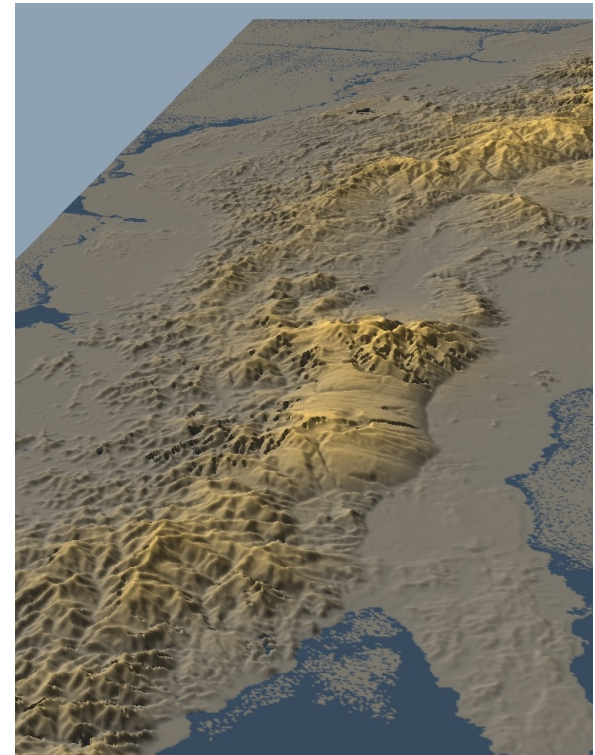
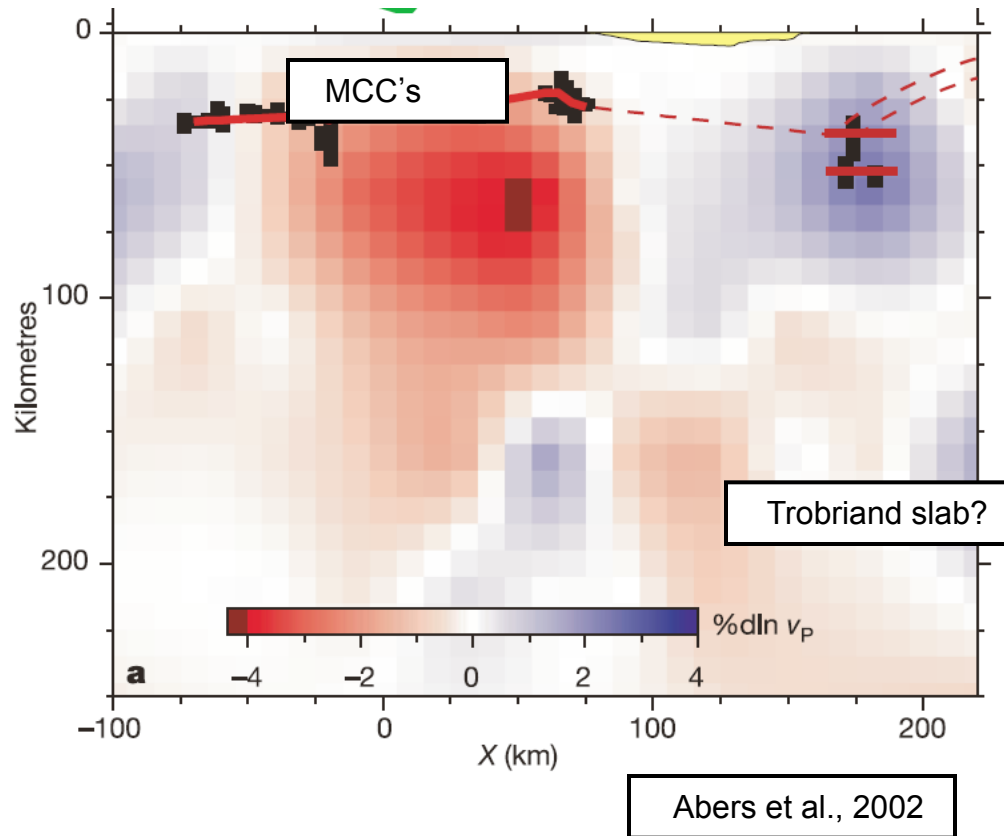
Discrepancy of oceanic spreading vs. continental extension about the pole of opening: where is it taken up? MCC's? Lower crust? Other unknown processes?



“Forcing parameters” of active plate boundary rifts can be inferred from GPS measurements (crustal) or tomography (mantle). Studies that show links between deformation and magmatism require combining data types from all depths.

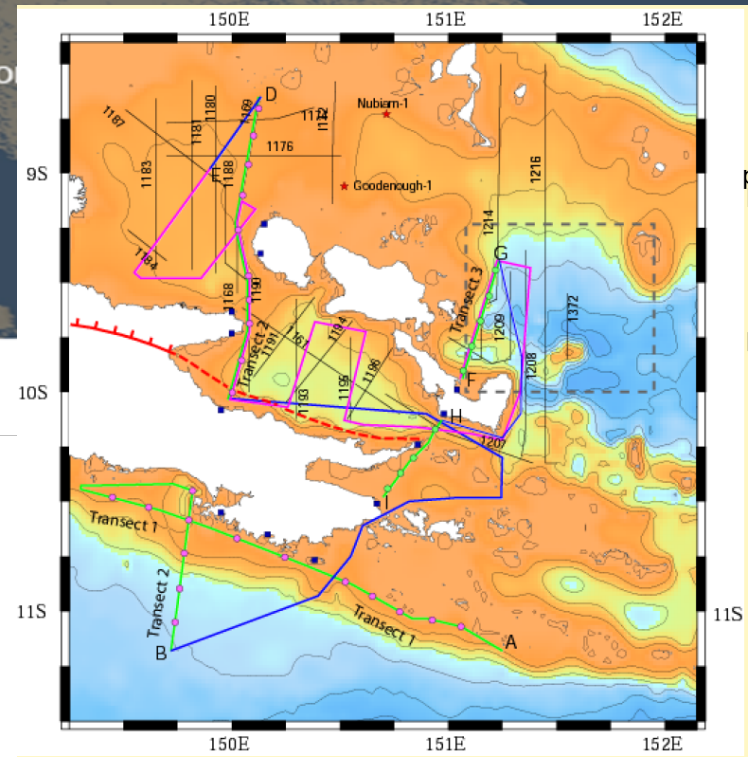
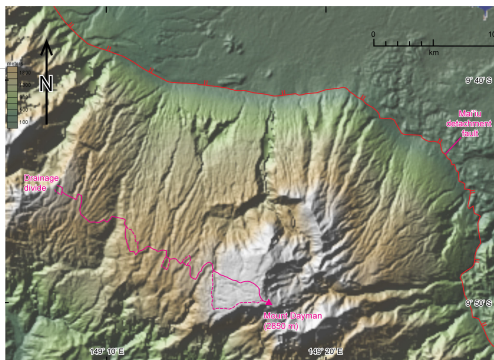
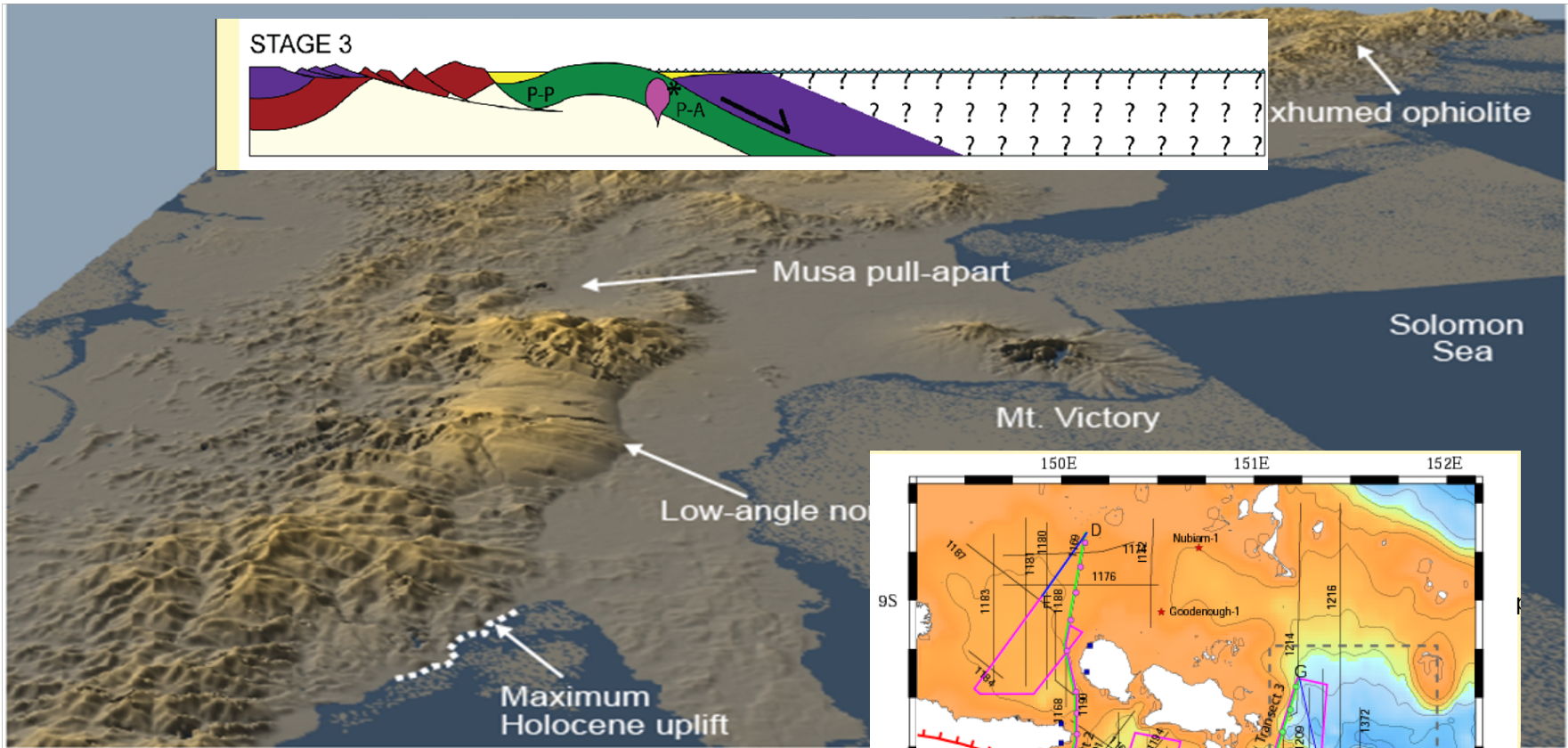


*"Few examples of low-angle detachments associated with continental breakup have been described".
Luc Lavier quoting ??.*



What controls locations of rifts?

Sites of former arcs or collisional zones? Inherited thrust surfaces?



Fast rates from basement rocks: Ferguson MCC uplifting at average rate of 12-17 mm/yr

➤ (U)HP rocks have been exhumed from beneath km-scale shear zones at plate tectonic rates (> 1 cm/yr)

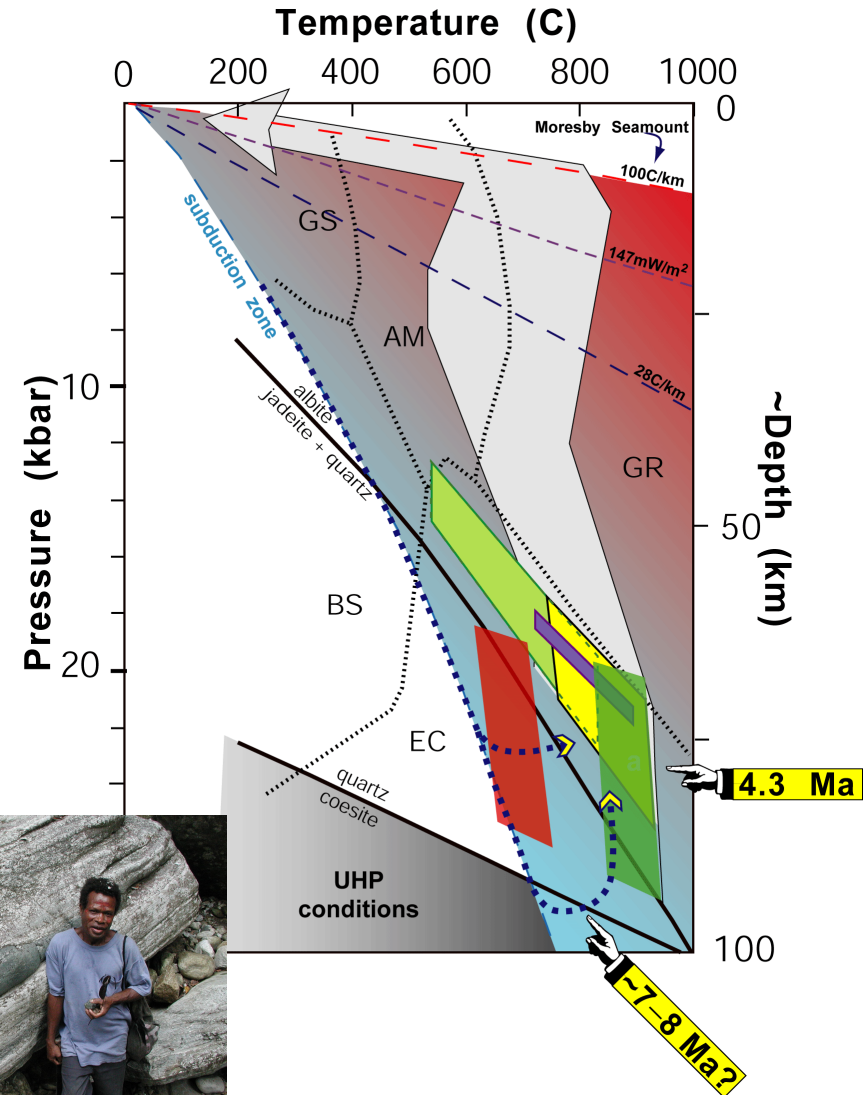
➤ Average minimum vertical exhumation rates (~ 12 - 17 mm/yr)

➤ 25-51 mm/yr assuming exhumation from beneath 20-30° NNE dipping shear zones

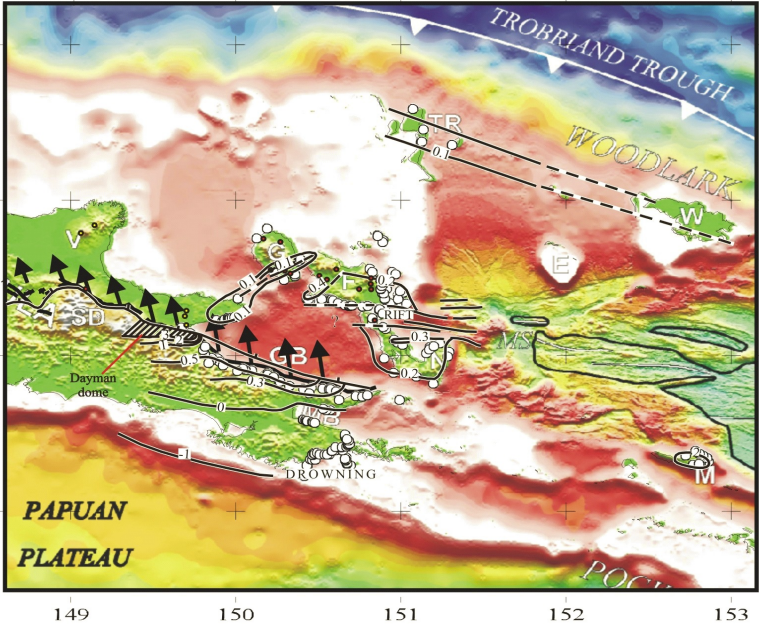
➤ Compare with:

➤ 25-40 mm/yr extension rates near rift tip (Abers, 2001)

➤ 12-40 mm/yr half-spreading rates from magnetic anomalies (Benes et al., 1994) and GPS crustal motion studies (Tregoning et al., 1998)

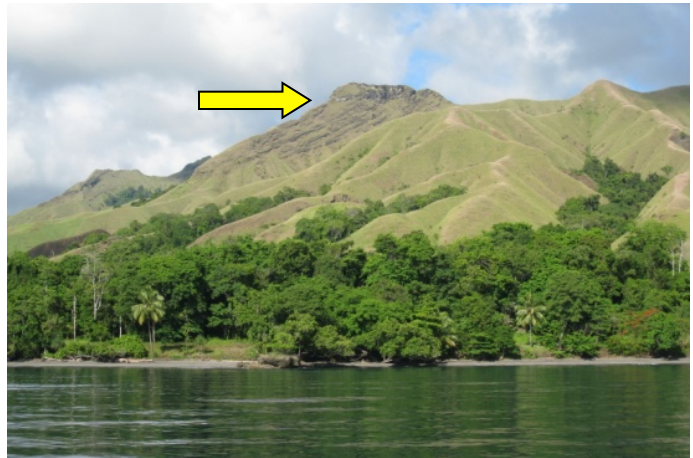
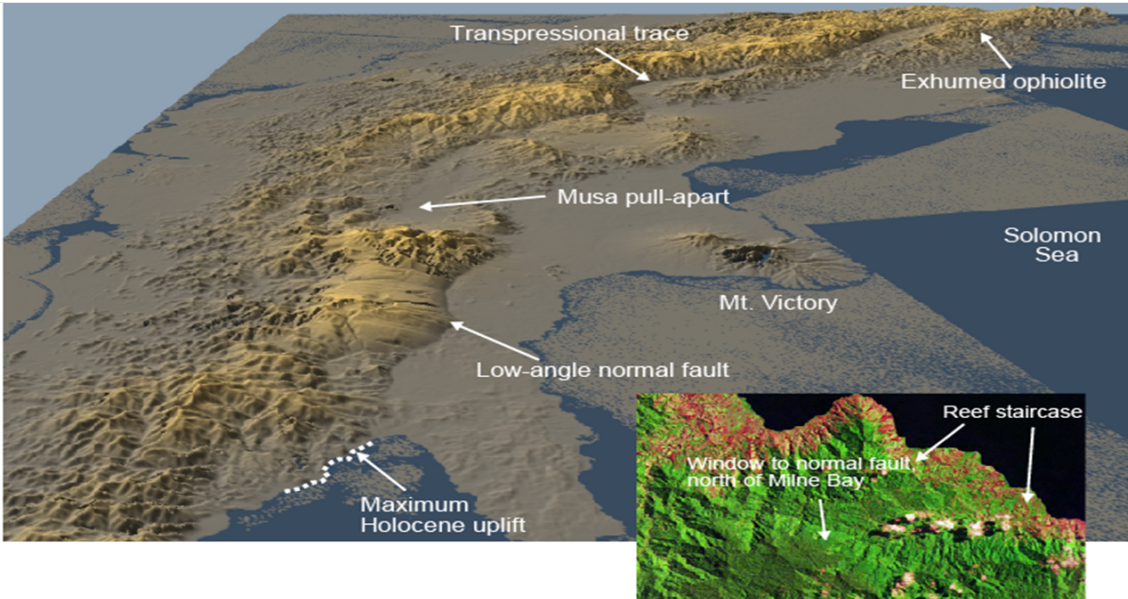


