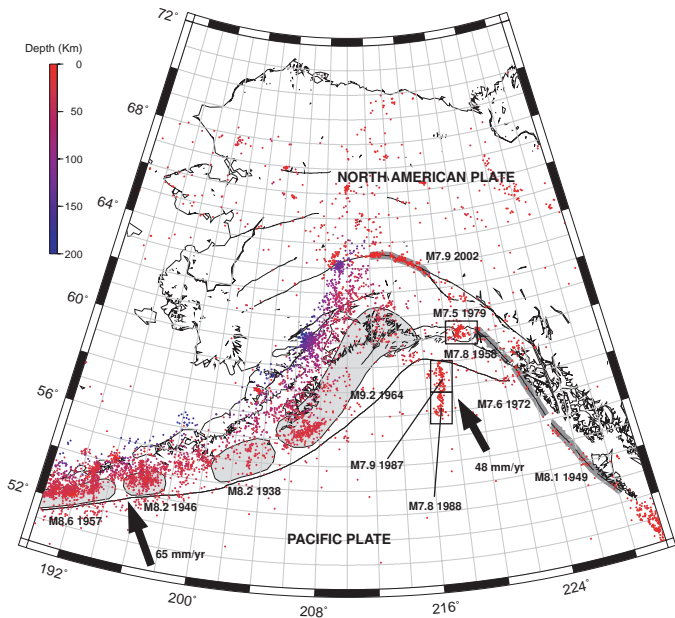


Contemporary deformation in southcentral Alaska

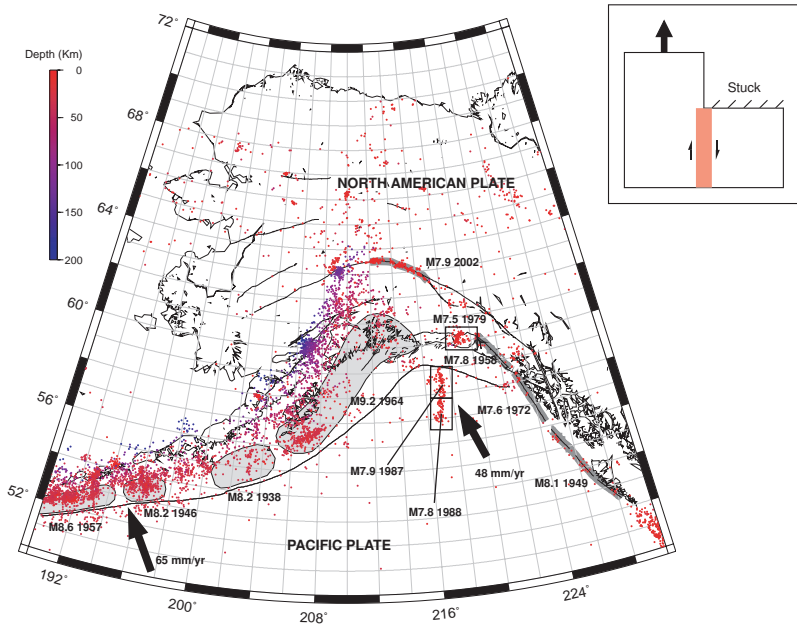
Tabrez Ali

University of Wisconsin - Madison

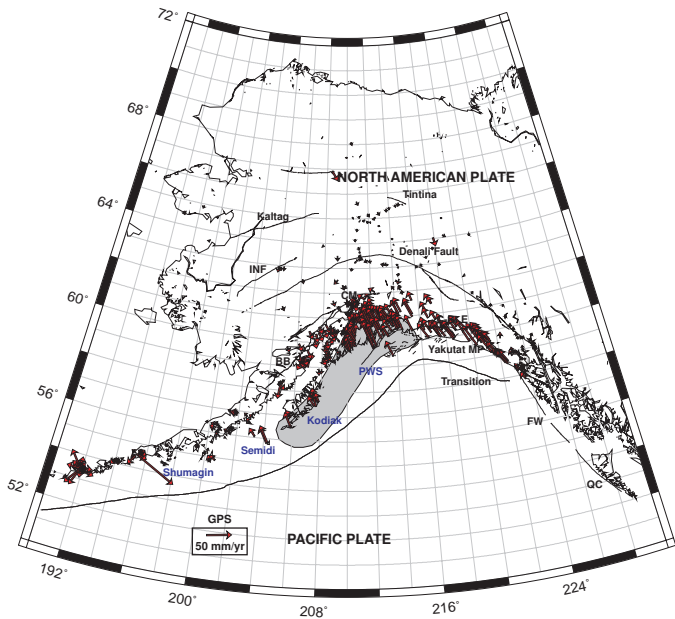
Alaska: Seismicity



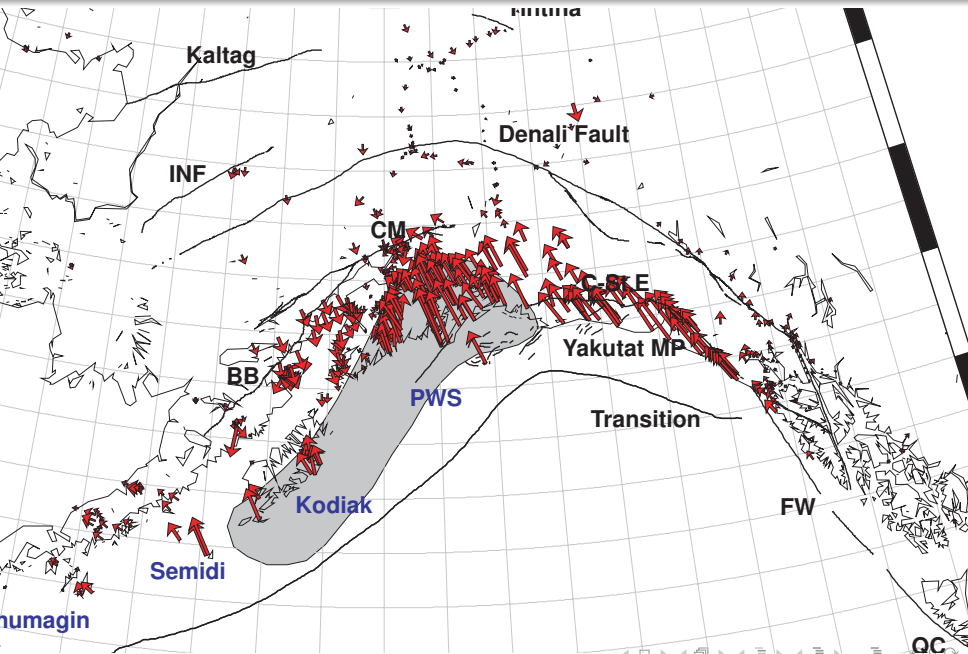
Alaska: Seismicity



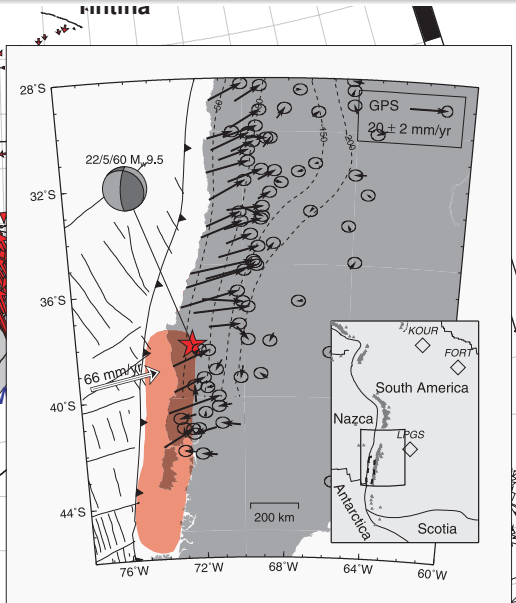
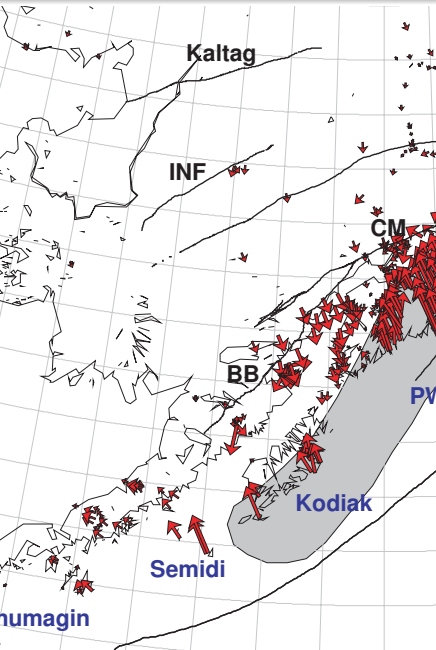
GPS