Fast rates from coral reefs: OSFZ footwall block uplifting at 4 mm/yr

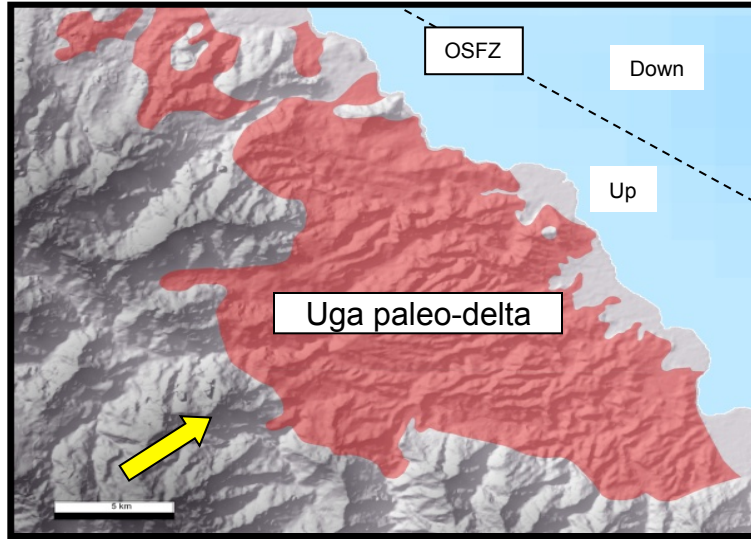


322 m high surface was at sea level ~127 ka

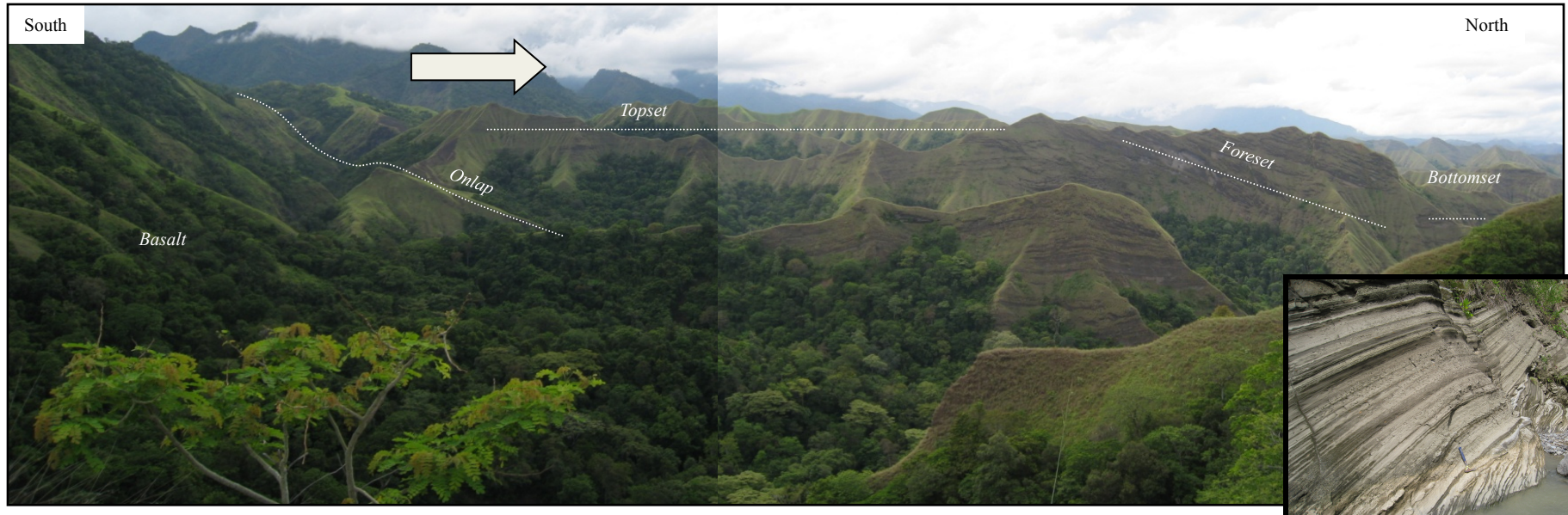
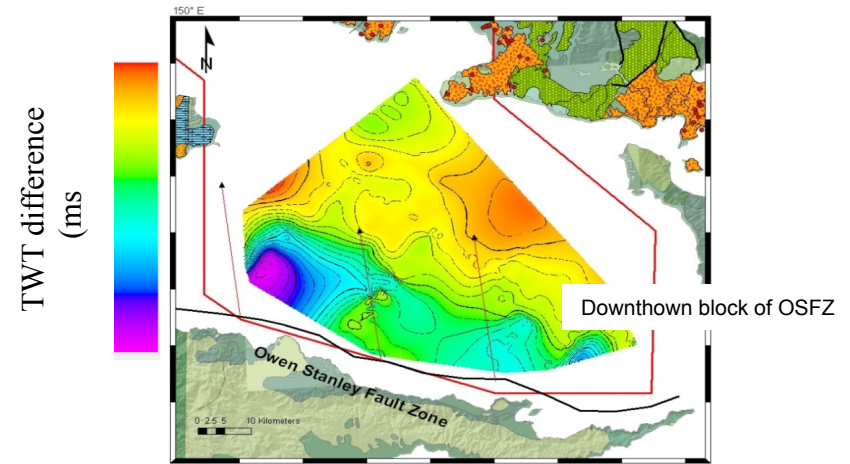
127 Ma coral unconformable on
Folded, Pleistocene deltaic sediments



Sediment flux into the rift: Pliocene-Pleistocene Uga delta uplifted on footwall of Owen-Stanley normal fault; thick seen on seismic data on hanging wall block

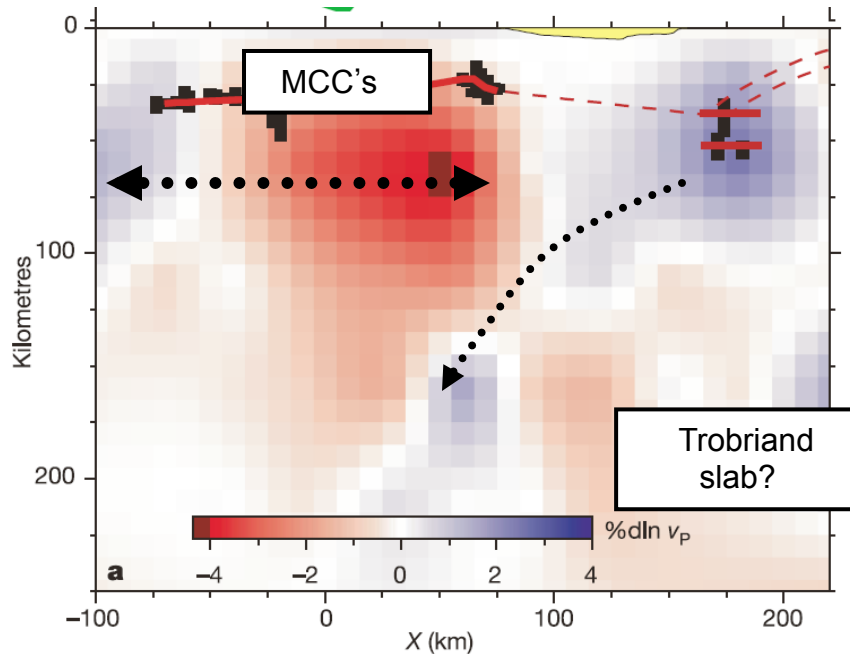


Pliocene to Plio-Pleistocene Isochron map of Goodenough basin from Fitz, in prep.



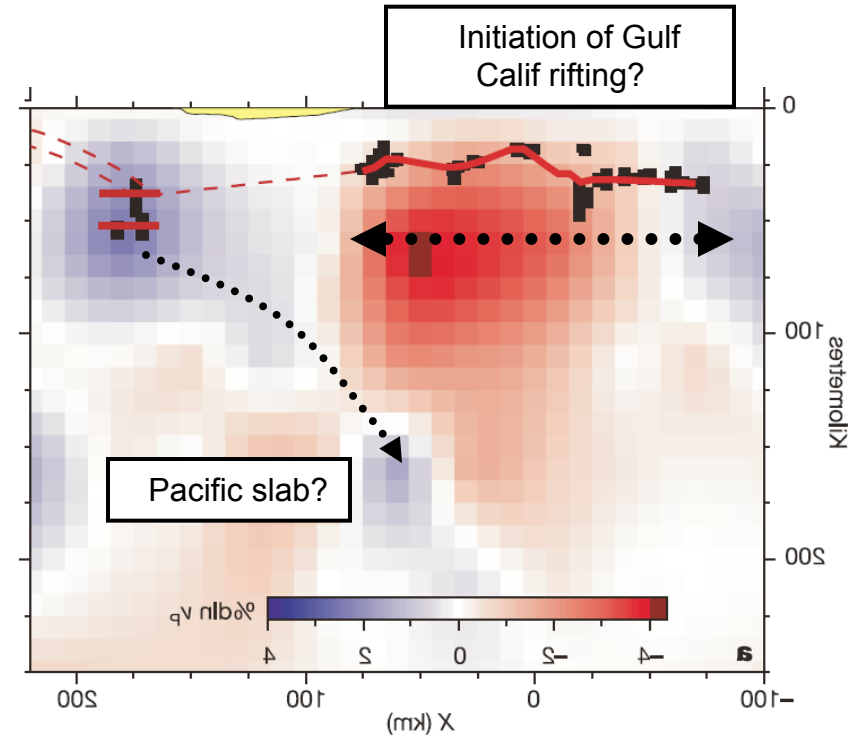
How is lower crust, mantle and deeply subducted slabs return to the surface during the rifting process? Is orogenic collapse triggered by rollback the answer?

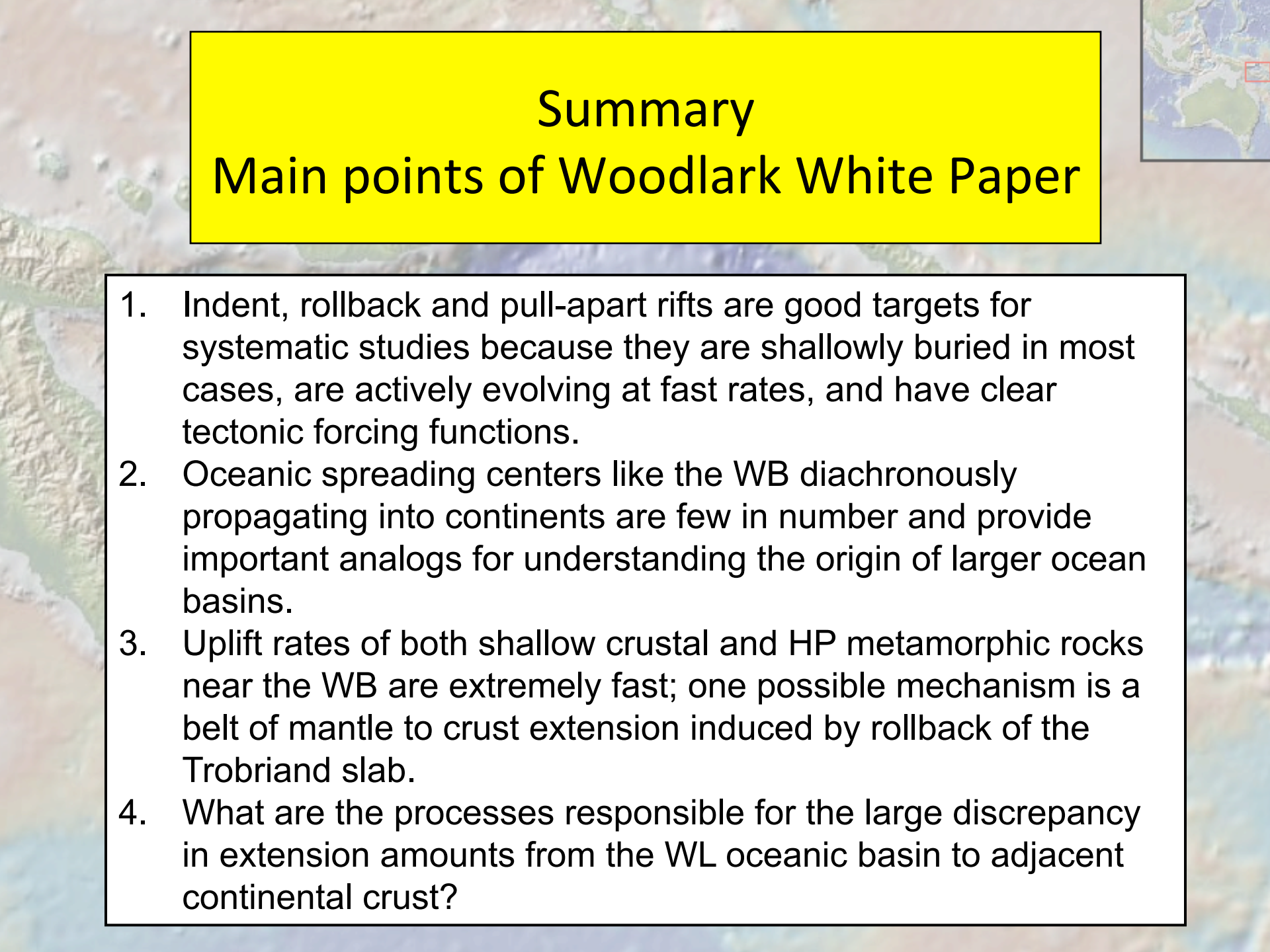
Papua New Guinea



Abers et al., 2002

Baja-Gulf of California analog





Summary

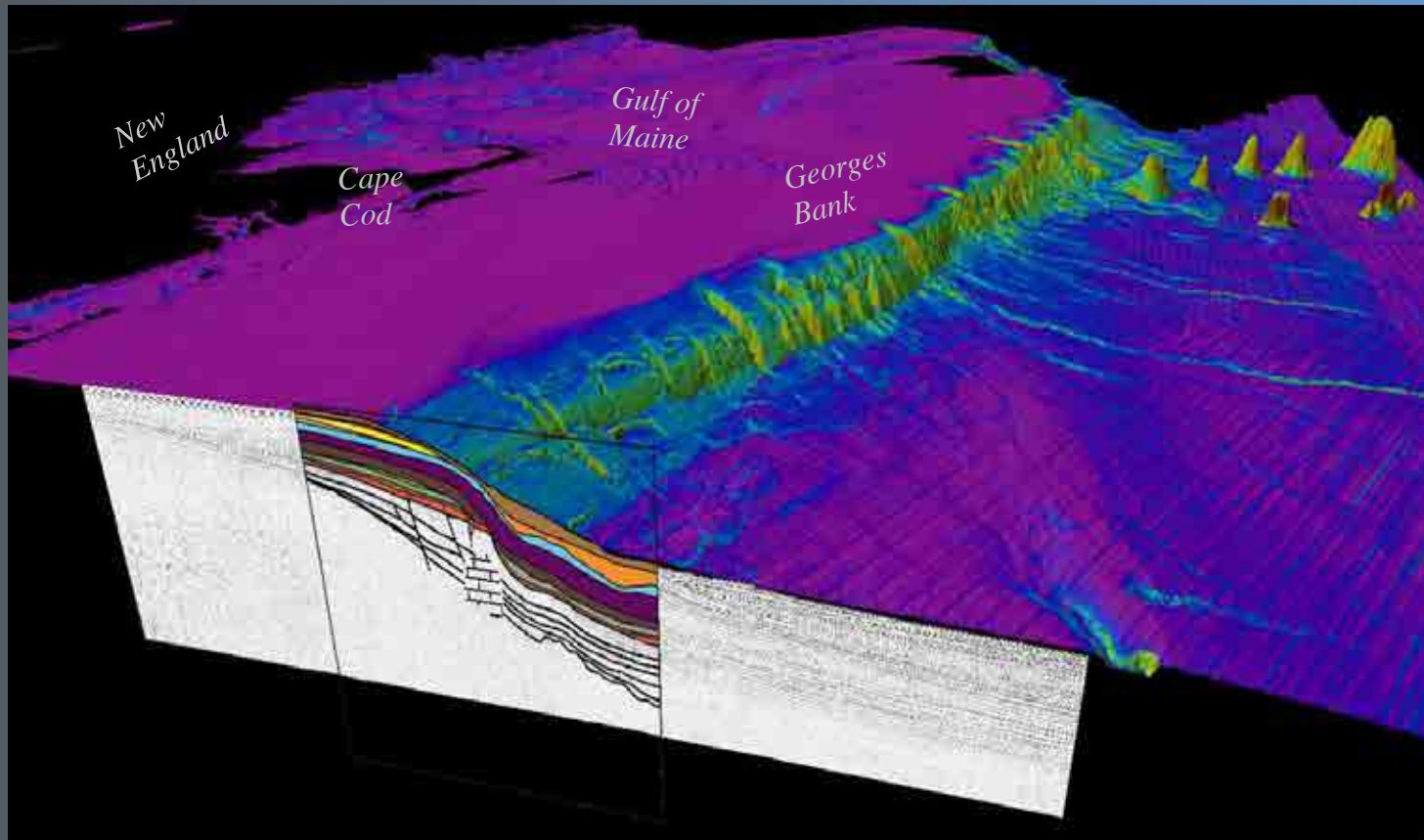
Main points of Woodlark White Paper

1. Indent, rollback and pull-apart rifts are good targets for systematic studies because they are shallowly buried in most cases, are actively evolving at fast rates, and have clear tectonic forcing functions.
2. Oceanic spreading centers like the WB diachronously propagating into continents are few in number and provide important analogs for understanding the origin of larger ocean basins.
3. Uplift rates of both shallow crustal and HP metamorphic rocks near the WB are extremely fast; one possible mechanism is a belt of mantle to crust extension induced by rollback of the Trobriand slab.
4. What are the processes responsible for the large discrepancy in extension amounts from the WL oceanic basin to adjacent continental crust?

US Atlantic Continental Margin

White Papers:

#7 - Gaherty et al. #9 – Hornbach et al. #12 – Olsen et al.