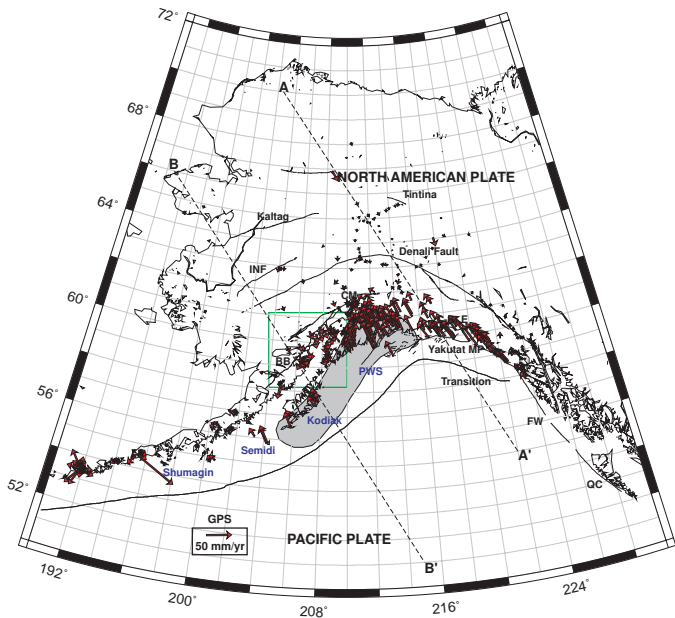
GPS



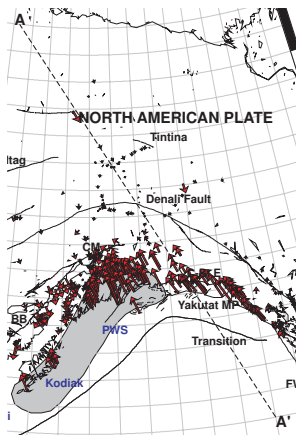
GPS



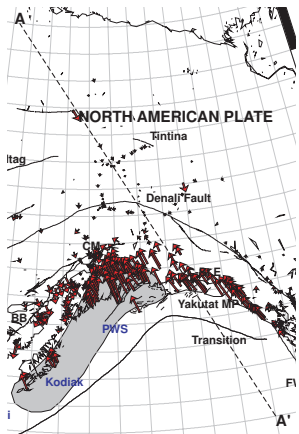
Cross-sections used for the 2D models



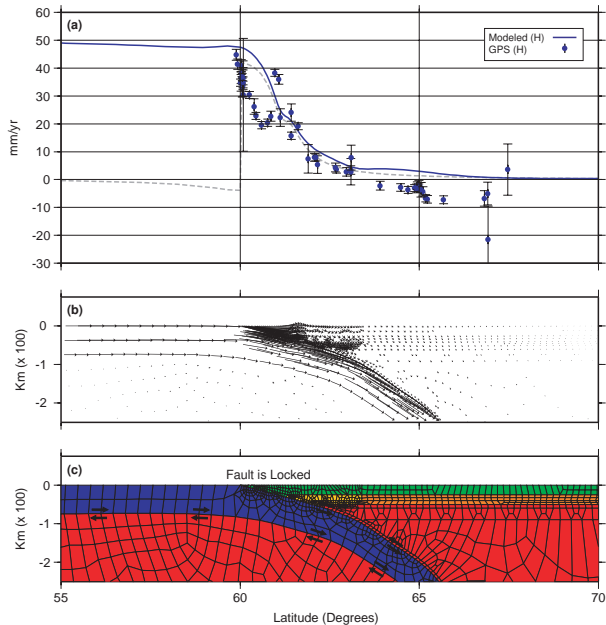
Interseismic deformation rate along A-A'



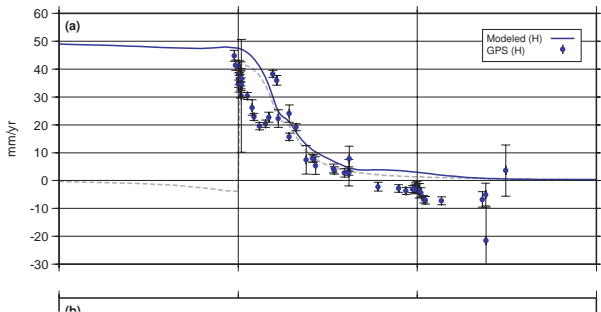
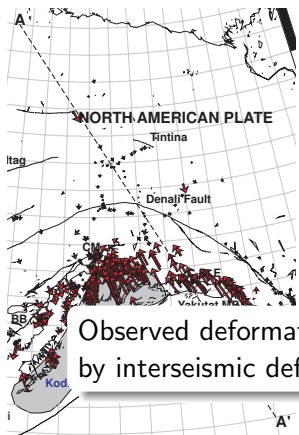
Interseismic deformation rate along A-A'



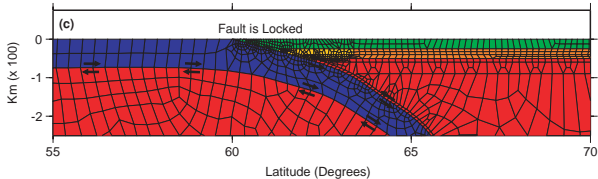
Modified Backslip Model



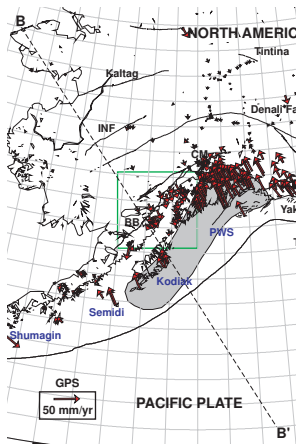