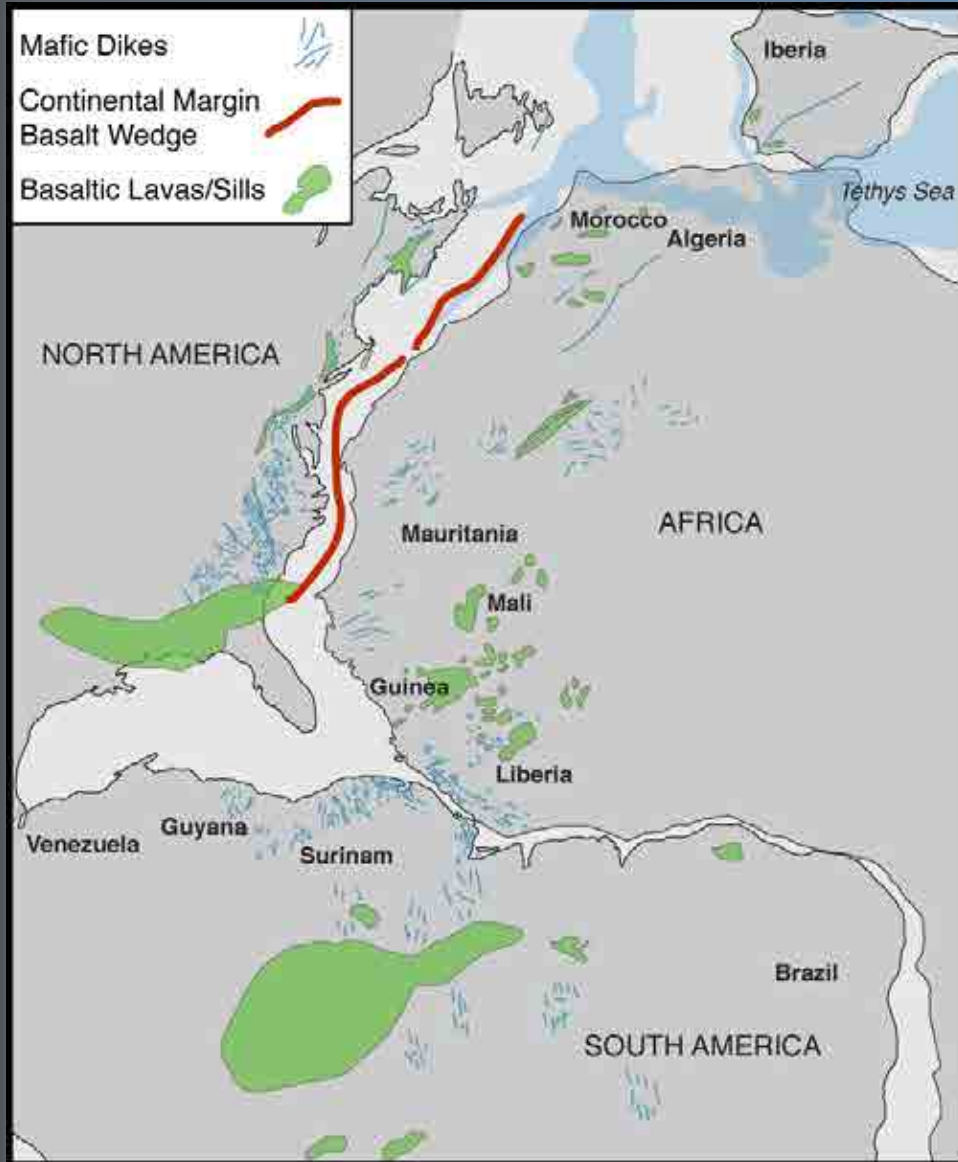
Outline

1. RIE Research Opportunities
(rift --> breakup --> post-rift)
2. Groundwork and Logistics
3. Synergistic Opportunities

Rift Initiation and Evolution: Compelling Science at a Drifting Margin

1. Distribution of magmatism and deformation in the mantle lithosphere
2. 3D rift structure and segmentation
3. Nature of transitional crust
4. Relationship between failed rift basins and successfully rifted margin
5. Post-rift evolution and growth of a passive margin

Central Atlantic Magmatic Province (CAMP)



McHone (Tectonophysics, 2000)

Rifting began ~230 Ma

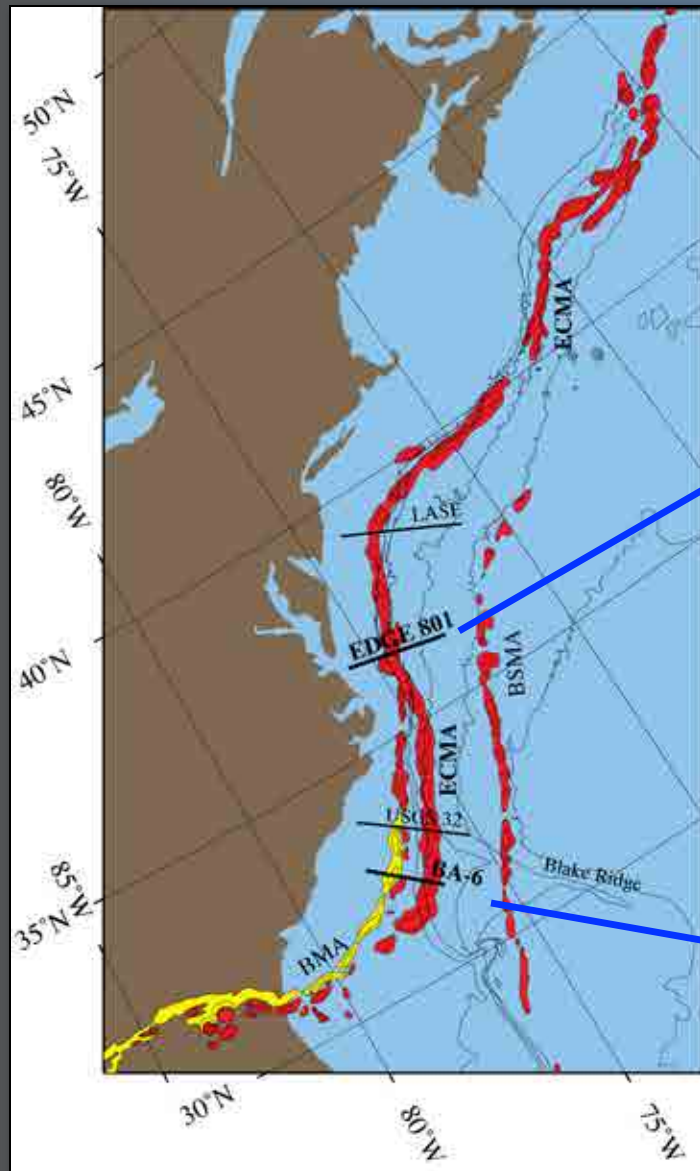
~200 Ma LIP emplaced -->
Approximately coincident w/
breakup and mass extinction
(end-Triassic)

Relationship between
deformation, magmatism and
syn-rift sedimentation onshore
and offshore is poorly
understood (failed versus
successful rifting)

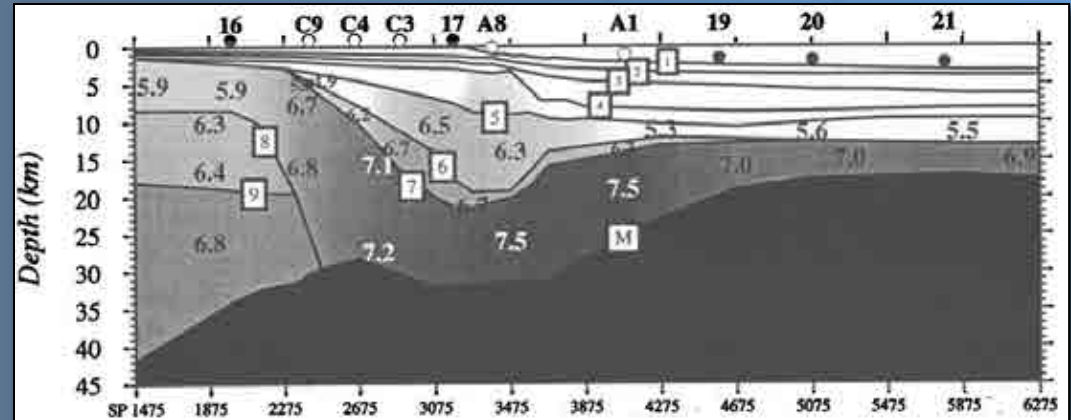
Unknown geodynamic source for
magma

See white papers by Gaherty et al. and Olsen et al.

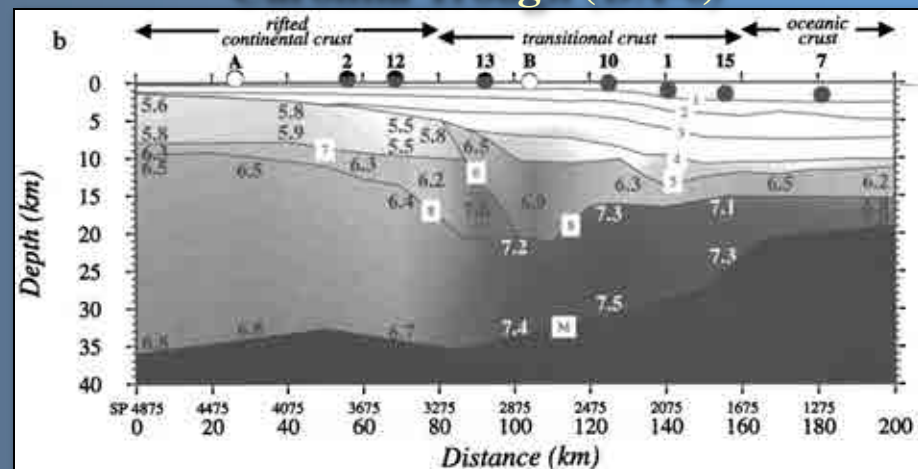
High lower crustal Vp – different melting regime than MOR



Mid-Atlantic (EDGE 801)



Carolina Trough (BA-6)



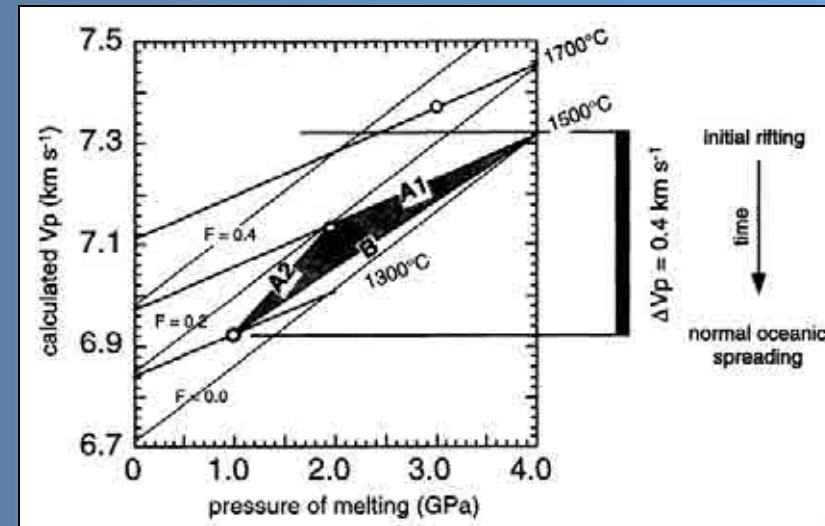
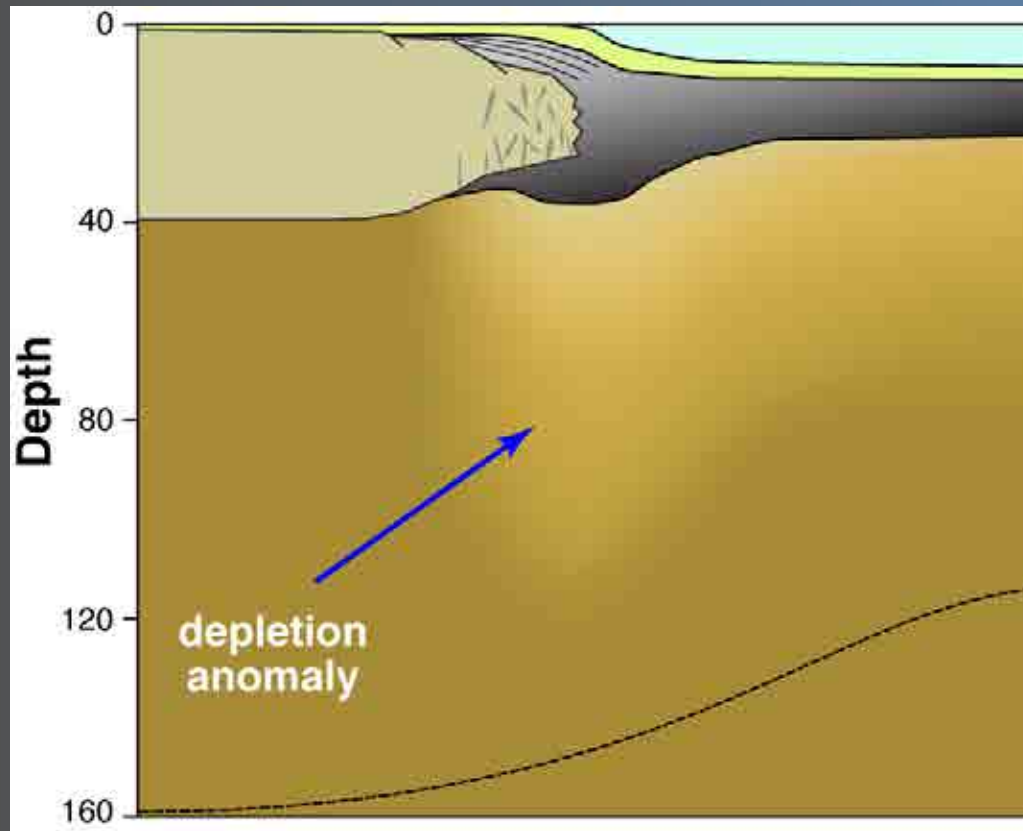
Holbrook et al. (JGR 1994a,b)

Need more constraints on the nature of transitional crust

High V_p > High pressure of melting > Depletion anomaly

New seismic imaging:

- Along-margin architecture
- Mantle source of melt

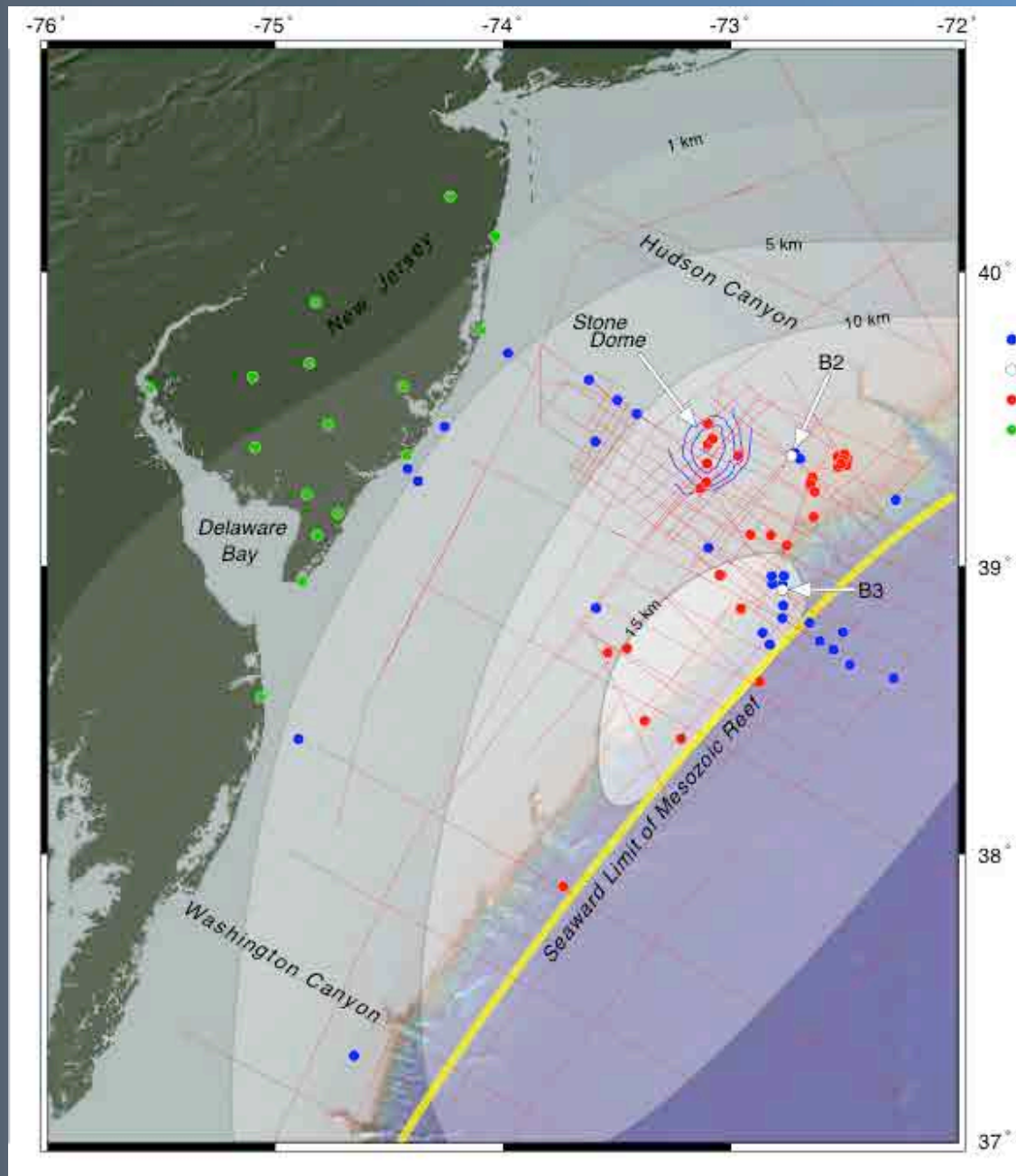


High velocities imply high average pressure of melting

This should leave behind a low density residue of melting that may stabilize the margin.