Interseismic deformation rate along A-A'



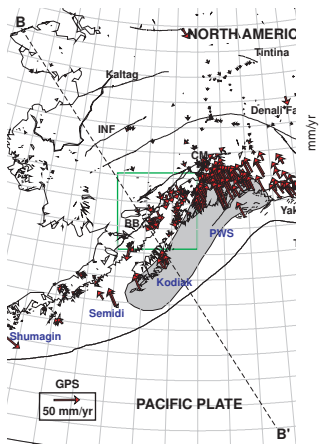
Observed deformation along A-A' can largely be explained by interseismic deformation alone



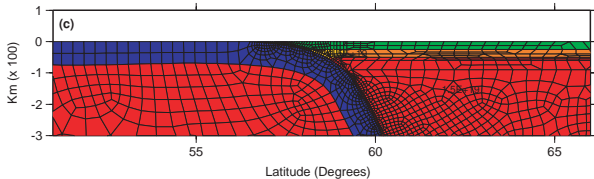
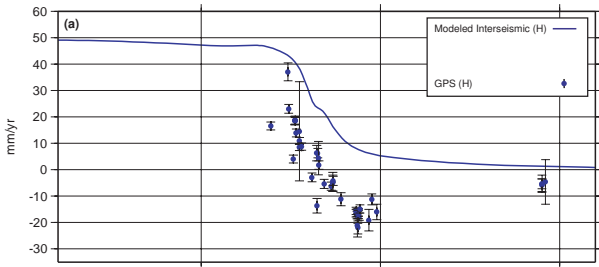
Modeling deformation rate along B-B'



Modeling deformation rate along B-B'

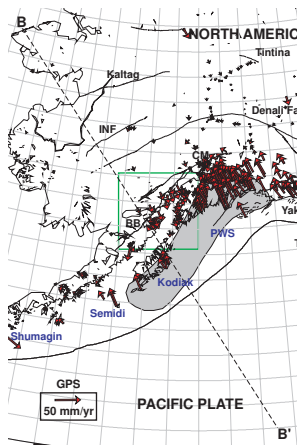


Interseismic
deformation rate

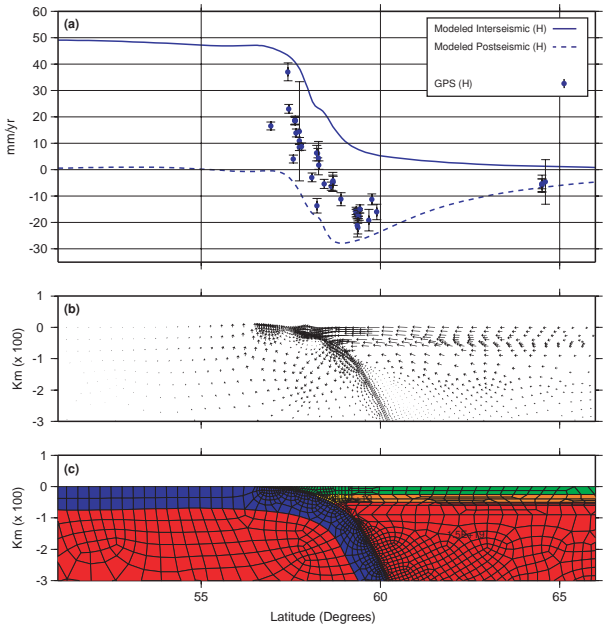


Latitude (Degrees)