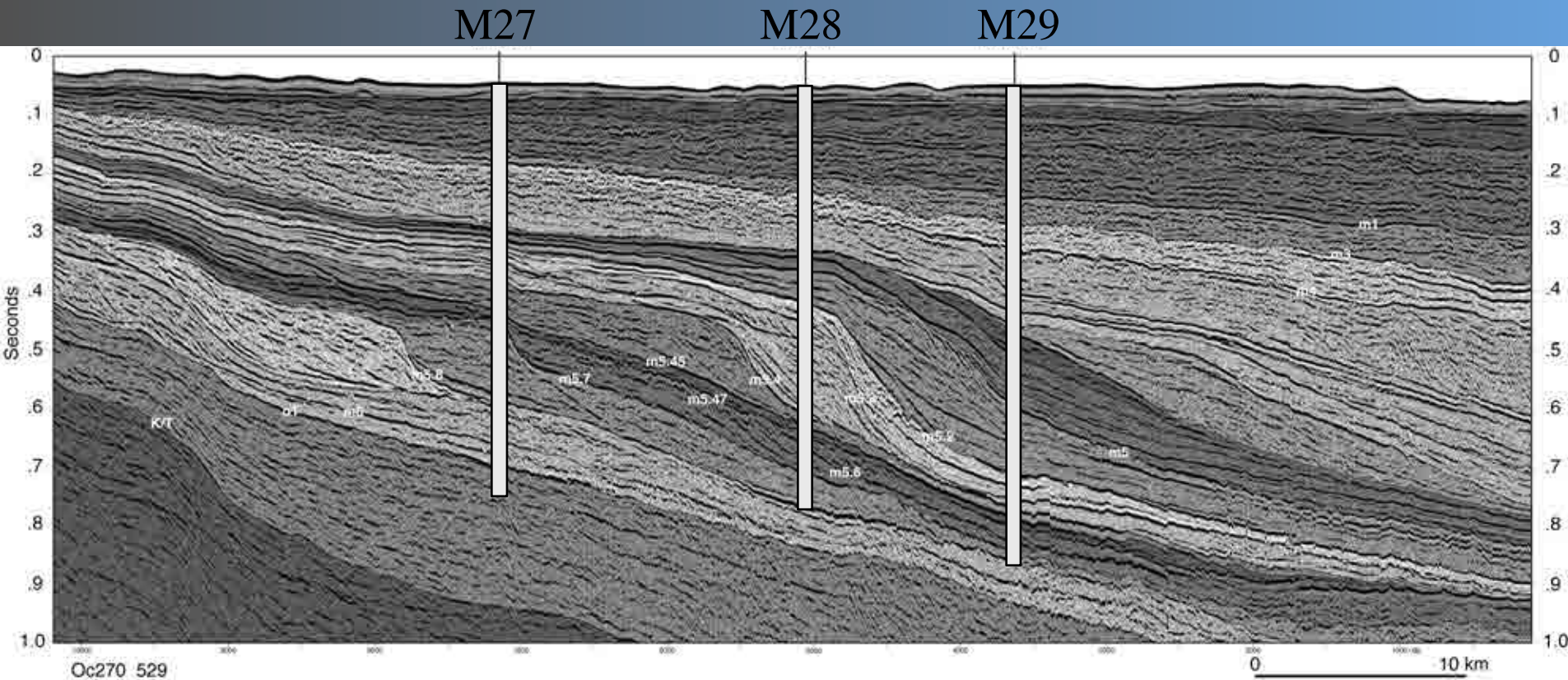
Kelemen and Holbrook (JGR 1994)

Post-rift Evolution



(Figure from Greg Mountain)

How has rift architecture and lithospheric evolution influenced post-rift sediment dispersal and accumulation?



(Figure from Greg Mountain)

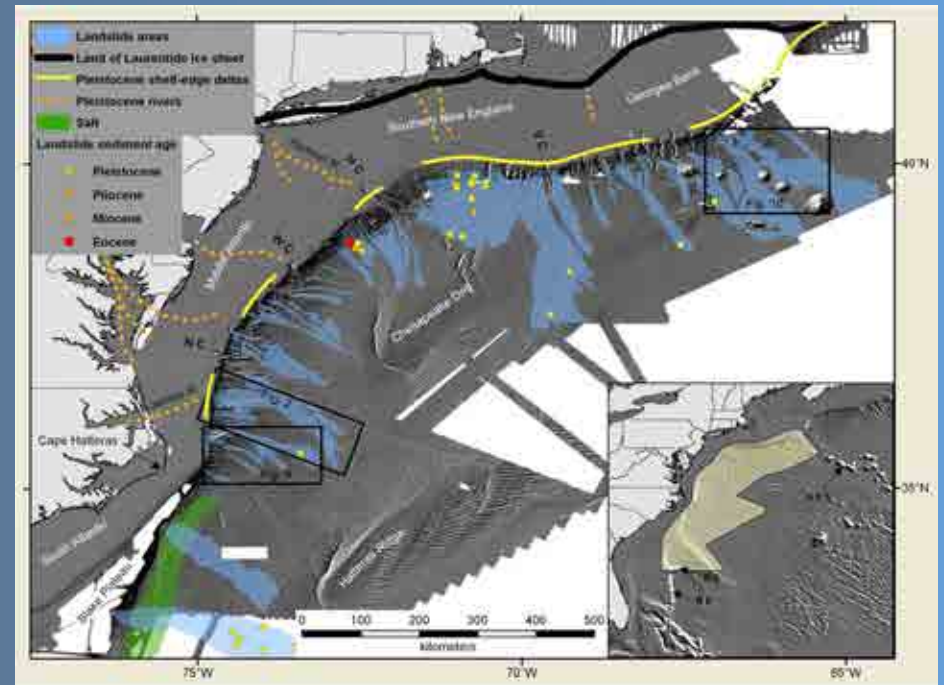
3D complexity requires 3D imagery

Surficial Processes:

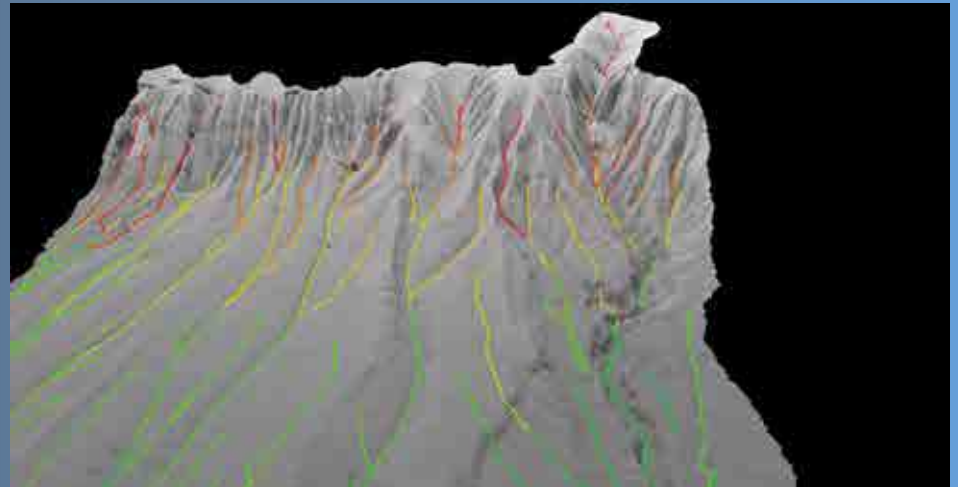
How is today's morphology influenced by early margin development ?

Quantitative geomorphology --> Sediment transport modeling, submarine Canyon formation, turbidity flows

Geohazards --> submarine landslides, slope stability, tsunamis, seismicity



Twichell et al., (2009)



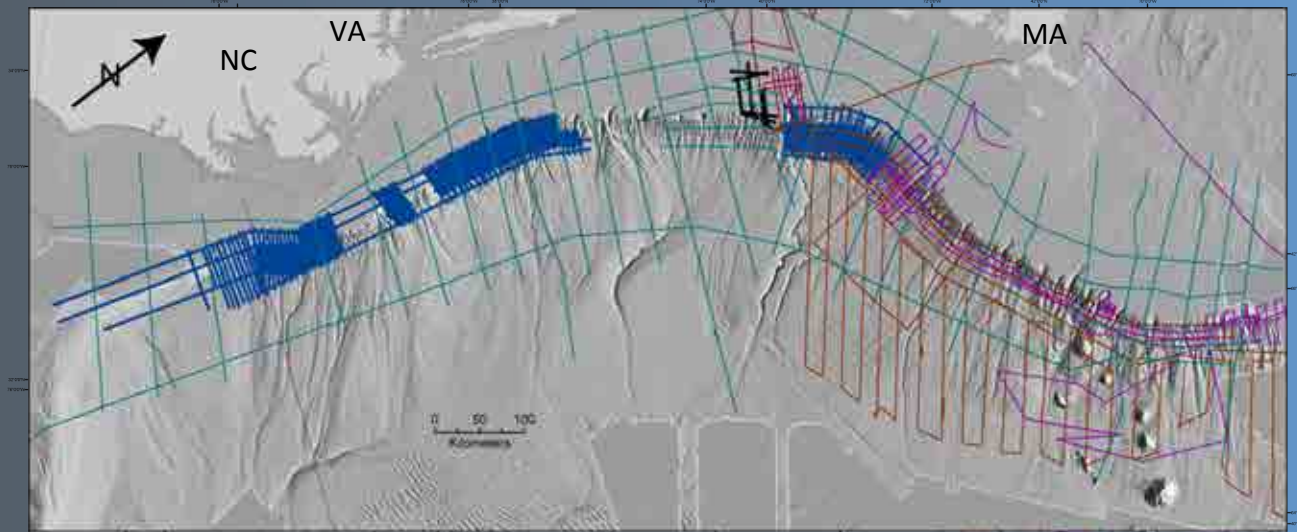
(See white paper by Hornbach et al.)

Groundwork and Logistical Considerations

Very heavily studied margin

- Dozens of hydrocarbon wells, 1-dozen IODP wells, AMCOR wells, many others
- 10's of thousands of line-km of reconnaissance-grade seismic reflection data
- Onshore seismic, well data, mapping
- Opportunity to synthesize onshore and offshore data
- Comprehensive potential field data

Need onshore/offshore synthesis



Synergistic Opportunities

Law of the Sea- ECS Project (USGS, NOAA)

- MCS spaced every 60 nautical miles (beyond 200 nautical mi)
- Coincident refraction (select lines)

US Array

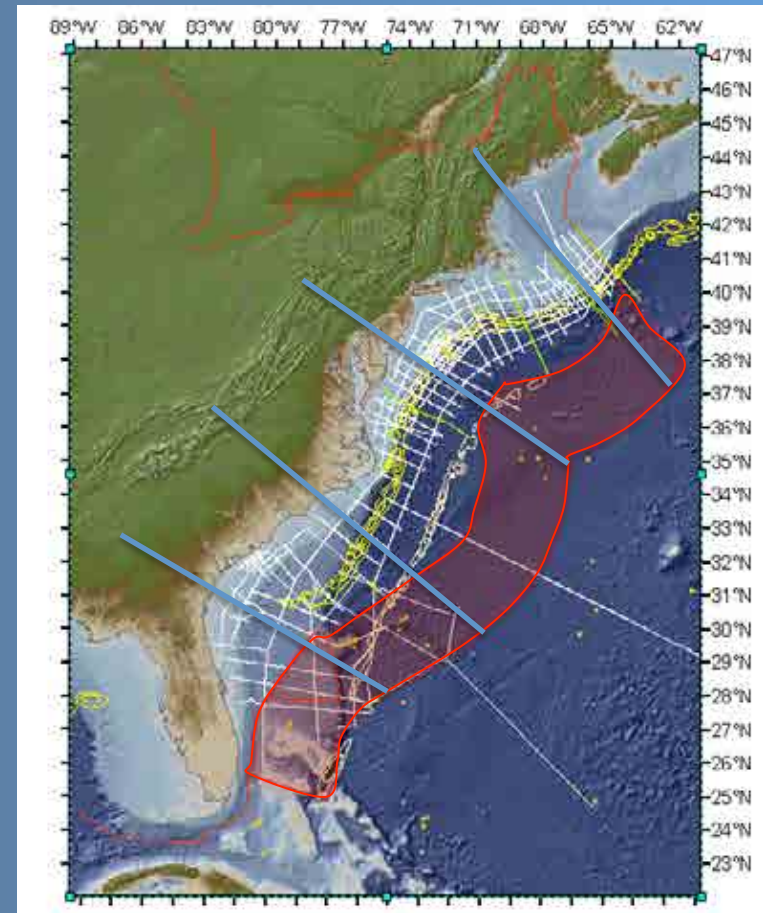
Eastern US transportable array 2013

Oil and Gas Exploration

Moratorium may be lifted...

Canadian-Atlantic and Conjugate Margins

International collaboration



Numerous Opportunities for Coordinates Studies

Summary

1. Late-stage rifting, transition to drifting
 - Magmatism (in time and space) and upper mantle deformation
 - Segmentation and along-axis variation
 - Rift localization and formation of transitional crust
 - Breakup unconformity
2. Post-rift evolution
 - Sediment dispersion patterns, passive margin evolution, climate control
 - Geomorphic evolution (progradation, clinoform development, submarine canyon formation)
 - Geohazards (submarine landslides, tsunami generation, seismicity)
3. Synergistic Opportunities
 - US Array
 - Extended Continental Shelf Project (Law of the Sea)
 - Industry and international collaboration
 - International collaboration

Continental Breakup and Formation of Rifted Margins: The Gulf of Mexico as a Natural Laboratory



D. Harry, *Colorado State University*

R. Stern, *University Of Texas At Dallas*

E. Anthony, *University Of Texas At El Paso*

G.R. Keller, *University Of Oklahoma*

I. Norton, *University Of Texas*

J. Van Wijke, *University Of Houston*

Community planning meetings

2008 Geological Society of London Meeting (Houston)

2009 Southcentral GSA (Dallas)

2009 Fall AGU Meeting Town Hall (San Francisco)

Continental Breakup and Formation of Rifted Margins: The Gulf of Mexico as a Natural Laboratory



PART I **Why the Gulf?**

Continental Breakup and Formation of Rifted Margins: The Gulf of Mexico as a Natural Laboratory



PART I **Why the Gulf?**

PART II **Tectonic Evolution Of The Gulf**

Continental Breakup and Formation of Rifted Margins: The Gulf of Mexico as a Natural Laboratory



PART I **Why the Gulf?**

PART II **Tectonic Evolution Of The Gulf**

PART III **Opportunities & Challenges**

Why the Gulf of Mexico?



- This is how continents break apart

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- **Inherited Tectonics – 2 Wilson Cycles**

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- Structural Style – Rift, transform, and oblique segments

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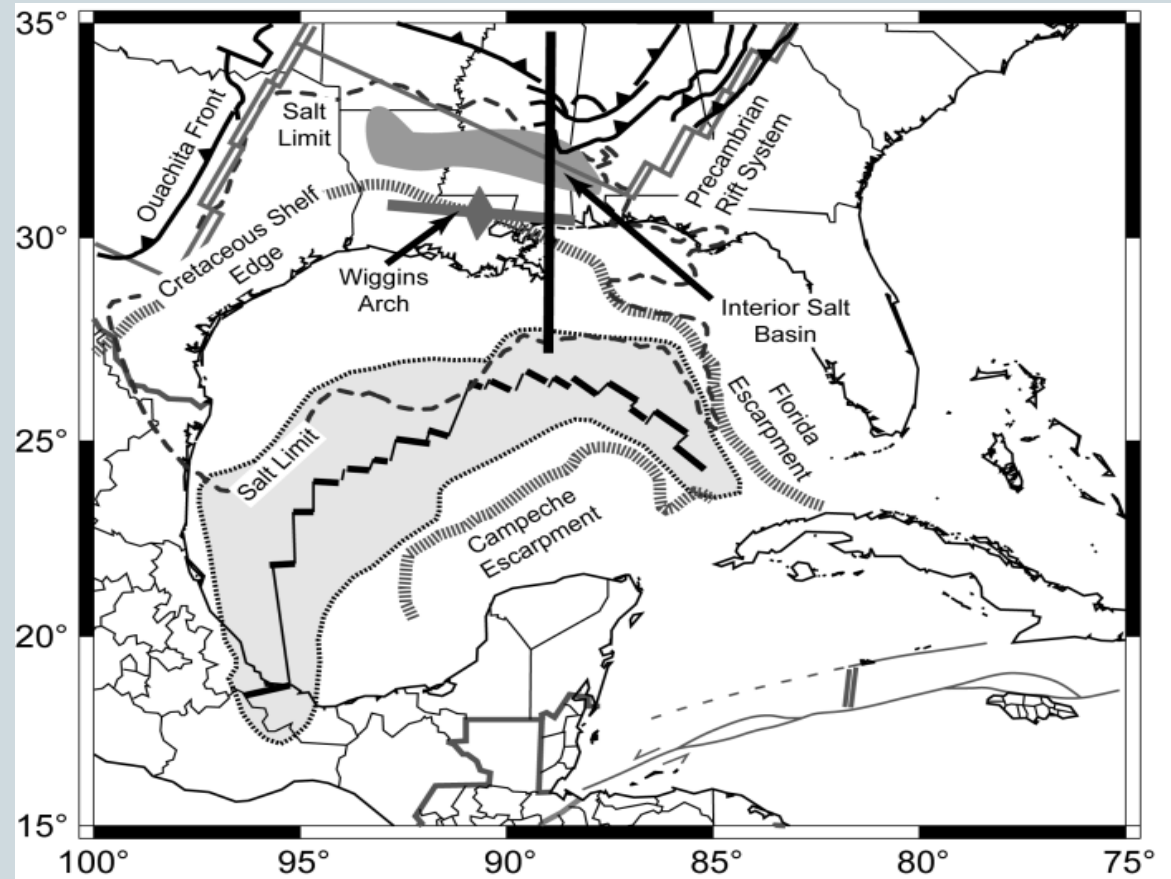
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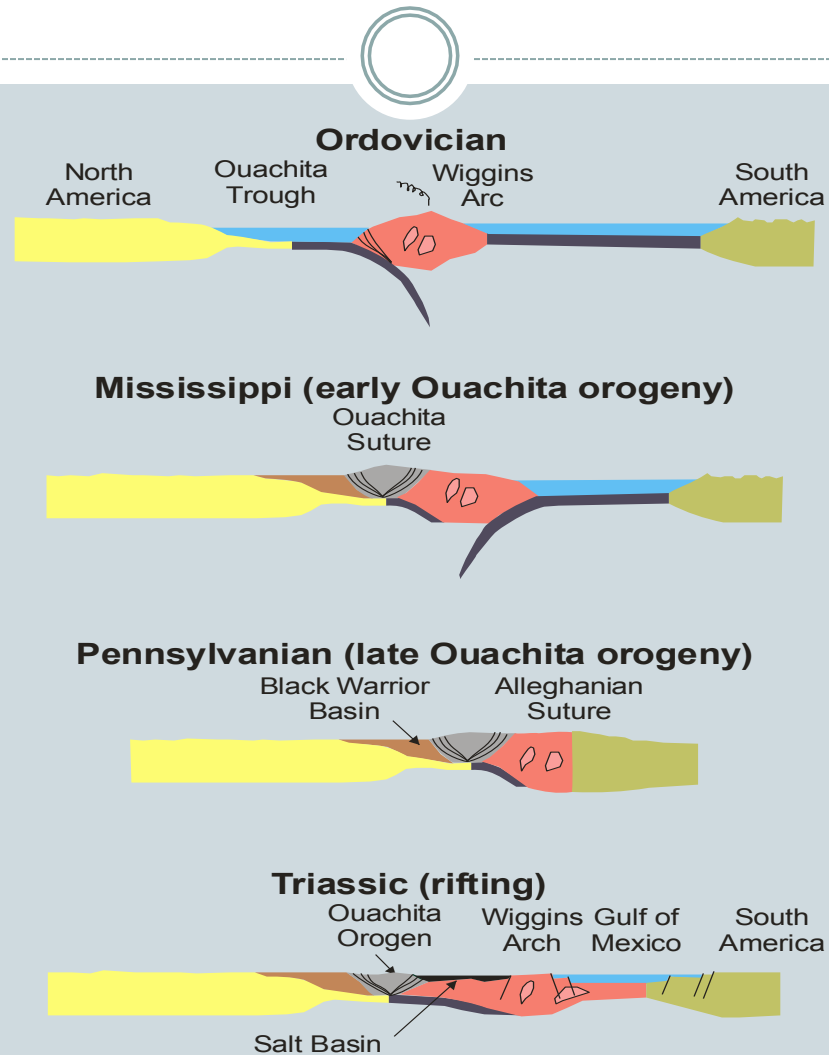
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- **Science Beyond RIE – a Mesozoic gateway between the Atlantic and Pacific Oceans**

Tectonic Evolution

- Rift initiation ~ 215 Ma (Eagle Mills and La Boca Fms.)
- Seafloor spreading began ~ 165 Ma (inferred – no direct evidence)
- Opening ceased ~ 138 Ma (inferred)
- Followed Ouachita orogenic trend (sort of)
- The widest known rifted margin

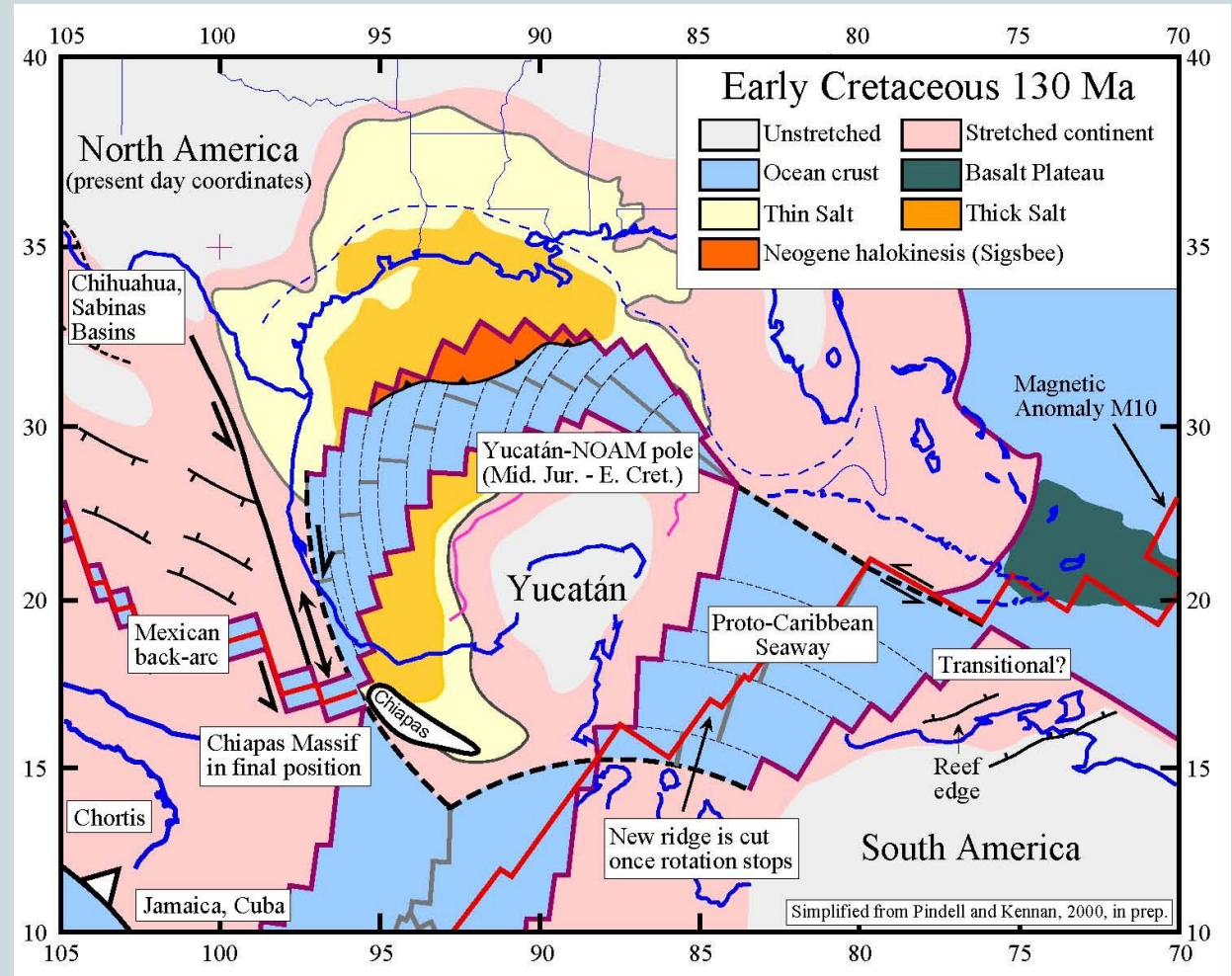


Tectonic Evolution



Tectonic Evolution

- Counterclockwise rotation of Yucatan away from North America
- Transform motion between Yucatan and eastern central Mexico
- Extreme stretching on central North American margin, oblique to trend of Precambrian transform margin and Paleozoic Ouachita orogen



Opportunities



- **Tectonic Inheritance** –The Gulf of Mexico provides an opportunity to compare rifting of a “soft” collision orogen in the central and eastern Gulf and a “hard” collision in the western Gulf: thin-skinned vs. thick-skinned.

Opportunities



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- **Fluid Evolution and Migration.** The GOM is a factory for generating a wide variety of fluids: CO₂, brines, and hydrocarbons. The GOM provides an opportunity for understanding how these fluids form, migrate, and interact.

Challenges



- Dodging industry and ship traffic

Challenges



- Dodging industry and ship traffic
- Access to international data & field sites

Challenges



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- **Transportable array is moving on**

Challenges



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Challenges



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- Great depth to basement & lack of exposure

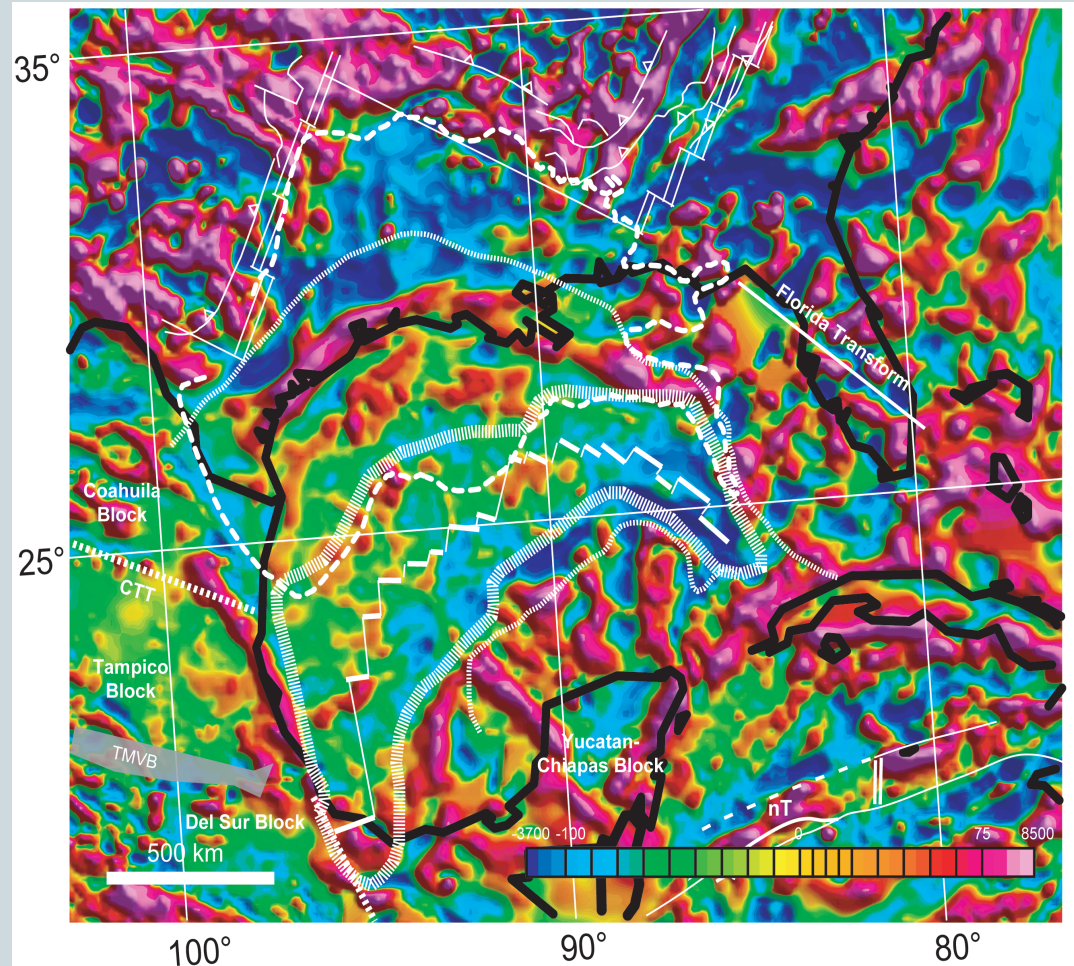
Challenges



- Dodging industry and ship traffic
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- Imaging through salt & separating Salt Tectonics from Rift Tectonics

Tectonic Evolution

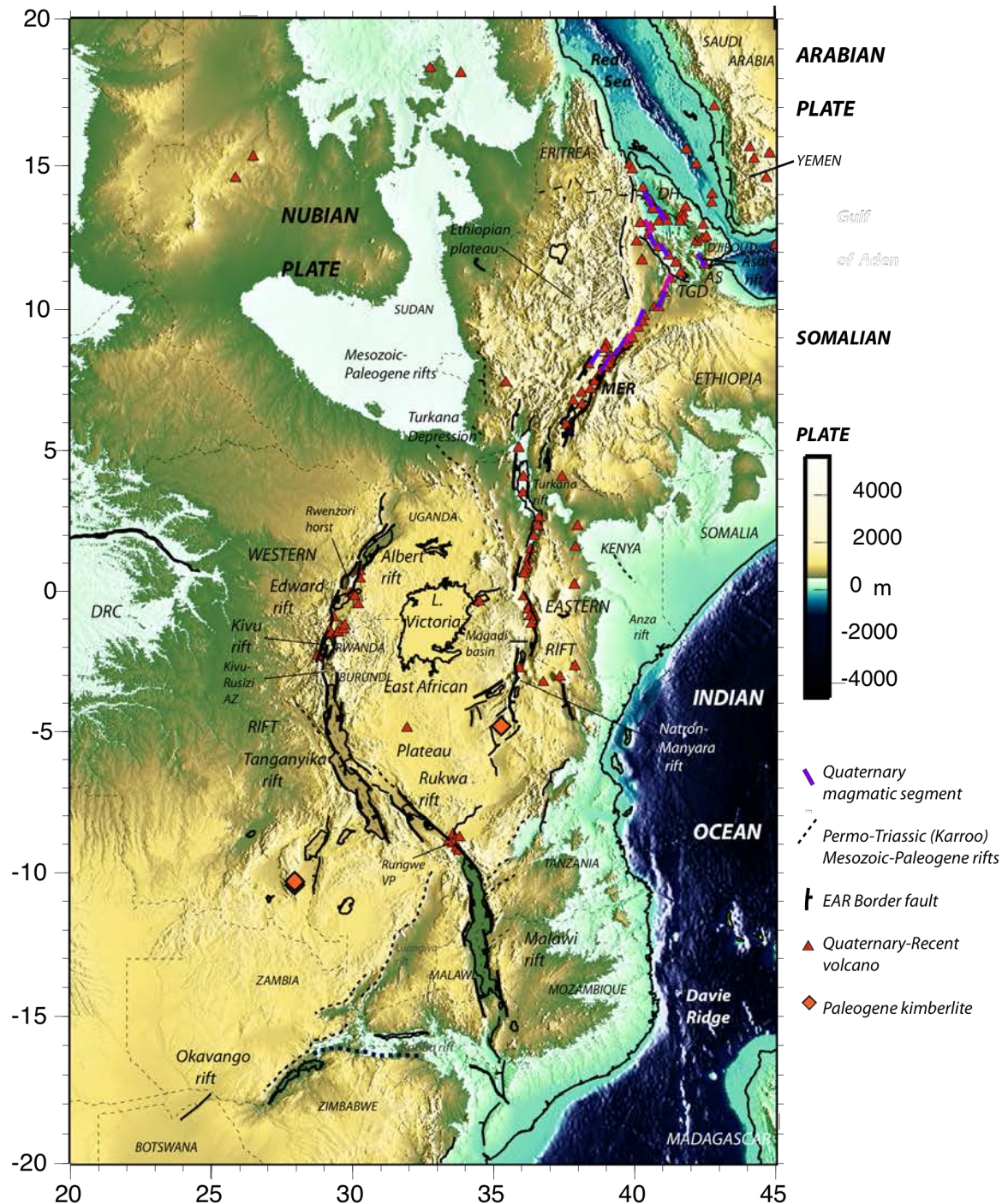
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The East African Rift System

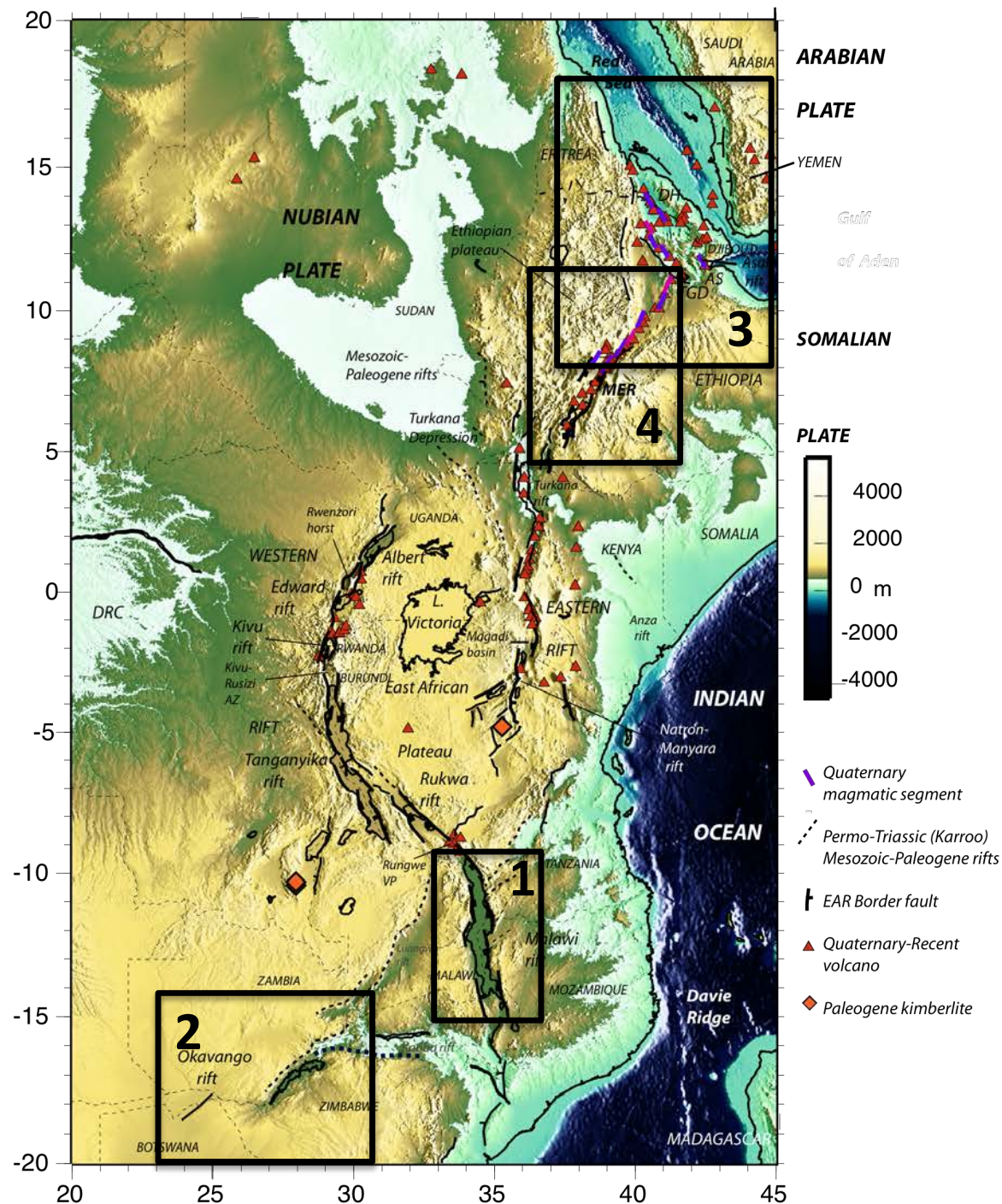
A possible primary or thematic site

Atekwana et al.
Ebinger
Gaherty et al.
Reilinger et al.
Rooney et al.