Modeling deformation rate along B-B'

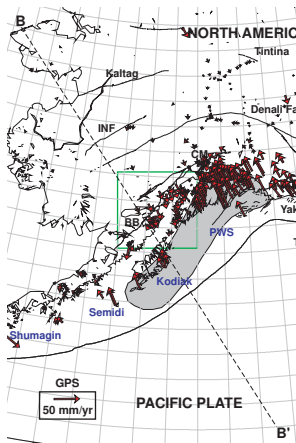


Postseismic signal
in 2001

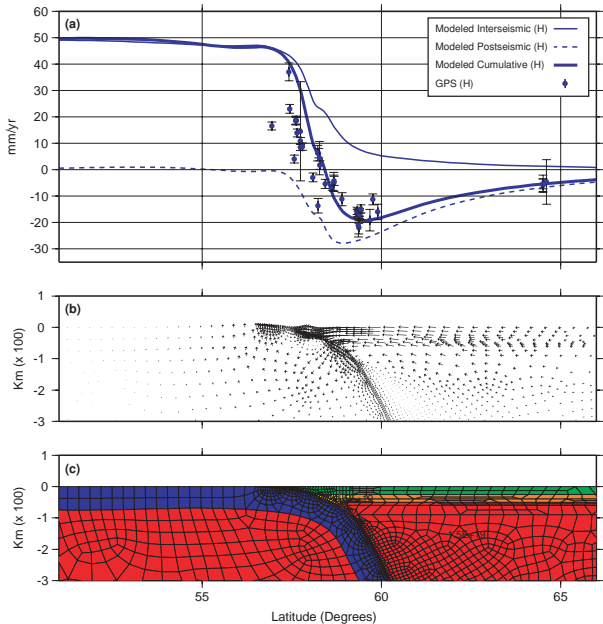


Latitude (Degrees)

Modeling deformation rate along B-B'

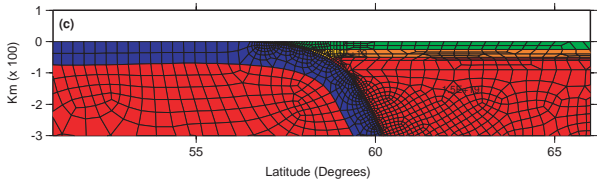
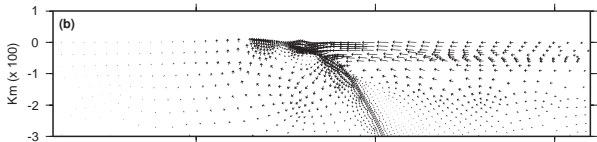
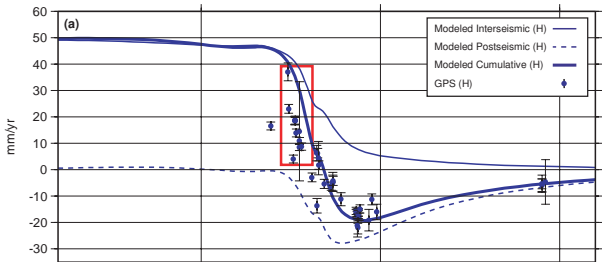
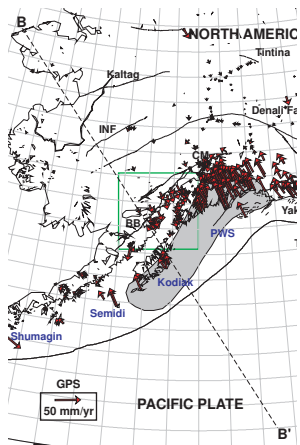


Cumulative signal
in 2001



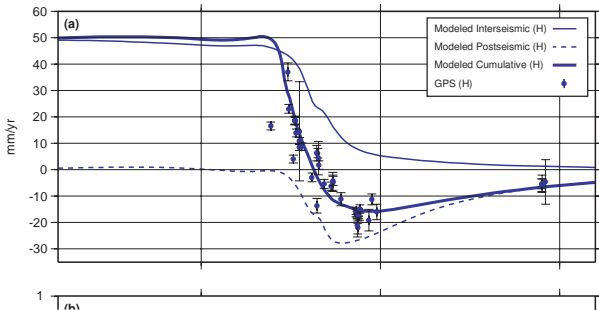
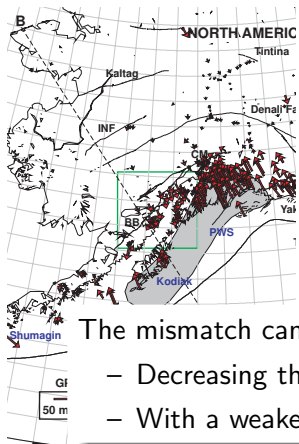
Latitude (Degrees)

Modeling deformation rate along B-B'



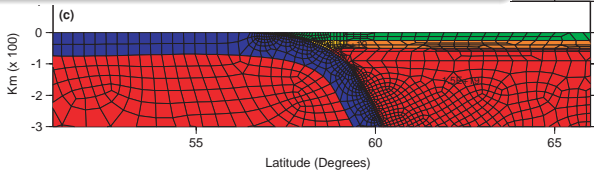
Latitude (Degrees)