Active continental rift

- Strain rates of 1-6 mm/yr
- Northern end – Afar plume
- Lithospheric blocks (e.g. Tanzanian, Kaapval cratons)
- Pan-African orogeny, older rifting events
- Volcanic, tectonic

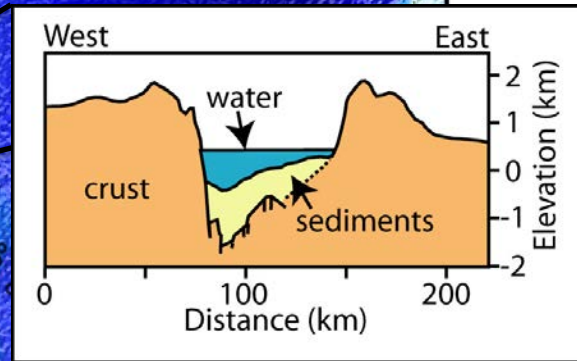


**East Africa Rift:
not so much
magma
everywhere?**

Only surface
expressions of
magma in the
Western Rift

km
0 500

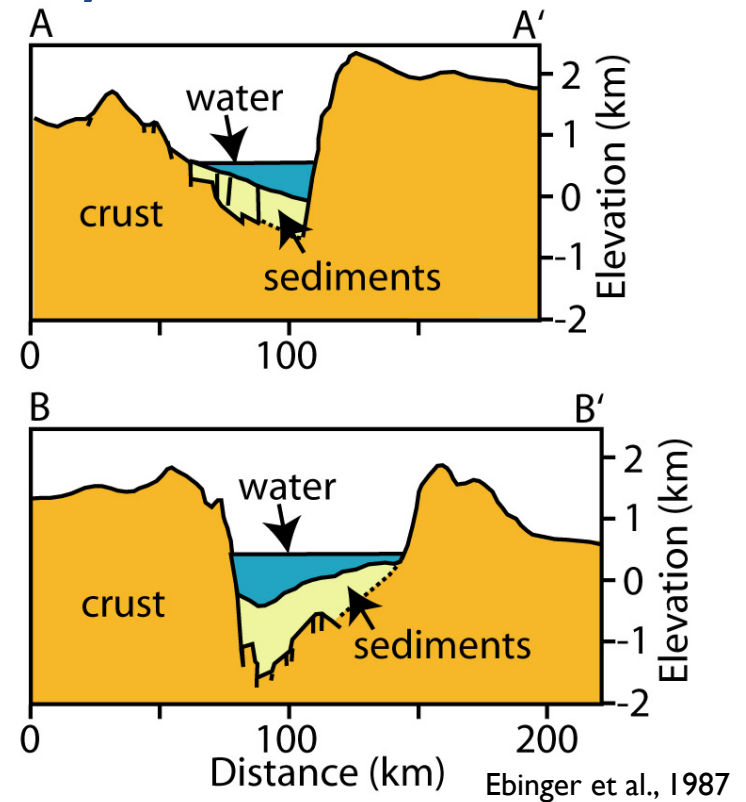
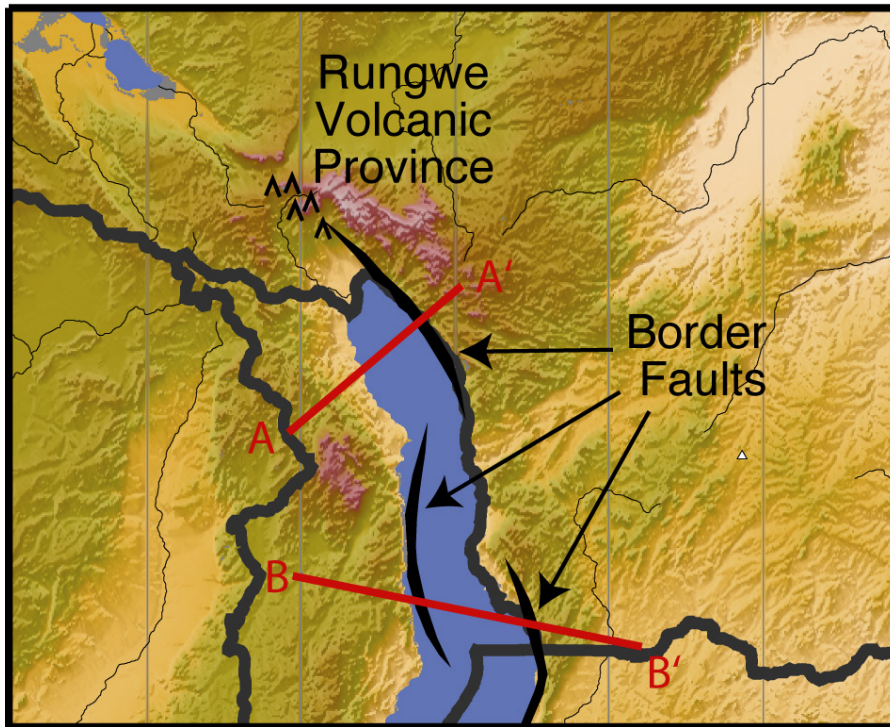
Indian
Ocean



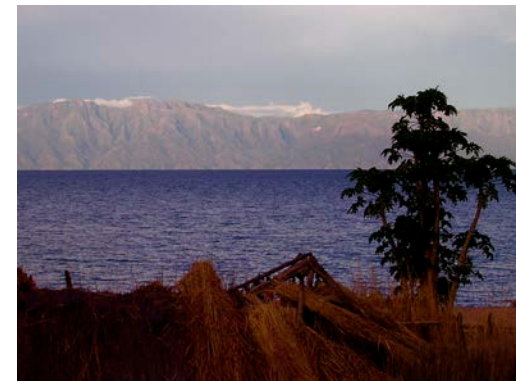
Ebinger et al., 1987

White paper by Gaherty et al.

Spatial patterns during early extension: Western Rift, East Africa Rift System

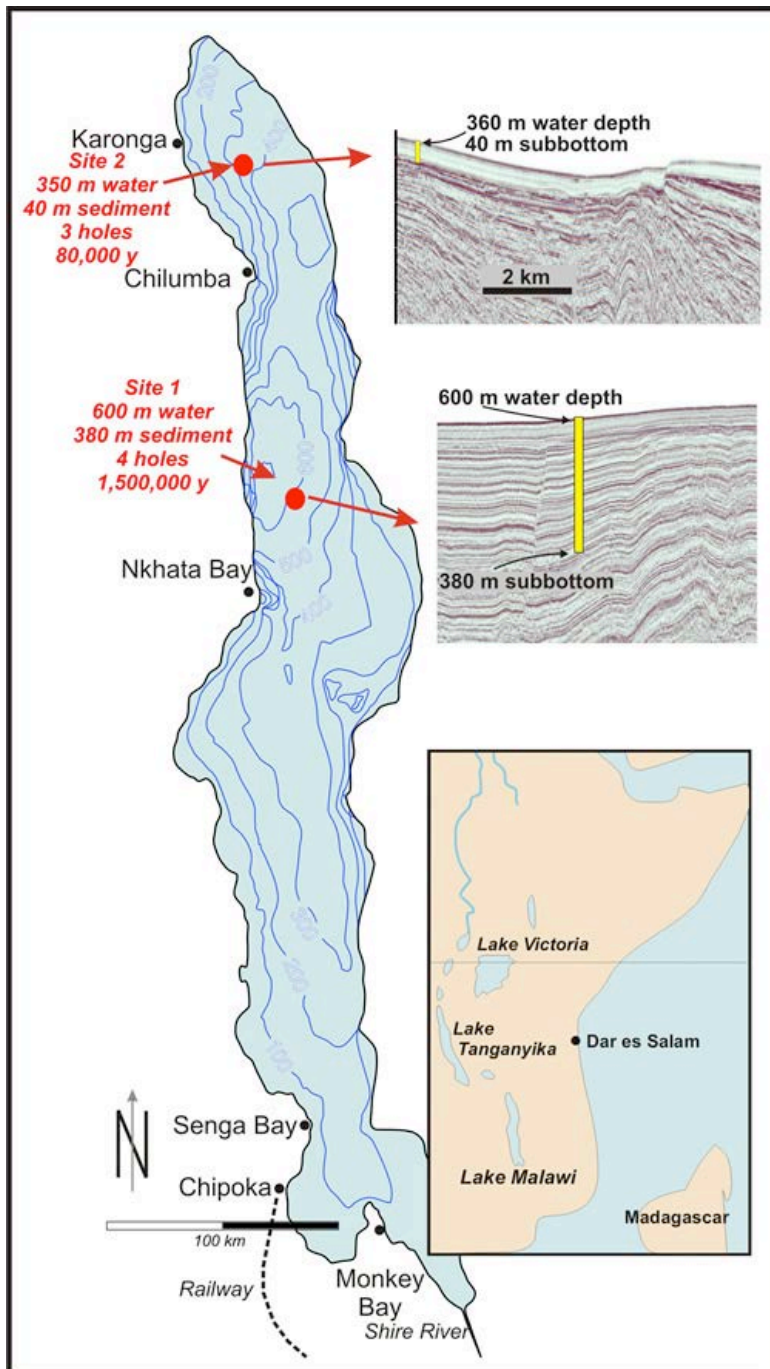


- Pronounced tectonic segmentation at the surface defined by ~100-km-long border faults and accommodation zones



Western branch:

Sedimentary record in lakes records tectonics and climate change



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B.R. ROSEND AHL ET AL.

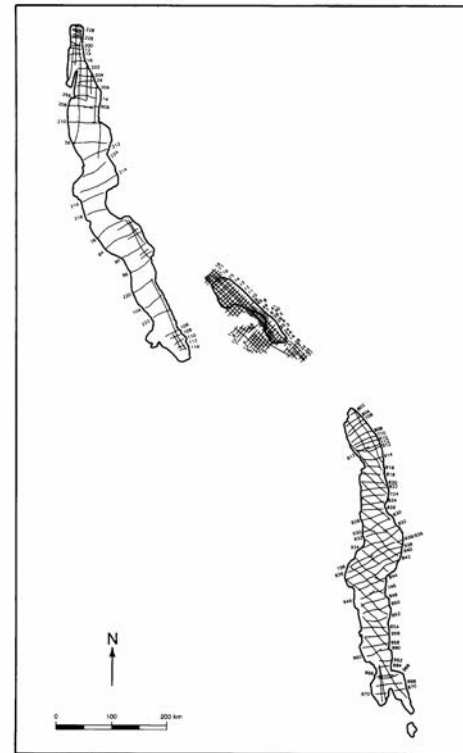
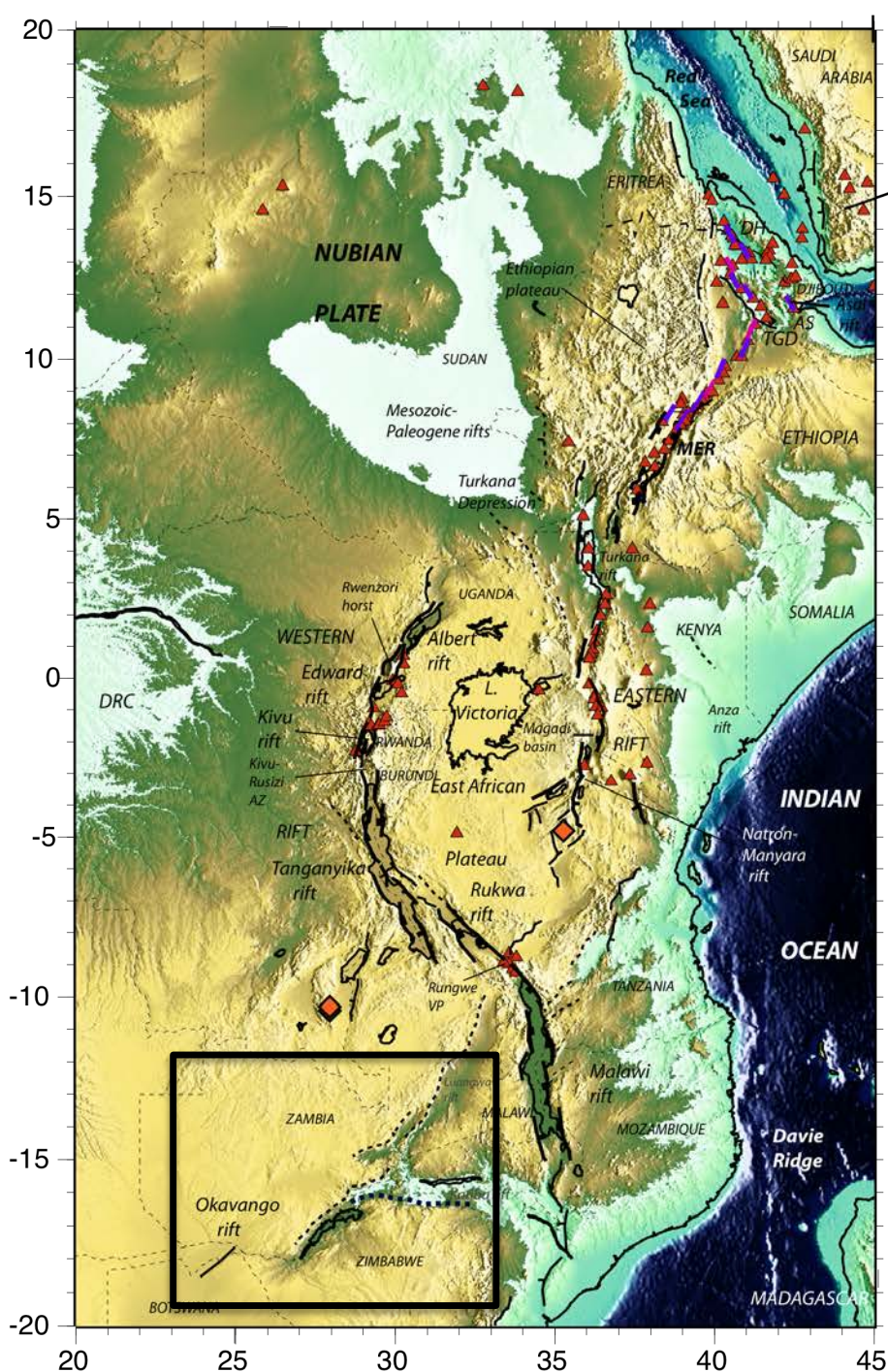


Fig. 2. Map portfolio of the Tanganyika, Rukwa and Malawi Rift Zones. (A) Multifield seismic reflection coverage used in analysis. Lines with numbers refer to line drawings shown in Fig. 3. (B) Pliocene-Tertiary, Cretaceous and Cenozoic sedimentary cover associated with known rift basins. (C) Structure map on acoustic basement. Compiled from Rosendahl (1987), Sholz and Rosendahl (1988), Versfelt and Rosendahl (1989) and Scott et al. (in press). (D) Possible pull-apart geometry linking Kalemie, Rukwa and Livingstone Basins.

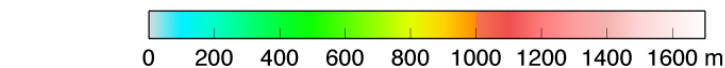
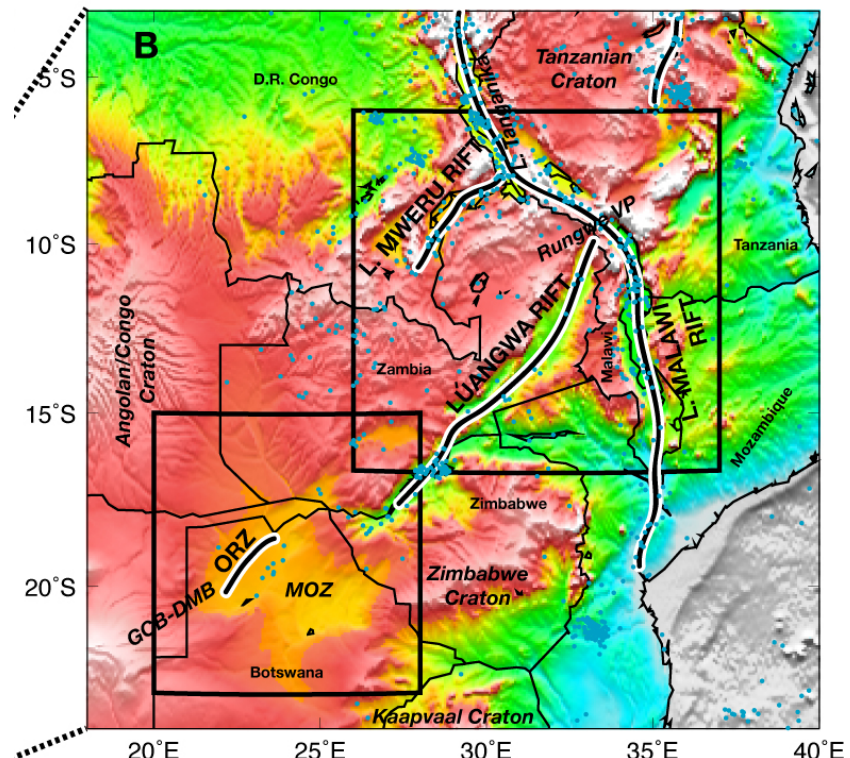
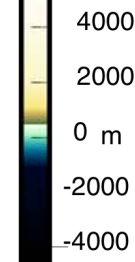
Rosendahl et al., 1992



**ARABIAN
PLATE**

**SOMALIAN
PLATE**

PLATE



East Africa Rift System seismic stations from previous and on-going studies

seismicity (ANSS catalogue)

ORZ: Okavango Rift Zone

MOZ: Makgadikgadi-Okavango-Zambezi Basin

GCB-DMB: Ghanzi-Chobe/Damara Mobile Belts

Fault scarp height comparison

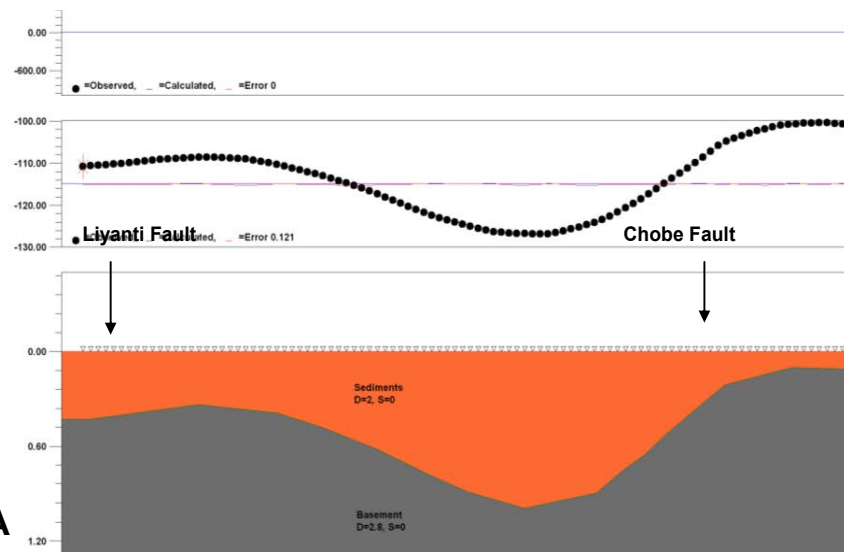
The SWB is geologically less evolved than either the eastern or western branches.



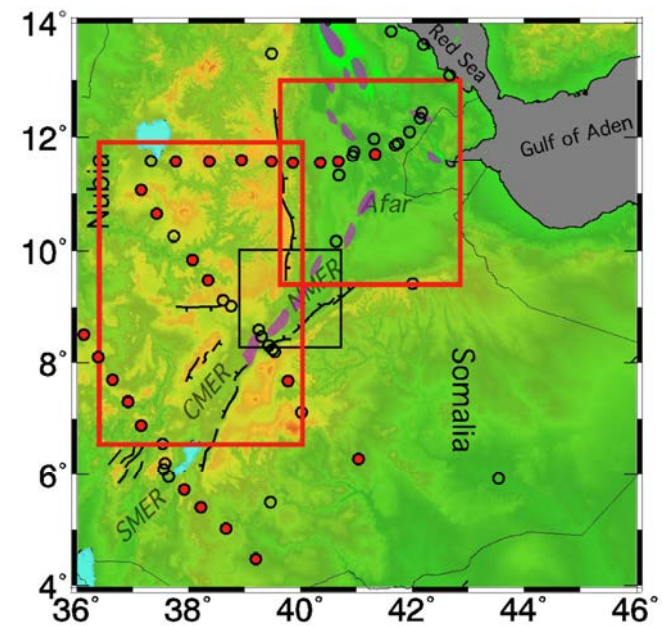
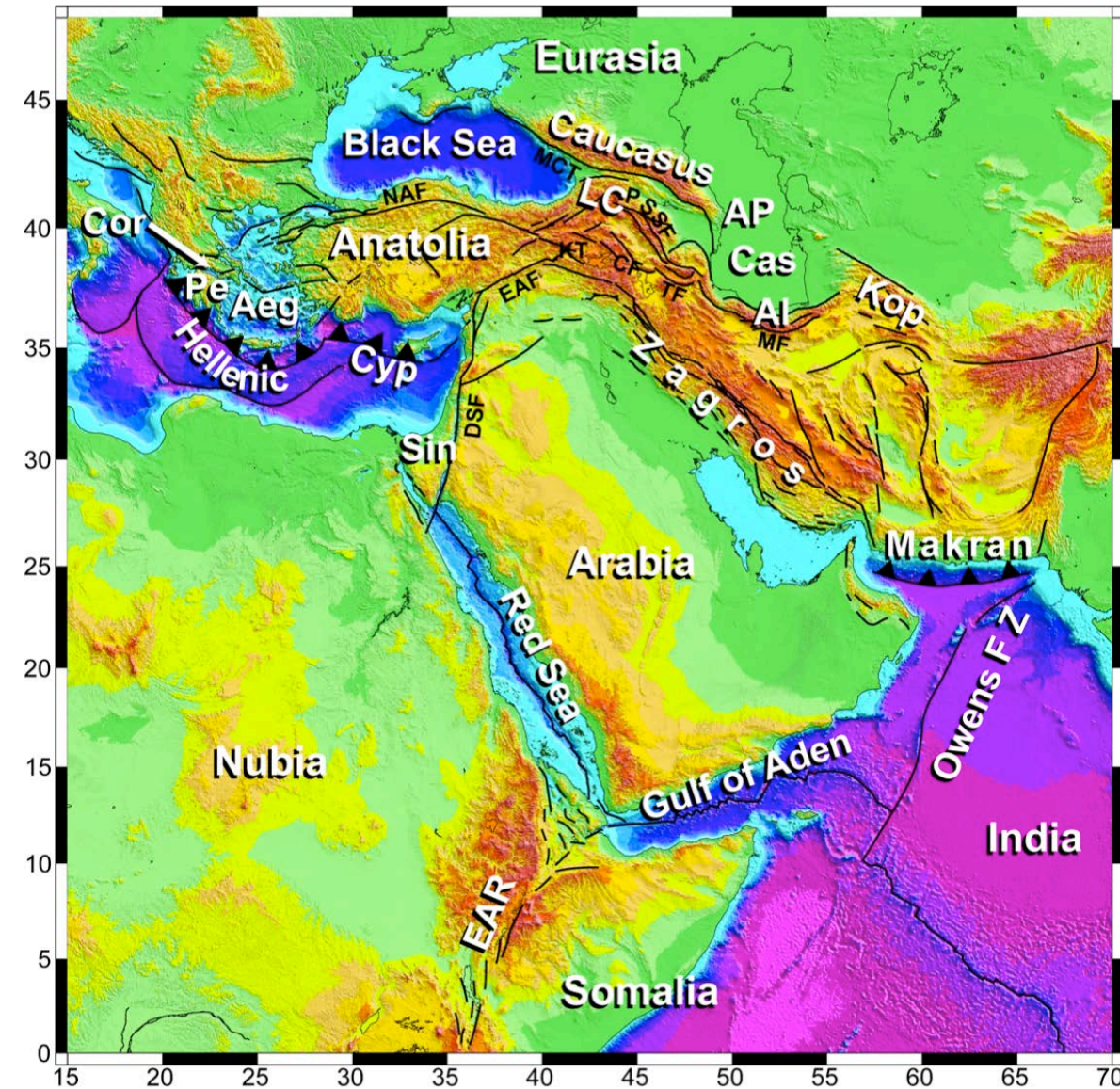
Livingstone Fault > 1 km –
L. Malawi



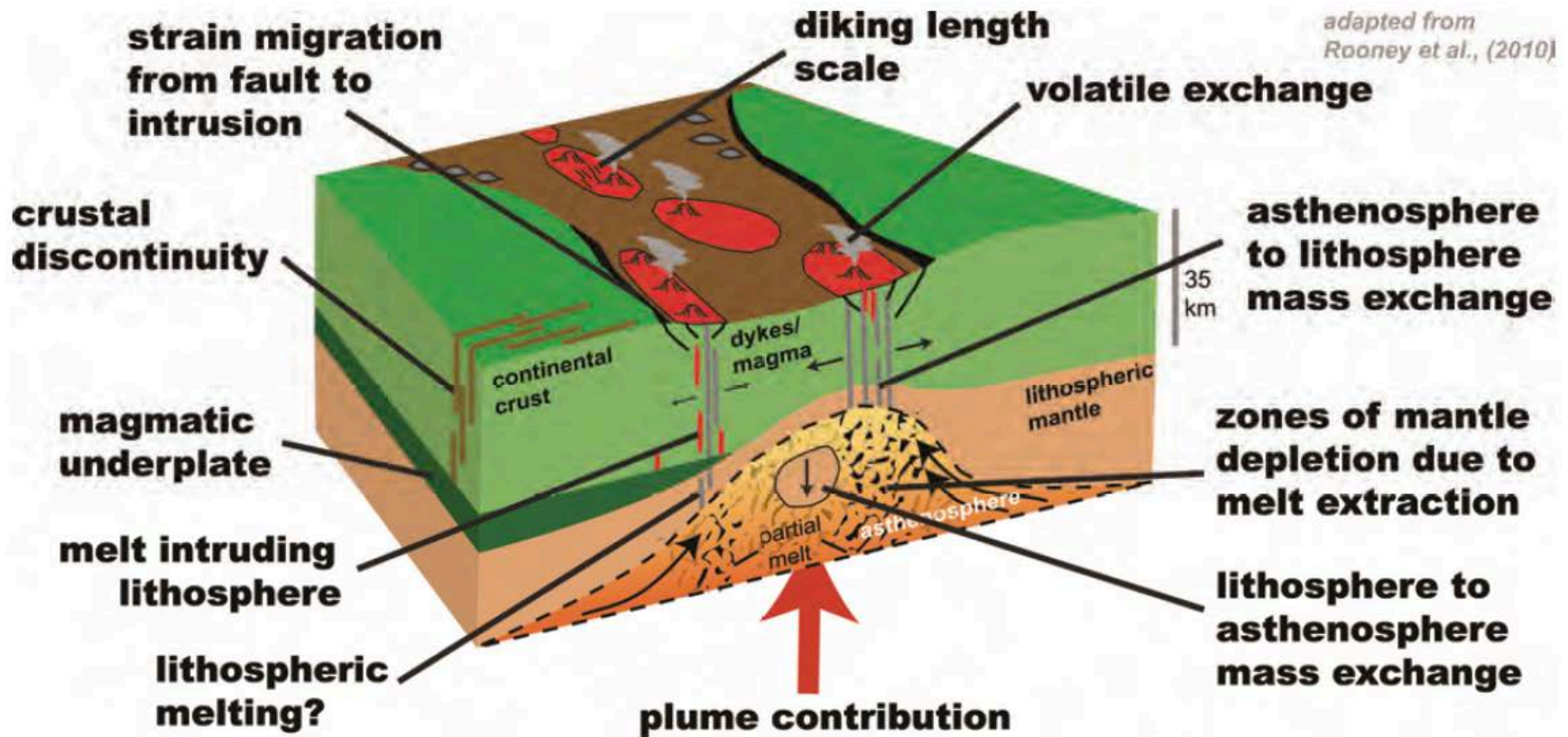
Mweru Fault– 50-200 m –
L. Mweru



Afar triple junction



White paper by Reilinger et al.

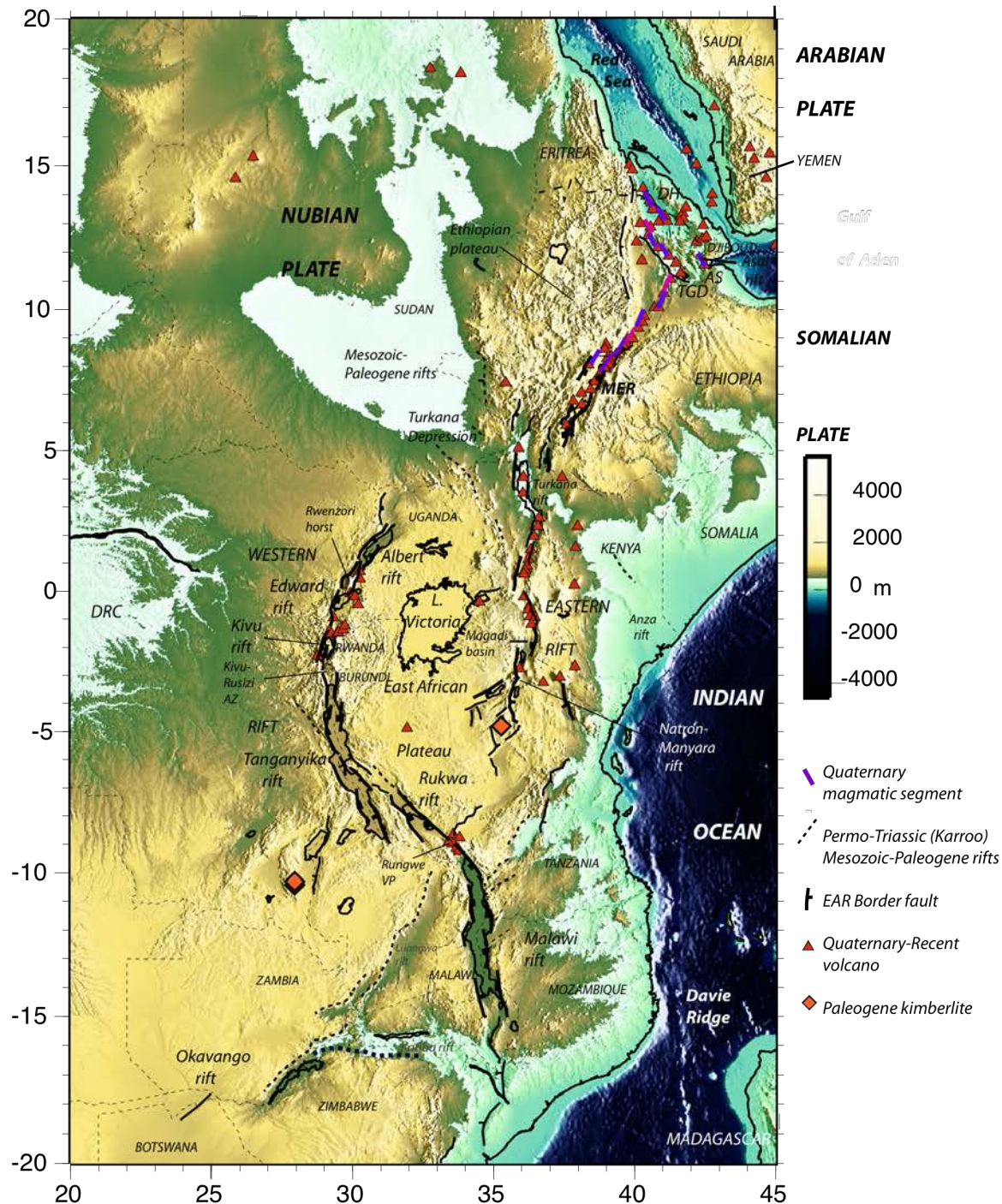


- Opportunity to characterize magma reservoirs at different crustal levels; spatial variability
- Combine geochemistry, geology, and geophysics

White paper by Rooney et al.

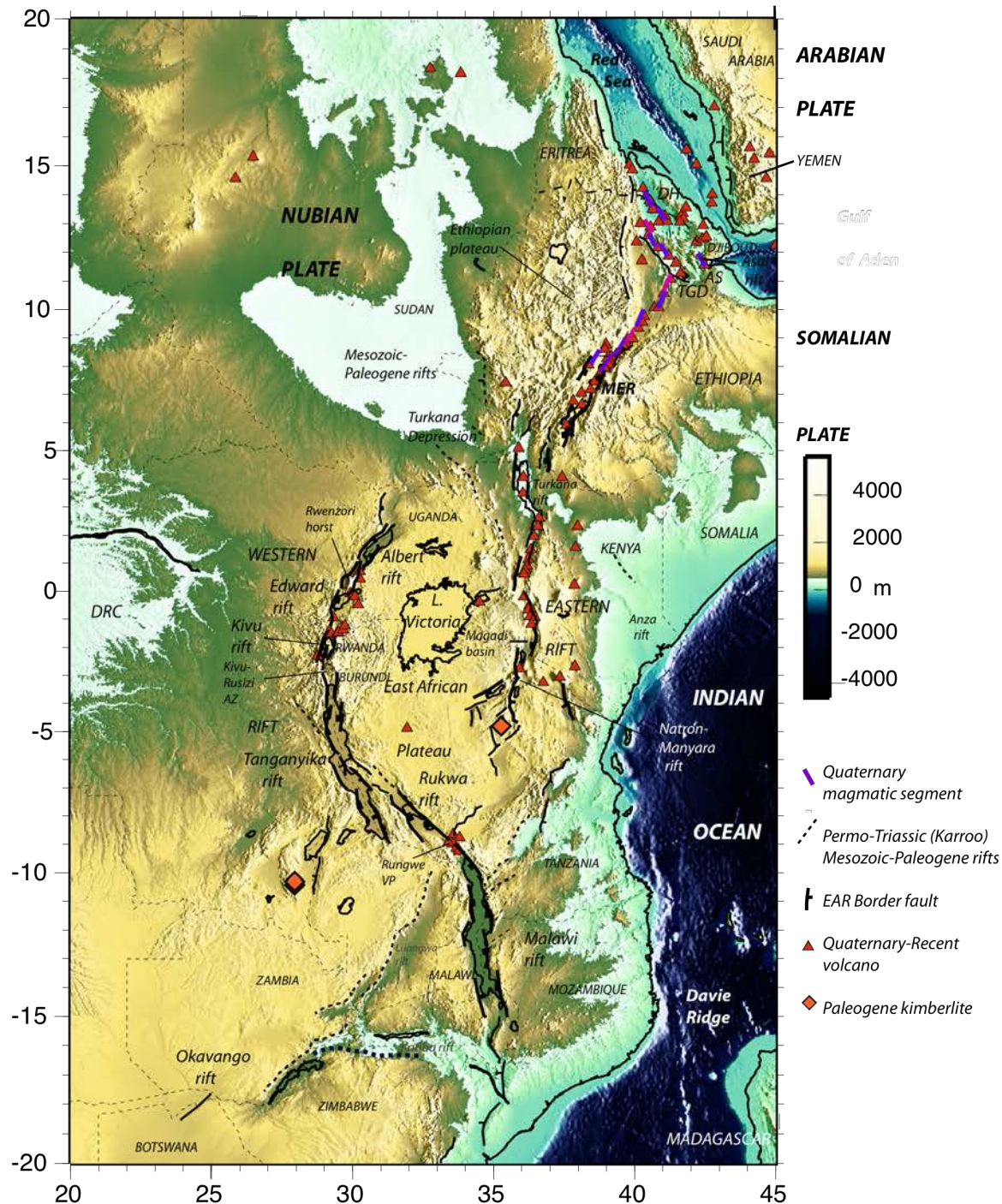
Also: NW branch, Kenya, Eastern Tanzania

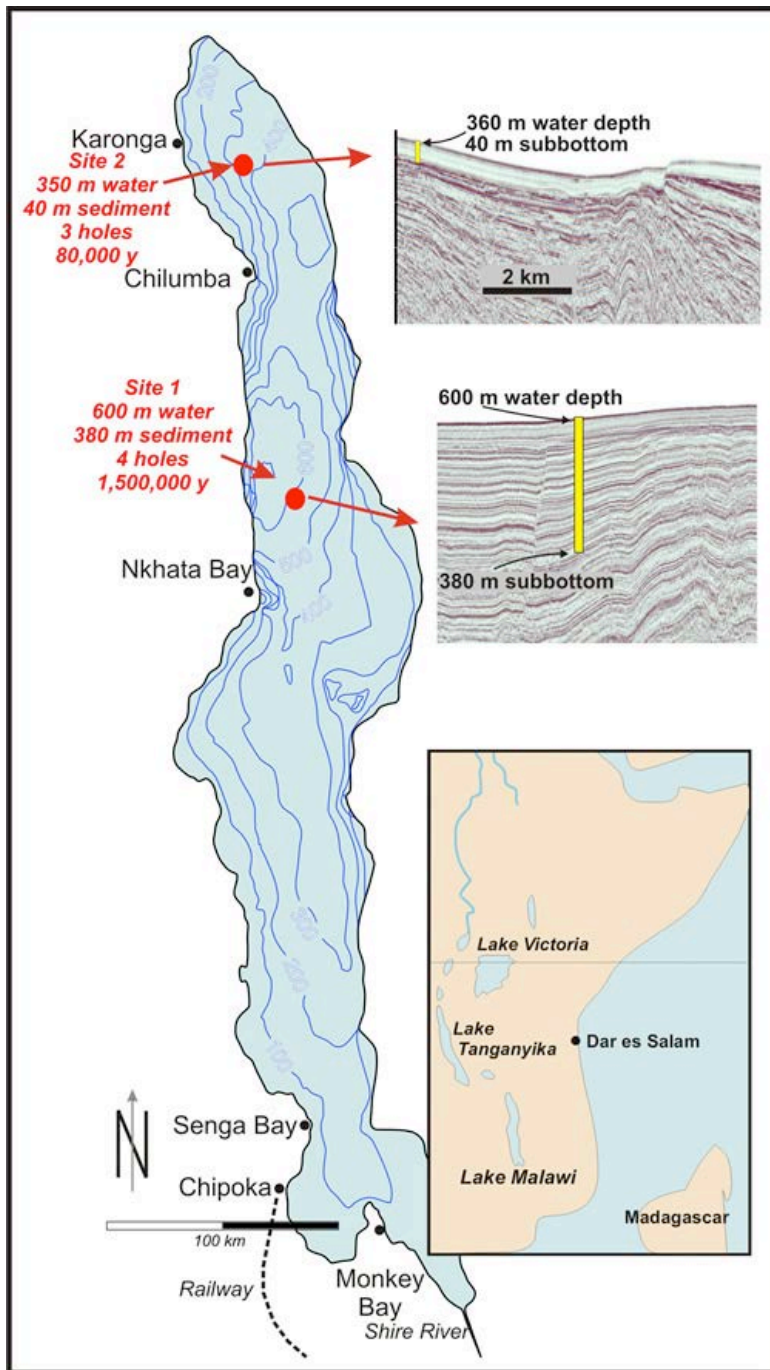
- Nyamulagira, Lake Kivu, etc.
- Resources: Petroleum, geothermal
- Deep earthquakes, sedimentation records
- Intermediate



Spatial variability; change in rift maturity

- Spectrum of fault system structures and magmatic influence
- Botswana to Afar
 - Incipient rifts to rift grabens to diking
 - Rift initiation with large faults, no surface volcanism
 - Plume, no plume
- Kenya to Rwanda
- **Comparative studies within one system**