Modeling deformation rate along B-B'



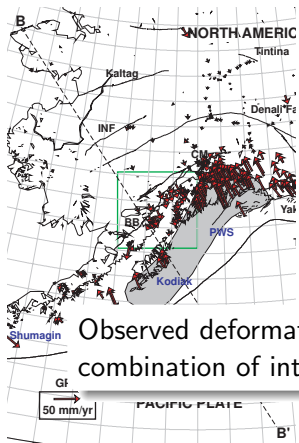
The mismatch can be fixed by either:

- Decreasing the locking depth
- With a weaker lower crust above the mantle wedge

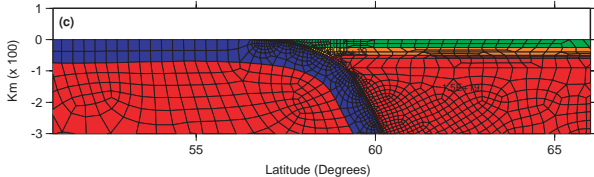
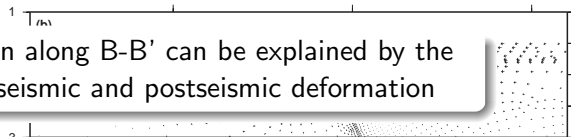
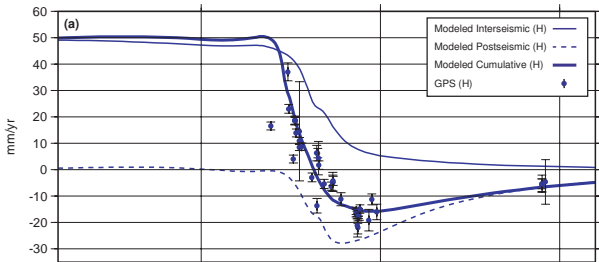


Latitude (Degrees)

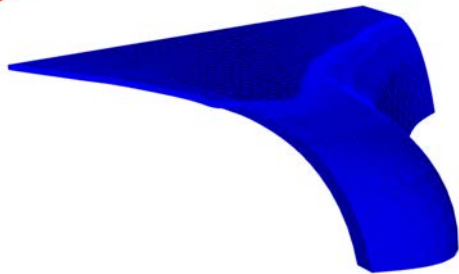
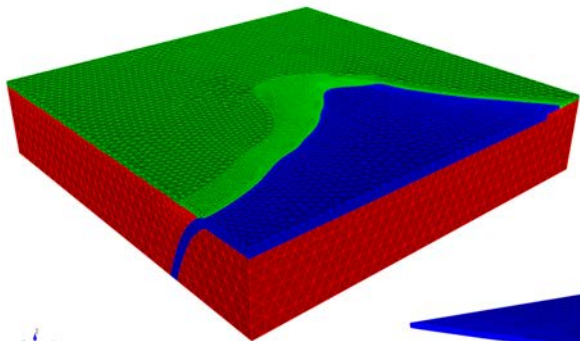
Modeling deformation rate along B-B'



Observed deformation along B-B' can be explained by the combination of interseismic and postseismic deformation