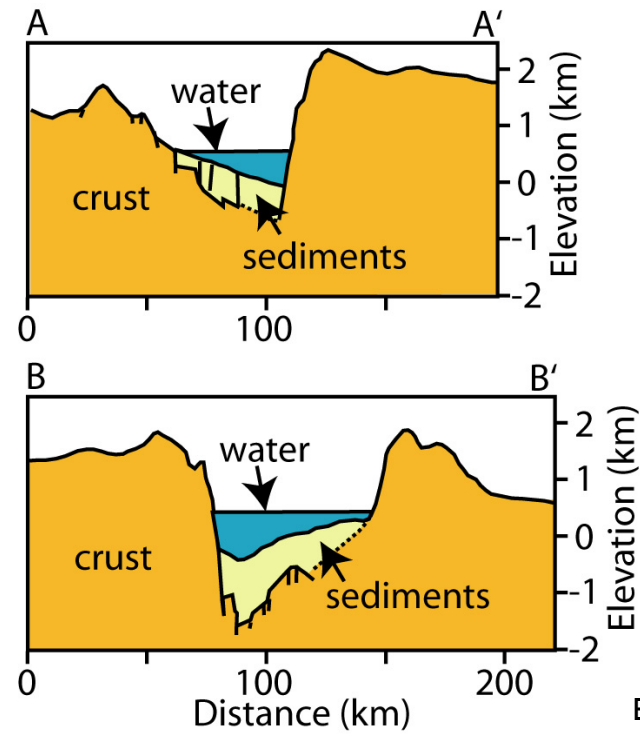




Fault-bounded basins along the length of the rift

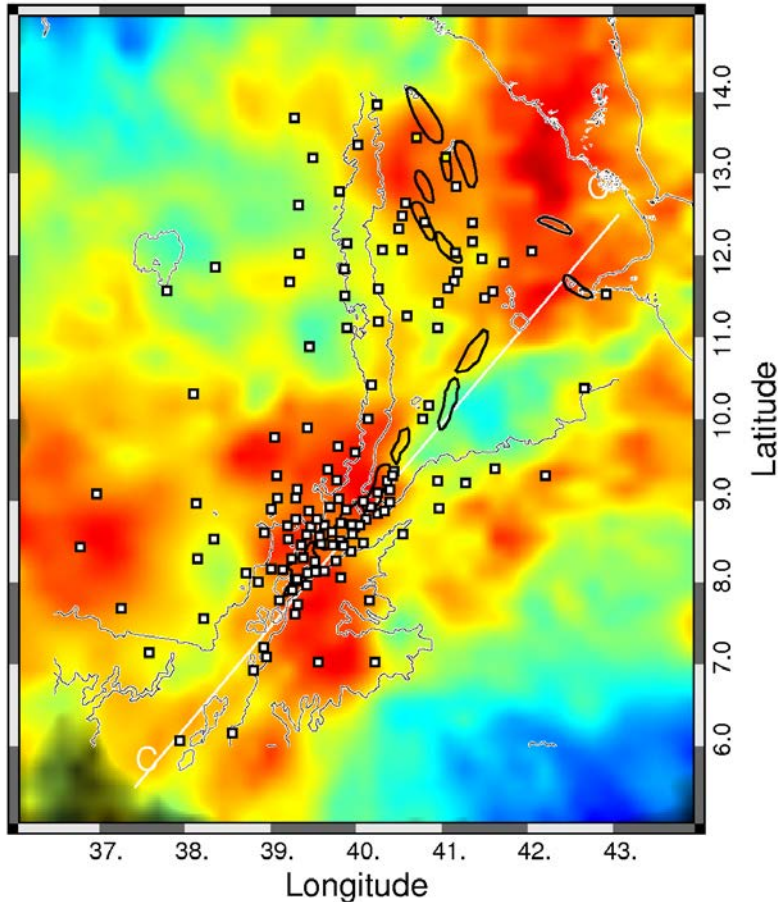
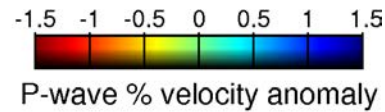
Sedimentary record of deformation and climate; represent all stages of rift evolution

Feedbacks between faulting, flank uplift, sedimentation, further deformation

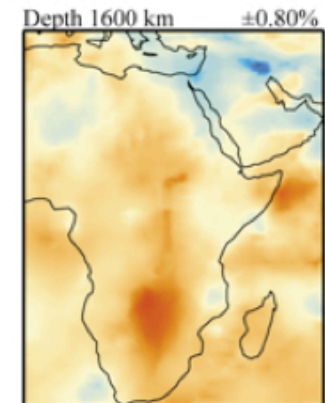
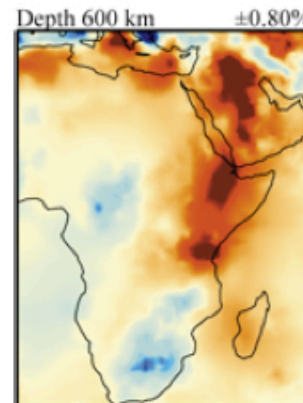
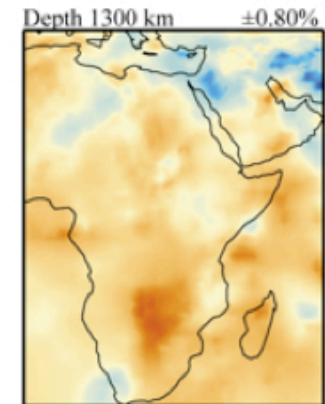
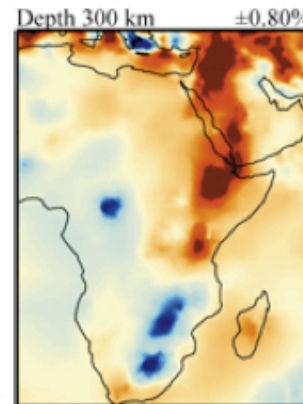
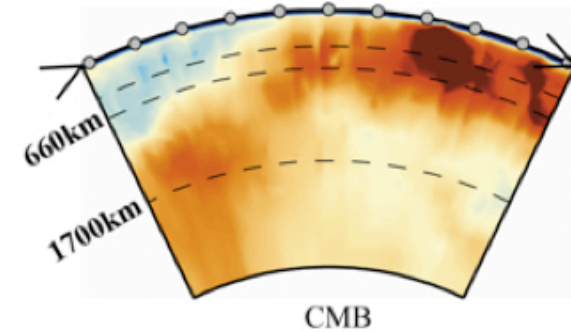


Backbone geophysics: Mantle tomography

depth =
550 km

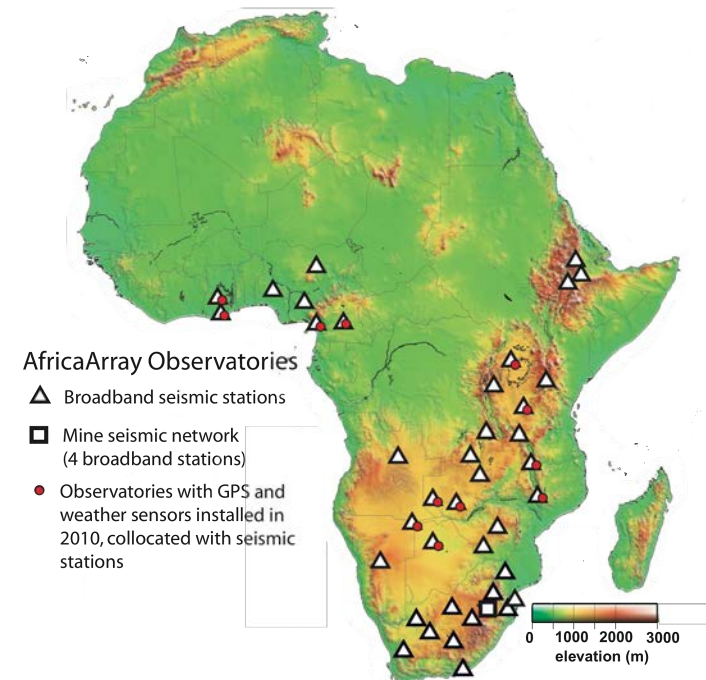
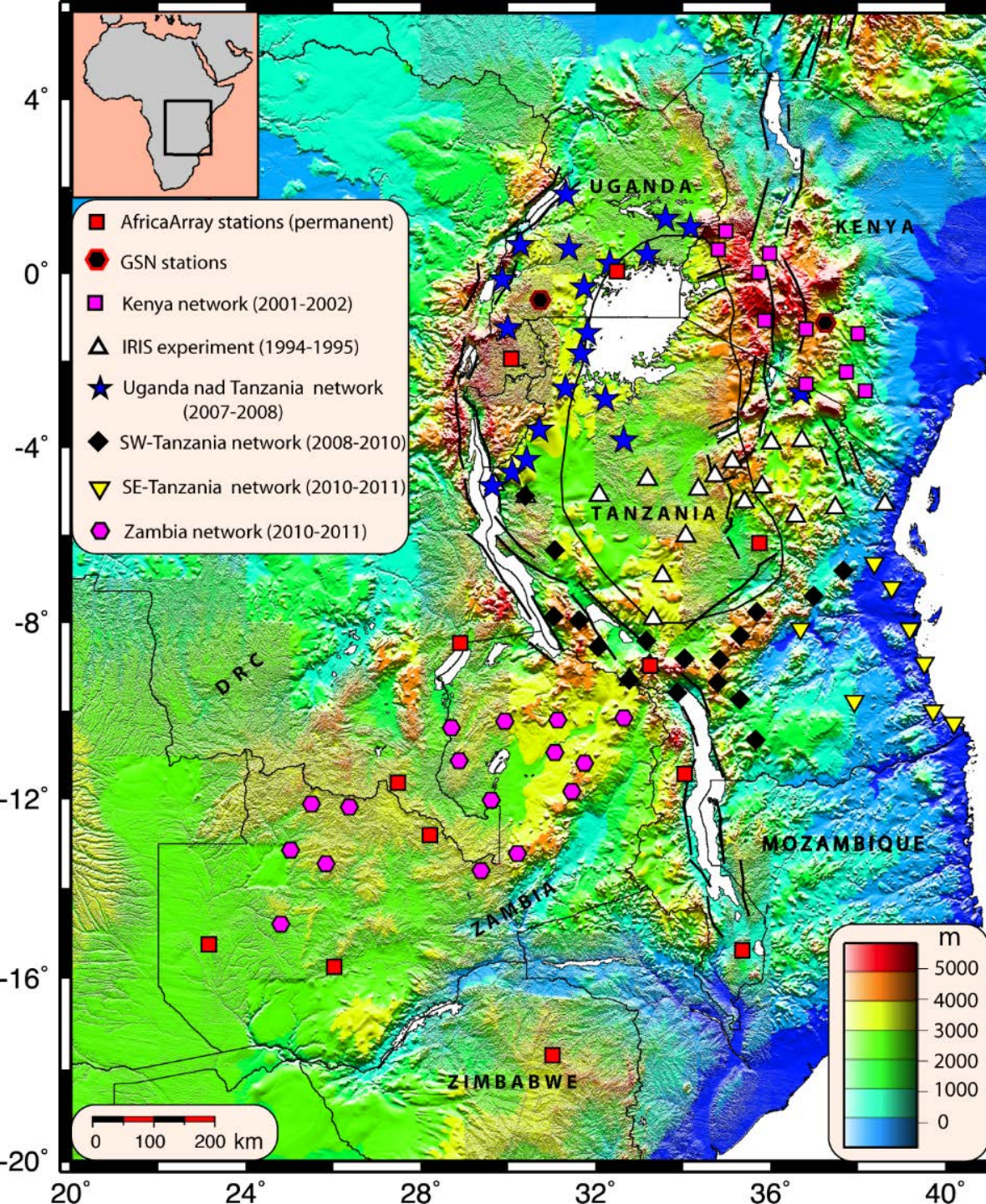


Africa

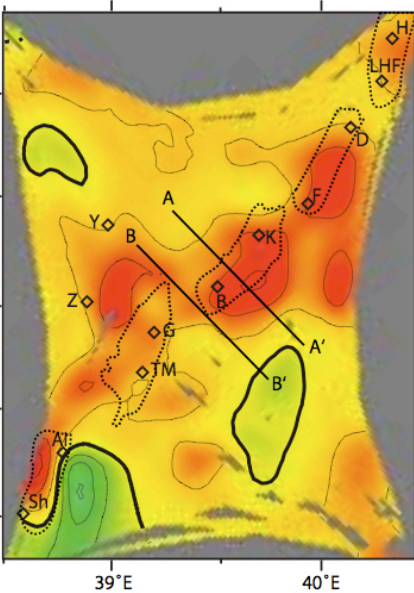


AfricaArray temporary and permanent seismic networks

- 1) Uganda/NW Tanzania
8/07-12/08
- 2) Southern Tanzania
1/09-7/10/8/10-8/11
- 4) SE Tanzania 2/10-3/11

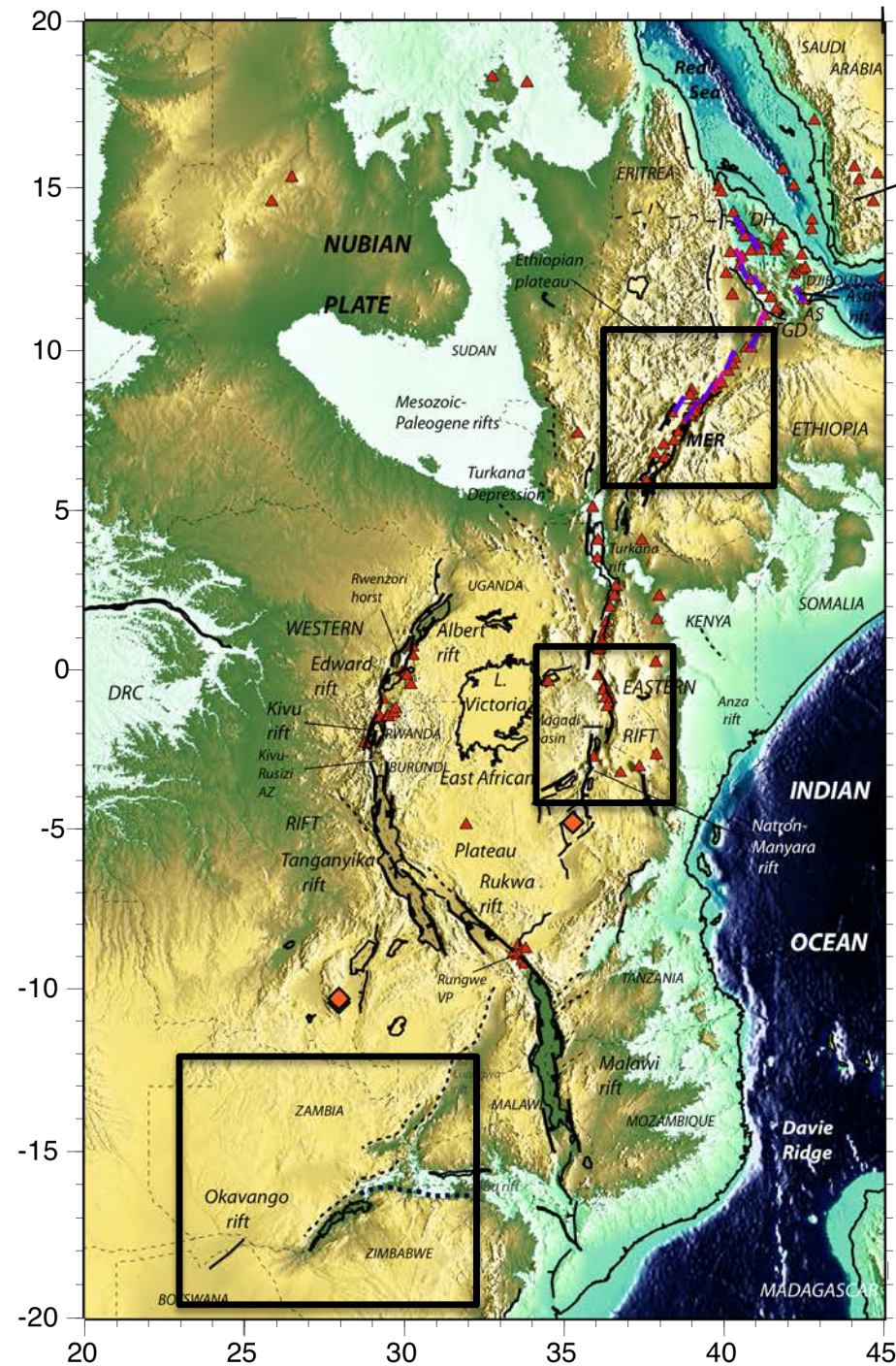
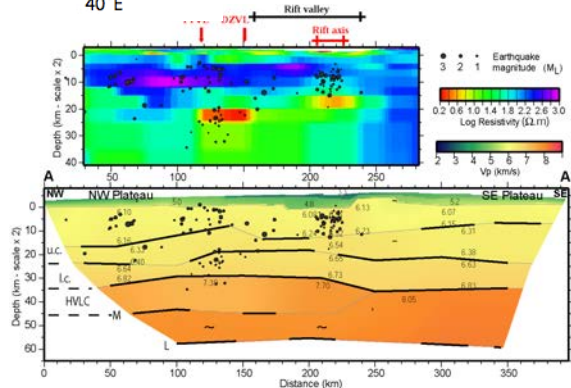


Backbone geophysics: Crustal seismic studies, MT



KRISP – Kenya Rift
EAGLE - Ethiopia
SAMTEX – S.Africa
CD Project – SW Rift*

Gravity, magnetics



Summary – ability to address science objectives of RIE within the EARS

- Where and when do continental rifts initiate?
- How do rift processes and feedbacks evolve in time and space?
- What controls the structural and stratigraphic architecture before and after breakup?
- What are the mechanisms and consequences of fluid and volatile exchange?

Summary – logistics, leveraging

- Amphibious (?)
 - Sub-aerially exposed, but crosses from continental to oceanic crust
- Readiness
 - Significant backbone geophysics, **Africa Array**
 - Existing ancillary studies, *but also a great opportunity for more work (immediate, long-term)*
- Accessibility and safety
- Availability of infrastructure
 - No EarthScope; Africa Array
- Foreign resources and collaboration
 - Strong existing relationships with African scientists at universities, geological surveys, etc.
 - Collaborations with European scientists working in East Africa
- Broader impacts
 - Geohazards – faulting (e.g. Malawi), volcanoes, CO2 emissions (Kivu)
 - Resources: petroleum, geothermal
 - International field experience and community-building

What can GeoPrisms do for East Africa?

- Bring together loosely-linked groups working on related problems throughout the rift system
 - Develop a strong community; enhance research results; leverage ongoing work
- Bring a new group of scientists, new methods, new enthusiasm
 - Fill gaps in geochron, paleoseismology, fault linkages, magmatic volumes, etc. to test and develop models of rift processes; rift hazards