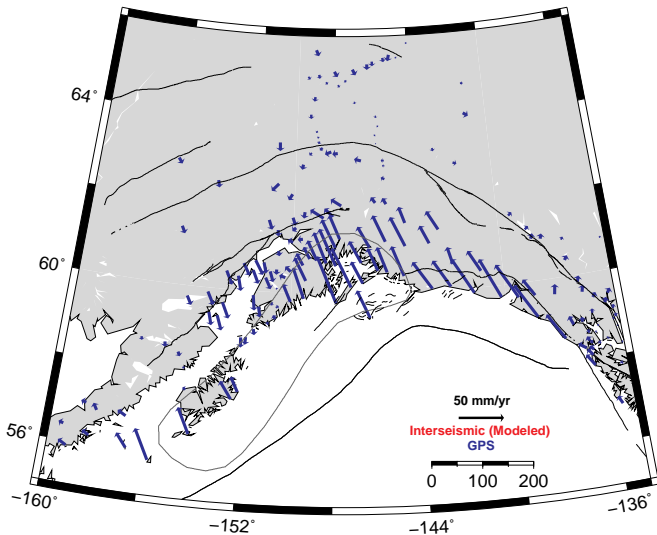


Modeling deformation in 3D: Discretized geometry



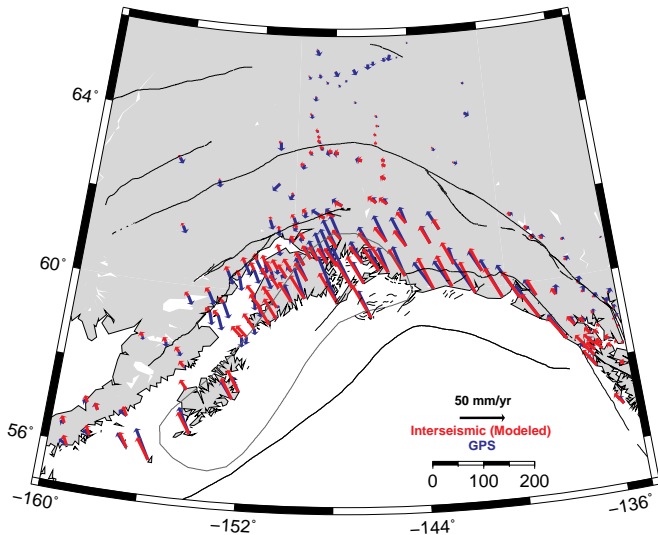
Interseismic deformation in 3D

GPS velocities

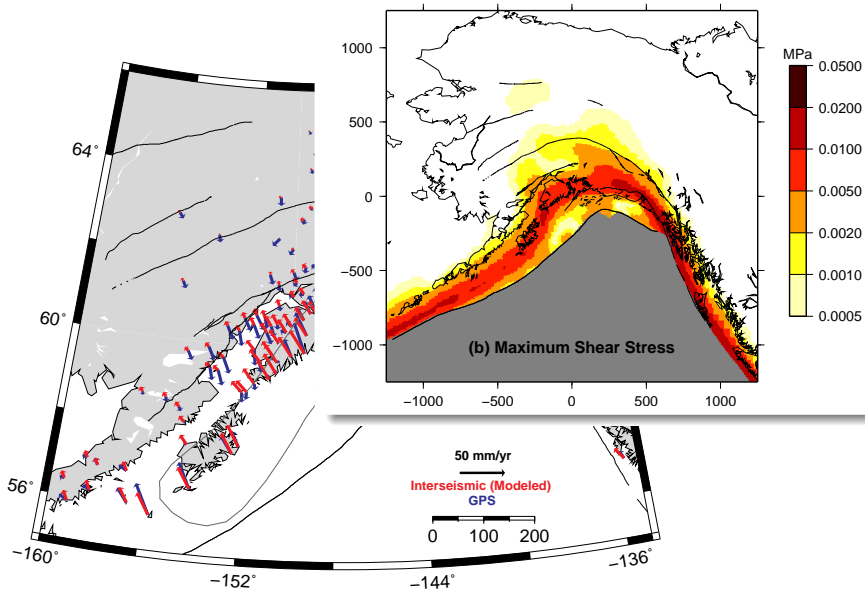


Interseismic deformation in 3D

GPS and modeled velocities

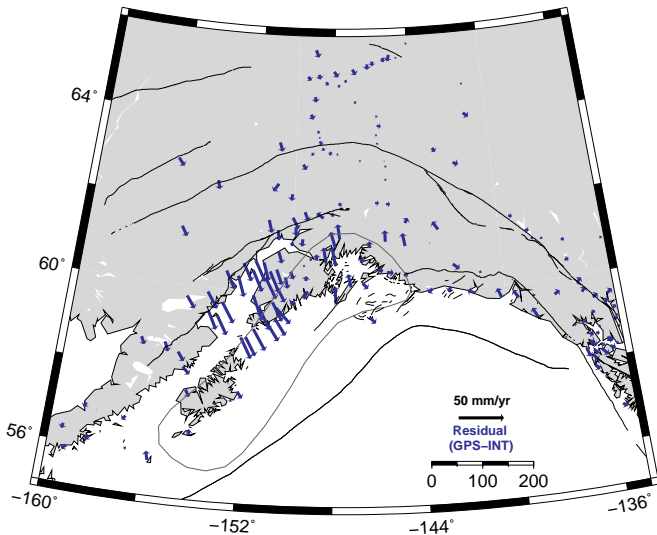


Interseismic deformation in 3D



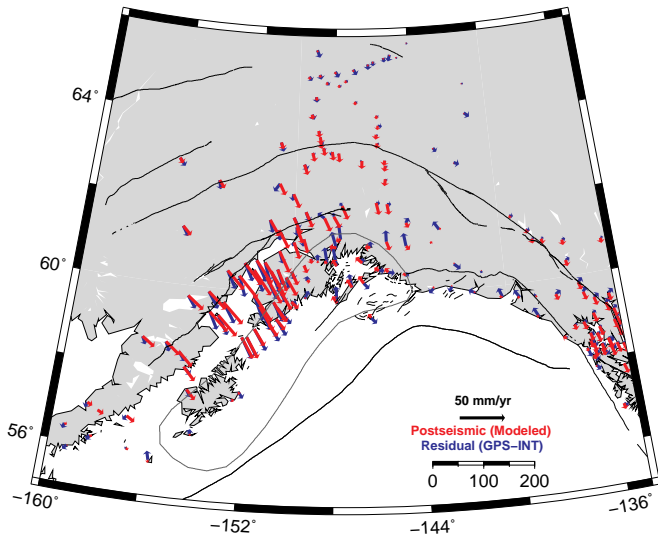
Residual velocities

GPS-Modeled velocities



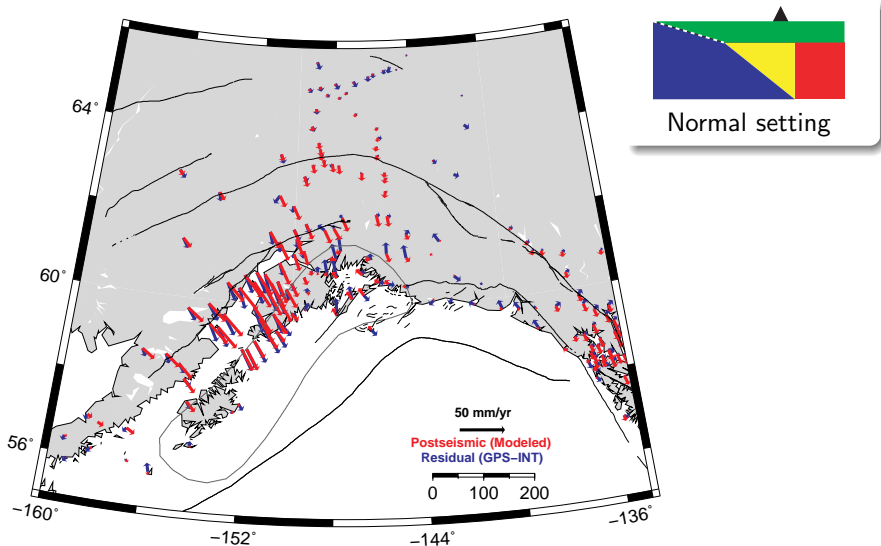
Postseismic deformation in 3D

Modeled postseismic velocities with viscous relaxation alone + a weaker lower crust above the mantle wedge



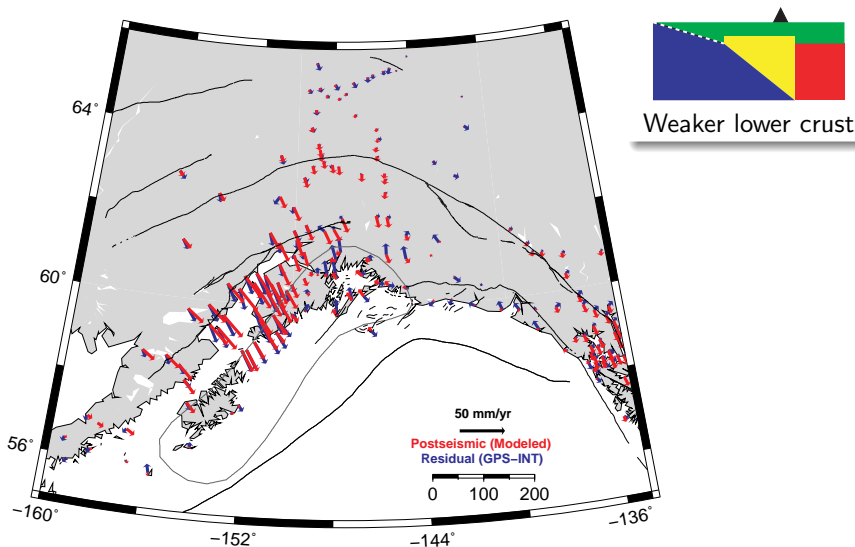
Postseismic deformation in 3D

Modeled postseismic velocities with viscous relaxation alone + a weaker lower crust above the mantle wedge



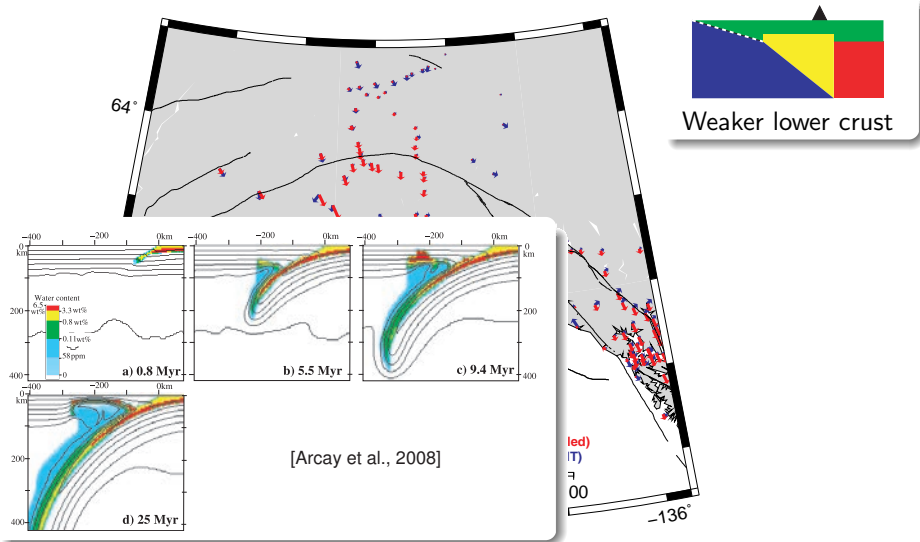
Postseismic deformation in 3D

Modeled postseismic velocities with viscous relaxation alone + a weaker lower crust above the mantle wedge



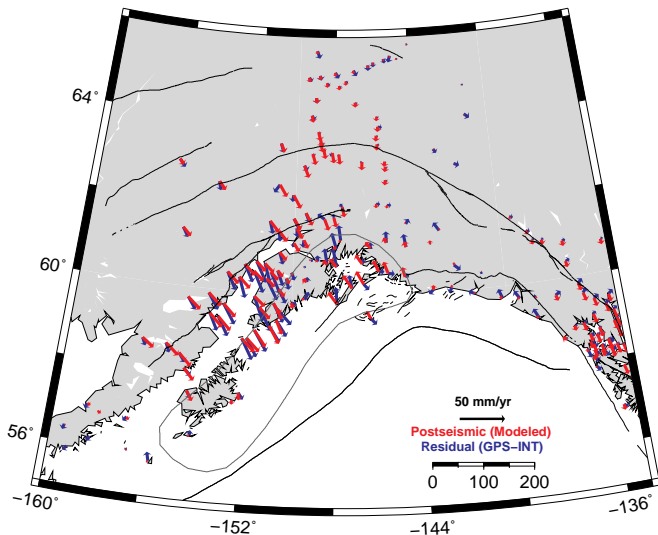
Postseismic deformation in 3D

Modeled postseismic velocities with viscous relaxation alone + a weaker lower crust above the mantle wedge



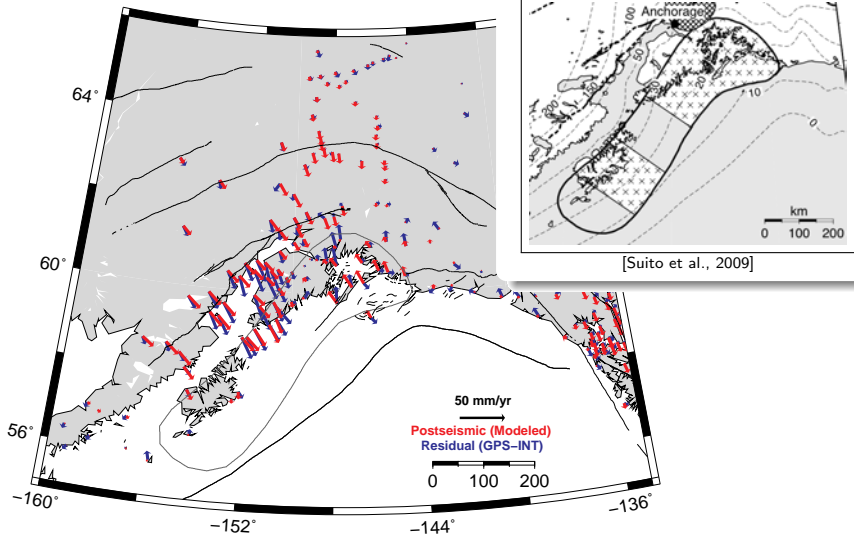
Postseismic deformation in 3D

Modeled postseismic velocities with viscous relaxation alone + shallower locking depth

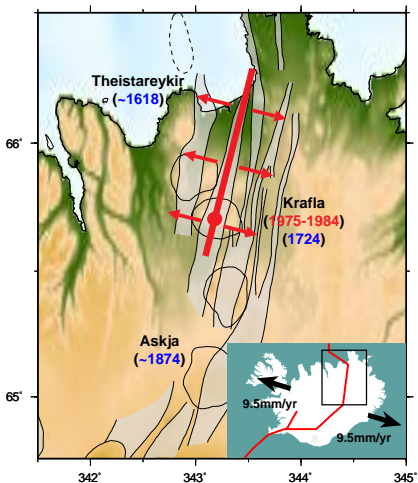


Postseismic deformation in 3D

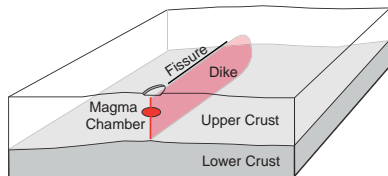
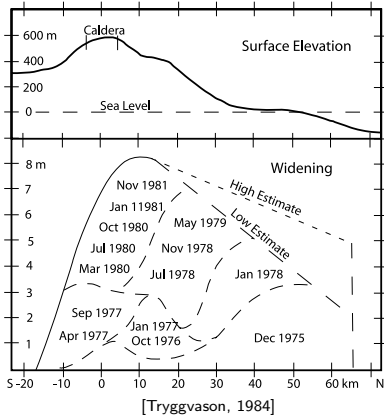
Modeled postseismic velocities with viscous relaxation alone + shallower locking depth



Iceland

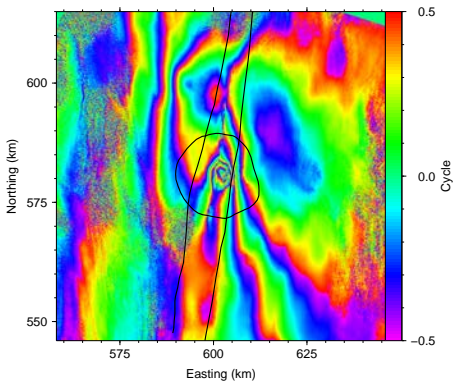
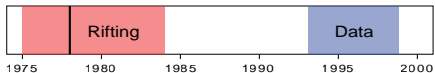


Northern Volcanic Zone and Krafla Volcanic System

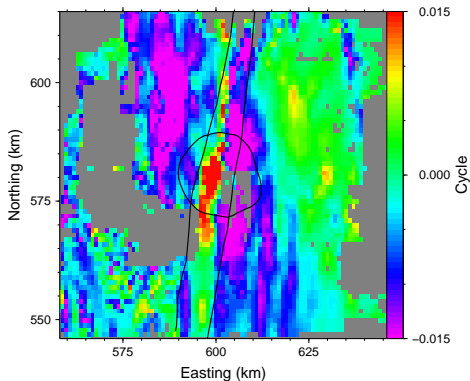


[Buck et al., 2006]

InSAR observations

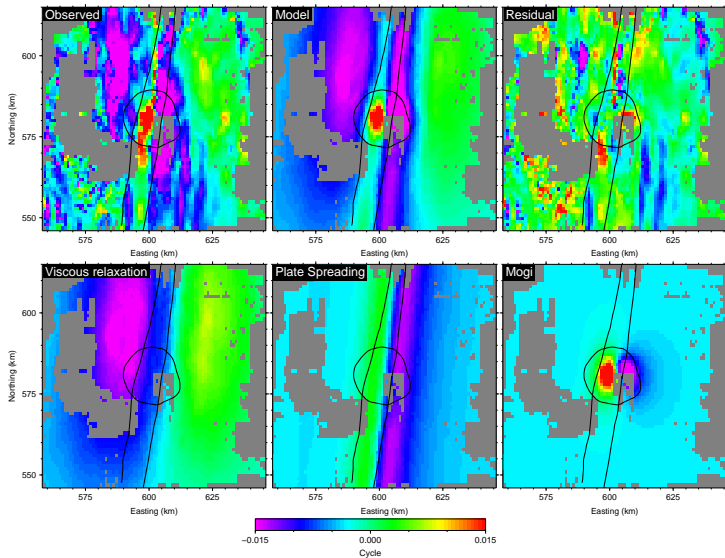


(Measures a component of displacement)



(Measures a component of strain)

Individual Contributions



Conclusions

- GPS velocities in southcentral Alaska can be explained by a combination of interseismic deformation and postseismic viscous relaxation following large recent earthquakes
- Best fitting model needed to match transient deformation requires postseismic viscous relaxation of mantle along with:
 - Shallow locking depth on part of the segment between the Kodiak and Kenai asperities, or
 - A weaker lower crust above the mantle wedge